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湘中包金山矿区花岗闪长斑岩的锆石 U-Pb 年龄、Hf-O 同位素组成及其地质意义



鲁玉龙^{1,3},彭建堂^{1,2},阳杰华²,李玉坤¹,陈宪佳¹,周 溪^{1,3},李干龙³

(1. 中南大学 地球科学与信息物理学院 有色金属成矿预测教育部重点实验室,长沙 410083;

2. 中国科学院 地球化学研究所 矿床地球化学国家重点实验室,贵阳 550002;

3. 湖南省有色地质勘查局二总队,湘潭 411102)

摘 要:包金山金矿床是近年在湘中地区探获的以金为主的多金属矿床,其成矿与印支晚期岩浆活动关系密切。 利用高精度 SIMS 锆石 U-Pb 定年方法得到矿区两条花岗闪长斑岩年龄分别为(225.1±1.5) Ma 和(223.3±1.4) Ma, 在误差范围内一致,属印支晚期大规模岩浆活动的产物;矿区两条花岗闪长斑岩的¹⁷⁶Hf/¹⁷⁷Hf 值均集中分布在 0.2825 附近,对应的 ε_{Hf}(*t*)值介于-5.9~3 之间,平均为-3.98,两阶段模式年龄 *T*_{DM2} 为 0.96~1.45 Ga,平均为 1.35 Ga; δ¹⁸O 值为 0.824%~0.973%; Hf-O 同位素研究表明花岗闪长斑岩主要为中元古代下地壳岩石重熔形成,且有部分 幔源物质参与成岩作用;与紫云山花岗岩体对比分析,表明两者为同源岩浆演化的产物,且间接指示包金山金矿 成矿时间为 227~223 Ma;花岗闪长斑岩形成于多板块汇聚的动力学背景,为印支地块向北挤压和太平洋板块向西 俯冲共同作用的结果。

关键词:SIMS 锆石 U-Pb 年龄;Hf-O 同位素示踪;花岗闪长斑岩;包金山矿床;湘中 文章编号:1004-0609(2017)-07-1441-14 中图分类号:P588.121;P597.3 文献标志码:A

随着一系列典型印支期成矿年龄的确认,如荷花 坪锡多金属矿((224±1.9) Ma)^[1]、云头界钨钼多金属矿 ((216.8±7.5) Ma)^[2]、锡田锡铅锌多金属矿((212.8±3.0) Ma)^[3],华南印支期的成矿作用越来越值得重视。在湘 中地区,金属(多金属)矿床多分布于印支期花岗岩体 周边,如白马山杂岩体附近的古台山石英脉型金锑矿 ^[4]和高家坳微细浸染型金矿床^[5–6];天龙山岩体附近的 大新破碎蚀变岩夹石英脉型金矿^[7];芙蓉岩体附近的 廖家坪金锑钨矿;紫云山岩体附近的包金山金矿等; 其成矿与印支期的岩浆活动有着密切的关系。

白马山-龙山-紫云山隆起带为湘中重要的金多 金属成矿带,其上的金属矿床主要赋存于受岩浆上侵 穹窿作用形成的放射状或环状断裂中^[7-11]。包金山金 矿位于该成矿带的东段(见图 1(b)),为典型的受紫云 山穹窿环状断裂和地层联合控制的矿床,以石英脉型 为主,其成矿流体属岩浆期后热液,与矿区内的花岗 闪长斑岩有着密切联系^[12-19]。但是以往地质工作对于 矿区内花岗闪长斑岩的研究还处于空白,严重制约了 对该区成矿作用的认识。 本文作者在野外实地调查的基础上,通过高精度 SIMS 锆石 U-Pb 测年及锆石原位 Hf-O 同位素分 析,精确构建区内花岗闪长斑岩脉的年代学格架,深 入探讨了花岗闪长斑岩的源区特征及其形成的动力学 背景,不仅为研究紫云山地区成岩成矿作用奠定了基 础,而且有助于进一步研究华南地区早中生代构造-岩浆演化过程。

1 矿区地质概况

包金山金矿床位于湘中盆地的东缘,紫云山岩体 北侧 3 km 处。在大地构造位置上,该区处在扬子地块 与华夏地块的对接带(见图 1(a)),长期以来经历了复杂 的构造岩浆活动,并伴有大规模的金属成矿活动。矿 区出露的地层为元古代板溪群马底驿组浅变质岩系, 赋矿地层为第二段倾向北的钙质板岩夹钙质条带状板 岩及灰岩透镜体。矿区主构造线方向为近东西向(见图 2),以断裂构造为主,近东西向断裂为紫云山穹窿环

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通信作者:彭建堂,教授,博士;电话:13787787177;E-mail:jtpeng@126.com



图1 湘中区域地质略图

Fig. 1 Regional geological map of central Hunan Province: 1—Cretaceous-Paleogene; 2—Carboniferous- Triassic; 3—Devonian; 4—Ordovician-Silurian; 5—Sinian- Cambrian; 6—Neoproterozoic; 7—Yanshanian granite; 8—Late Indosinian biotite (two-mica) granite; 9—Late Indosinian porphyroid monzonitic granite; 10—Caledonian granite; 11—Major deep fracture; 12—Geological boundary; 13—Unconformable boundary line; 14—Ore deposit(ore spot); 15—Study area



图 2 包金山矿区地质略图

Fig. 2 Regional geological map of Baojinshan gold deposit: 1—Xikuangshan formation, Devonian group; 2—Shetianqiao formation, Devonian group; 3—Qiziqiao formation, Devonian group; 4—Tiaomajian formation, Devonian group; 5—Third member of Madiyi formation, Banxi group; 6—Second member of Madiyi formation, Banxi group; 7—The late Indosinian granite; 8—Granodiorite –porphyry; 9—Ore belt; 10—Fault; 11—Sampling position and number

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;状断裂的一部分;近南北向、北东向及北西向破矿断 裂。区内花岗闪长斑岩分布广泛,但产出面积较小, 多为脉状,呈NW走向,倾向北东。本区主成矿期与 印支晚期中酸性岩浆的热液活动有关[16-18]。矿区内的 花岗闪长斑岩与矿体相互穿插,在其转弯处常有富矿 体存在,其对成矿具有重要的叠加改造作用,并为成 矿物质活化运移及矿体的进一步富集提供了热量和能 量,并提供了部分的成矿物质^[13,17-19]。围岩蚀变以绢 云母化、硅化、黄铁矿化与金成矿的关系最为密切。

本区金矿以石英脉型为主,主要有脉状和筒状两 类矿体。其中脉状矿体多呈群脉雁形排列,走向 NW, 倾向 SW, 倾角 45°~67°。该类矿体走向延伸较短, 一 般在 20 m 左右,倾向延深相对较长,矿体厚度 0.2~2.0 m。筒状矿体的走向和倾向与脉状矿体相近,但规模 相对较大,厚度几米到十几米,倾向延深可达百米。 金矿化主要位于石英脉与围岩的接触部位,近围岩的 石英脉内常见明金,金品位变化较大。

2 岩石岩相学特征

通过系统的野外地质观察及镜下显微鉴定分析 , 本次所采集的样品均为花岗闪长斑岩。岩石呈灰黑色, 斑状结构,块状构造(见图 3(a)和(b)),其浅色矿物为 长石、石英,暗色矿物以黑云母为主,斑晶含量约 20%~25%,主要由长石(15%~20%)和石英(5%~8%)组 成,基质含量约75%~80%,主要由它形粒状石英、长 石和黑云母。副矿物主要为锆石、磁铁矿等。长石主 要是斜长石、次为正长石。长石常沿边缘、粒间发生 不同程度的绢云母化,整体蚀变强烈,长石可呈残余 状产出,局部甚至全部蚀变为绢云母而保留原长石的 晶体外形(见图 3(d)、(e)和(f));石英分布零星,主要 呈椭圆状、圆状以斑晶的形式分布,集合体粒度一般 0.5~1.5 mm(见图 3(c)和(e))。

3 样品采集及测试分析

用于锆石测年研究的花岗闪长斑岩样品共有 2 件,样品 BJS-1 采自钻孔,样品 BJS-2 采自 10 中段的 44 线附近,样品均未发生风化。

3.1 锆石 U-Pb 年龄测定

锆石的挑选由河北省廊坊市诚信地质服务有限公 司完成,样品破碎至 0.3 mm 以下,用人工淘洗和电

磁选方法富集锆石,再在双目镜下精选锆石,未使用 任何化学药剂。将精选的锆石用无色透明环氧树脂固 定并进行抛光,使内部暴露,进行透射光和反射光下 照相,并用阴极发光扫描电子显微镜照相,以了解锆 石的内部结构,系统对比的基础上选出最理想的供分 析的锆石颗粒。锆石阴极发光(CL)内部结构照相在中 国科学院地质与地球物理研究所 Cameca 电子探针仪 器上完成,分析电压为 50 kV,电流为 15nA。锆石 U-Pb 定年在中国科学院地质与地球物理研究所离子 探针实验室的 Cameca IMS-1280 型二次离子质谱仪 (SIMS)上进行,详细的分析流程见文献[20]。锆石样 品的 Pb/U 比值用标准锆石 TEMORA 2(417 Ma)^[19]的 ln(²⁰⁶Pb/²³⁸U)与 ln(²³⁸U¹⁶O₂/²³⁸U)之间的线性关系校正 ^[22]; Th 和 U 含量用标准锆石 91500^[23]计算; 普通 Pb 用测量的 ²⁰⁴Pb 进行校正;用现代地壳的平均 Pb 同位 素组成^[24-25]作为普通 Pb 组成进行校正。单点分析的 同位素比值及年龄误差为 1σ(σ 为绝对误差), U-Pb 平 均年龄误差为 95%置信度。

3.2 锆石原位 Hf、O 同位素分析

锆石微区原位 Hf 同位素分析在中国地质大学(武 汉)地质过程与矿产资源国家重点实验室(GPMR)利 用激光剥蚀多接收杯等离子体质谱(LA-MC-ICP-MS) 完成。激光剥蚀系统为 GeoLas 2005 (Lambda Physik, 德国), 配备了信号平滑装置, 使用氦气作为载气, 并 引入少量氮气来提升元素的灵敏度^[26]。详细仪器操作 条件和分析方法参照文献[27]。锆石样品自身的 $\beta_{\rm Yb}$ 用于干扰校正^[28]; ¹⁷⁹Hf/¹⁷⁷Hf = 0.7325 和 ¹⁷³Yb/¹⁷¹Yb= 1.132685 ^[29]用于计算 Hf 和 Yb 的质量分馏系数 $\beta_{\rm Hf}$ 和 $\beta_{\rm Yb}$;¹⁷⁹Hf/¹⁷⁷Hf 和 ¹⁷³Yb/¹⁷¹Yb 的比值用于计算 Hf($\beta_{\rm Hf}$) 和 Yb (β_{Yb})的质量偏差;使用 ¹⁷⁶Yb/¹⁷³Yb=0.79639^[29] 来扣除¹⁷⁶Yb对¹⁷⁶Hf的同量异位干扰;使用 ¹⁷⁶Lu/¹⁷⁵Lu=0.02656 ^[30]来扣除干扰程度相对较小的 ¹⁷⁶Lu 对 ¹⁷⁶Hf 的同量异位干扰;采用 Yb 的质量分馏 系数 β_{vb}来校正 Lu 的质量分馏行为。分析数据的离线 处理(包括对样品和空白信号的选择、同位素质量分馏 校正)采用软件 ICP-MS Data Cal 完成^[31]。

锆石微区原位 O 同位素分析在中国科学院地质 与地球物理研究所离子探针实验室的 Cameca 进行。 将做过 SIMS 锆石 U-Pb 定年的样品靶再次磨去约 5 μm,以消除前期在 U-Pb 定年时造成的氧污染,详细 的分析流程参见文献[20]。仪器质量分馏校正采用 91500 锆石标准,其中 91500 标准锆石的 δ^{18} O= 0.99%^[32], 测量的 ¹⁸O/¹⁶O 比值通过 VSMOW 值 (¹⁸O/¹⁶O= 0.0020052)校正后,加上仪器质量分馏校正



图 3 花岗闪长岩岩相学特征

Fig. 3 Petrographies of granodiorite-porphyry: (a) Granodiorite-porphyry of BJS-1; (b) Granodiorite-porphyry of BJS-2; (c) Spherical quartz(BJS-1); (d) Spherical quartz and sericitization of plagioclase(BJS-2); (e) Spherical quartz and sericitization of plagioclase(BJS-1); (f) Sericitization of plagioclase(BJS-2). Qtz- Quartz; Pl- Plagioclase; Srt-Sericite

因子 IMF 即为该点的 δ^{18} O 值:

 $(\delta^{18}O)_M = (({}^{18}O/{}^{16}O)_M / 0.0020052 - 1) \times 100\%$; IMF= $(\delta^{18}O)_{M(\text{standard})} - (\delta^{18}O)_{VSMOW}$; $\delta^{18}O_{\text{sample}} = (\delta^{18}O)_M + \text{IMF}_{\circ}$

4 分析结果

4.1 锆石 U-Pb 年龄

供 U-Pb 同位素定年的 2 件花岗闪长斑岩样品中 的锆石呈灰色长柱状或短柱状,颗粒较大,长度 100~200 μm,长宽比一般为 1.5~4,晶体自形程度较 好,大部分具有清晰韵律环带结构(见图 4),并且锆石 的 Th/U 比值主要介于 0.10~0.90 之间(见表 1 和表 2), 具有岩浆锆石的特点^[33]。利用 Isoplot^[34]绘制锆石的谐 和曲线及谐和年龄的投影图(见图 5),数据点多分布在 谐和线上或在谐和线附近呈线状分布,表明所测的锆 石颗粒在形成后 U-Pb 同位素体系是封闭的,基本未 发生 U 或 Pb 的加入和丢失^[35-37]。

样品 BJS-1 进行了 25 个点的定年分析,其中有 4 个测点(8、15、17、18)年龄为离群值,其余 21 个测 点 ²⁰⁶Pb/²³⁸U 数据得到加权平均年龄(225.5±3.3) Ma(95% 置信度, MSWD=0.21),谐和年龄为 (225.1±1.5) Ma(95%置信度,MSWD=0.40)(见图 5(a))。



图 4 包金山金矿床花岗闪长斑岩锆石 CL 图像及分析点位图

Fig. 4 Zircon cathodeluminescence (CL) images and analytical point bitmaps of granodiorite -porphyry in Baojinshan gold deposit: (a) BJS-1; (b) BJS-2





Fig. 5 Zircon U-Pb concordia age diagrams of granodiorite -porphyry in Baojinshan gold deposit: (a) BJS-1; (b) BJS-2

样品 BJS-2 进行了 25 个点的定年分析,其中有 4 个测点(1、9、18、21)年龄为离群值,其余 21 个测点 ²⁰⁶Pb/²³⁸U 数据得到的加权平均年龄为(223.5±3.2) Ma(95% 置信度, MSWD=0.25),谐和年龄为 (223.3±1.4) Ma(95%置信度,MSWD=1.9)(见图 5(b))。

因此,包金山矿区两条花岗闪长斑岩的形成年龄 分别为(223.3±1.4) Ma 和(225.1±1.5) Ma,其形成时 间在误差范围内一致,属印支晚期的产物。

4.2 锆石的 Hf、O 同位素组成特征

对样品 BJS-1 的锆石进行微区原位 Hf 同位素分 析,初始¹⁷⁶Hf/¹⁷⁷Hf 比值为 0.282494~0.282729 之间, 平均为 0.282538,对应的 $\varepsilon_{\text{Hf}}(t)$ 为-5.3~3(见图 6(a)), 平均为-3.4,两阶段模式年龄 T_{DM2} 为 0.96~1.41 Ga, 平均为 1.32 Ga。锆石微区原位 O 同位素分析, δ^{18} O 为 0.824%~0.973%(见图 6(b)),平均为 0.907%。

对样品 BJS-2 的锆石进行微区原位 Hf 同位素分

析,初始¹⁷⁶Hf/¹⁷⁷Hf 比值为 0.282286~0.282592,平均 为 0.282498,对应的 $\varepsilon_{\text{Hf}}(t)$ 为-5.9~2.3(见图 6(a)),平均值-4.5,两阶段模式年龄 T_{DM2} 为 1.25~1.45 Ga,平均为 1.37 Ga。锆石微区原位 O 同位素分析, δ^{18} O 为 0.866%~0.948%(见图 6(b)),平均为 0.916%。

对花岗闪长斑岩样品 BJS-1 和 BJS-2 的锆石颗粒 (包括定年的锆石)进行了微区原位 Hf-O 同位素测定, 分析结果分别见表 3 和 4。

5 讨论

5.1 成岩年代及意义

本实验中采用高精度 SIMS 锆石 U-Pb 测年方法获得的包金山矿区花岗闪长斑岩年龄为(225.1±1.5) Ma

和 (223.3±1.4) Ma,在误差范围内一致。其与临近的 紫云山花岗岩体(锆石 SIMS U-Pb(227.0±2.2)~ (225.2±1.7) Ma^[38],锆石LA-ICP-MS U-Pb(222.5±1.0) Ma~(222.3±1.8) Ma^[39])、歇马花岗岩体(锆石 SHRIMP U-Pb(218±3) Ma^[40];锆石 SHRIMP U-Pb(214.1±5.9) Ma^[41]) 以及与区域上的南岳花岗岩体(锆石 LA-ICP-MSU-Pb(215.5±1.5) Ma^[42]、沩山花岗岩体(黑 云母 Rb-Sr(227.0±13) Ma 和(221.9±5.8) Ma^[43])、大 神山花岗岩体(锆石 LA-ICP-MS(224.3±1.0) Ma^[44])、 白马山花岗岩体(锆石 LA-ICP-MS U-Pb(224.3±2.4) Ma、(221.4±4.0) Ma 和(226.6±4.1) Ma^[45])的成岩时限 基本一致,表明矿区内花岗闪长斑岩属于印支晚期 (230~200 Ma)大规模岩浆活动的组成部分^[37,46]。该类 花岗闪长斑岩在区域上广泛发育,在紫云山岩体主体 花岗岩中也偶见出露,其与紫云山岩体应属同期不同

表1 包金山矿区花岗闪长斑岩样品 BJS-1 锆石 U-Pb(SIMS)分析结果

Table 1Zircon U-Pb (SIMS) analytical results of No. BJS-1 sample from granodiorite-porphyry in Baojinshan gold deposit

Analytical	Th/	U/	Th/II	207 Dh /206 Dh	1	207 Db /235 I I	1	206 Db /238 t I	1	²⁰⁷ Pb/ ²⁰⁶ Pb/	1-	²⁰⁷ Pb/ ²³⁵ U/	1	206Pb/238U/	1
position	10^{-6}	10^{-6}	TH/U	P0/ P0	10	P0/ U	10	PD/ U	10	Ma	10	Ma	10	Ma	10
BJS-1-1	831	3180	0.26	0.05058	0.00067	0.24791	0.00499	0.0355	0.00054	221.7	30.4	224.9	4.1	225.2	3.3
BJS-1-2	476	1894	0.25	0.04973	0.00044	0.24219	0.00421	0.0353	0.00053	182.4	20.3	220.2	3.4	223.8	3.3
BJS-1-3	599	2093	0.29	0.05107	0.00116	0.24806	0.00678	0.0352	0.00054	243.8	51.6	225.0	5.5	223.2	3.3
BJS-1-4	365	1552	0.23	0.05106	0.00057	0.24428	0.00459	0.0347	0.00052	243.4	25.6	221.9	3.7	219.9	3.3
BJS-1-5	227	633	0.36	0.04954	0.00082	0.23814	0.00558	0.0349	0.00058	173.3	38.1	216.9	4.6	220.9	3.6
BJS-1-6	350	1720	0.20	0.05055	0.00052	0.24092	0.00476	0.0346	0.00059	220.4	23.4	219.2	3.9	219.1	3.6
BJS-1-7	1078	3257	0.33	0.05054	0.00064	0.24523	0.00482	0.0352	0.00053	220.0	29.0	222.7	3.9	222.9	3.3
BJS-1-8	718	1812	0.40	0.05378	0.00380	0.21302	0.01831	0.0287	0.00141	361.7	152.1	196.1	15.4	182.6	8.8
BJS-1-9	253	2621	0.10	0.05070	0.00036	0.25128	0.00422	0.0359	0.00055	227.4	16.2	227.6	3.4	227.6	3.4
BJS-1-10	614	2826	0.22	0.05005	0.00033	0.25039	0.00415	0.0363	0.00055	197.4	15.0	226.9	3.4	229.7	3.4
BJS-1-11	351	1712	0.21	0.05056	0.00050	0.25200	0.00462	0.0361	0.00056	220.8	22.6	228.2	3.7	228.9	3.5
BJS-1-12	1099	1228	0.90	0.05077	0.00051	0.24715	0.00449	0.0353	0.00053	230.5	23.2	224.3	3.7	223.7	3.3
BJS-1-13	368	1707	0.22	0.05030	0.00059	0.24696	0.00473	0.0356	0.00054	208.7	27.1	224.1	3.9	225.6	3.3
BJS-1-14	619	942	0.66	0.05003	0.00070	0.25335	0.00567	0.0367	0.00064	196.4	32.2	229.3	4.6	232.5	4.0
BJS-1-15	465	1074	0.43	0.05269	0.00139	0.24480	0.00894	0.0337	0.00085	315.3	58.8	222.3	7.3	213.7	5.3
BJS-1-16	282	1205	0.23	0.05039	0.00049	0.25078	0.00448	0.0361	0.00054	213.2	22.3	227.2	3.6	228.6	3.4
BJS-1-17	1439	6454	0.22	0.05035	0.00834	0.24595	0.04092	0.0354	0.00056	211.1	344.4	223.3	33.9	224.4	3.5
BJS-1-18	265	726	0.36	0.05014	0.00098	0.26432	0.00778	0.0382	0.00084	201.6	44.6	238.1	6.3	241.9	5.2
BJS-1-19	504	1892	0.27	0.05101	0.00043	0.24985	0.00436	0.0355	0.00054	241.2	19.3	226.5	3.5	225.0	3.4
BJS-1-20	356	1088	0.33	0.05070	0.00045	0.24511	0.00430	0.0351	0.00053	227.1	20.4	222.6	3.5	222.2	3.3
BJS-1-21	491	1420	0.35	0.04986	0.00073	0.24744	0.00520	0.0360	0.00054	188.4	33.7	224.5	4.2	228.0	3.4
BJS-1-22	485	1705	0.28	0.05096	0.00037	0.25003	0.00424	0.0356	0.00054	239.1	16.7	226.6	3.4	225.4	3.4
BJS-1-23	607	2376	0.26	0.05079	0.00033	0.24860	0.00417	0.0355	0.00055	231.4	14.7	225.4	3.4	224.9	3.4
BJS-1-24	811	2285	0.35	0.05106	0.00032	0.24788	0.00410	0.0352	0.00054	243.5	14.5	224.9	3.3	223.1	3.4
BJS-1-25	856	2234	0.38	0.05054	0.00032	0.25544	0.00416	0.0367	0.00055	220.0	14.6	231.0	3.4	232.1	3.4

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表 2 包金山矿区花岗闪长斑岩样品 BJS-2 锆石 U-Pb(SIMS)分析结果

Table 2 Zircon U-FU (Shvis) analytical results of No. DJS-2 sample from granoutome-polphyty in Daofinshan gold

Analytical	Th/	U/	701 (7)	207 21 206 21		20751 (2357)		206-51 (238-1		²⁰⁷ Pb/ ²⁰⁶ Pb/		²⁰⁷ Pb/ ²³⁵ U/		²⁰⁶ Pb/ ²³⁸ U/	
position	10^{-6}	10^{-6}	ιh/U	²⁰⁷ Pb/ ²⁰⁰ Pb	1σ	²⁰⁷ Pb/ ²³³ U	lσ	²⁰⁰ Pb/ ²³⁸ U	lσ	Ma	lσ	Ма	lσ	Ма	lσ
BJS-2-1	439	1820	0.24	0.05093	0.00096	0.23342	0.00643	0.0332	0.00067	237.7	42.7	213.0	5.3	210.8	4.2
BJS-2-2	627	1898	0.33	0.05057	0.00038	0.24263	0.00409	0.0348	0.00052	221.3	17.4	220.6	3.3	220.5	3.3
BJS-2-3	221	733	0.30	0.05045	0.00063	0.24296	0.00475	0.0349	0.00053	215.8	28.5	220.8	3.9	221.3	3.3
BJS-2-4	356	1644	0.22	0.05129	0.00050	0.24774	0.00445	0.0350	0.00053	254.0	22.2	224.7	3.6	221.9	3.3
BJS-2-5	367	1327	0.28	0.05010	0.00046	0.23829	0.00420	0.0345	0.00052	199.5	21.1	217.0	3.4	218.6	3.2
BJS-2-6	379	1251	0.30	0.05033	0.00067	0.23906	0.00480	0.0344	0.00052	210.3	30.5	217.7	3.9	218.3	3.2
BJS-2-7	1151	3021	0.38	0.05067	0.00031	0.25195	0.00409	0.0361	0.00054	226.0	14.3	228.2	3.3	228.4	3.4
BJS-2-8	445	1533	0.29	0.05060	0.00042	0.24998	0.00430	0.0358	0.00054	222.8	19.0	226.6	3.5	226.9	3.4
BJS-2-9	865	2170	0.40	0.04725	0.00563	0.22243	0.02671	0.0341	0.00053	61.6	261.6	203.9	22.4	216.4	3.3
BJS-2-10	1073	5127	0.21	0.05027	0.00023	0.25012	0.00393	0.0361	0.00054	207.5	10.6	226.7	3.2	228.5	3.4
BJS-2-11	592	2282	0.26	0.05046	0.00044	0.24904	0.00433	0.0358	0.00054	216.2	20.1	225.8	3.5	226.7	3.3
BJS-2-12	866	2629	0.32	0.05003	0.00043	0.25121	0.00437	0.0364	0.00055	196.4	19.7	227.6	3.6	230.6	3.4
BJS-2-13	420	1163	0.36	0.05060	0.00049	0.24443	0.00437	0.0350	0.00053	222.7	22.2	222.0	3.6	222.0	3.3
BJS-2-14	534	1817	0.29	0.05044	0.00039	0.24202	0.00416	0.0348	0.00053	215.1	18.0	220.1	3.4	220.5	3.3
BJS-2-15	693	2393	0.28	0.05067	0.00042	0.24769	0.00424	0.0355	0.00053	225.9	18.9	224.7	3.5	224.6	3.3
BJS-2-16	390	1460	0.26	0.05070	0.00043	0.25158	0.00435	0.0360	0.00054	227.3	19.7	227.9	3.5	227.9	3.4
BJS-2-17	656	1621	0.40	0.05012	0.00053	0.24522	0.00452	0.0355	0.00053	200.4	24.5	222.7	3.7	224.8	3.3
BJS-2-18	586	1096	0.54	0.05037	0.00253	0.25214	0.01325	0.0363	0.00055	212.1	112.6	228.3	10.8	229.9	3.4
BJS-2-19	647	2207	0.29	0.05074	0.00036	0.24500	0.00413	0.0350	0.00054	229.1	16.4	222.5	3.4	221.9	3.3
BJS-2-20	392	1155	0.34	0.05212	0.00161	0.24839	0.00854	0.0346	0.00052	290.7	69.1	225.3	7.0	219.0	3.2
BJS-2-21	1128	3768	0.30	0.05044	0.00049	0.23904	0.00427	0.0344	0.00052	215.4	22.3	217.6	3.5	217.8	3.2
BJS-2-22	567	1750	0.32	0.05076	0.00042	0.24419	0.00420	0.0349	0.00052	230.0	19.2	221.8	3.4	221.1	3.3
BJS-2-23	823	1834	0.45	0.05074	0.00040	0.24955	0.00422	0.0357	0.00054	229.2	18.0	226.2	3.4	225.9	3.3
BJS-2-24	1105	2808	0.39	0.04854	0.00074	0.23974	0.00512	0.0358	0.00054	125.8	35.3	218.2	4.2	226.9	3.3
BJS-2-25	509	1495	0.34	0.05006	0.00054	0.24214	0.00447	0.0351	0.00053	197.6	24.7	220.2	3.7	222.3	3.3

Test method: SIMS; Test unit: Beijing Nano SIMS Lab, Institute of Geology and Geophysics, Chinese Academy of Sciences.





Fig. 6 Hf isotope compositions (a) and O isotope compositions (b) for granodiorite-porphyry of Baojinshan gold deposit

1	Λ	Λ	Q
- 1	4	4	0

Analytical position	Age/Ma	¹⁷⁶ Hf/ ¹⁷⁷ Hf	1σ	¹⁷⁶ Yb/ ¹⁷⁷ Hf	1σ	¹⁷⁶ Lu/ ¹⁷⁷ Hf	1σ	$arepsilon_{ m Hf}/t$	1σ	T _{DM2} / Ma	$\delta^{18}\mathrm{O}/\%$	2σ
BJS-1-1	225.2	0.282495	0.000009	0.022826	0.000162	0.000948	0.000007	-5.0	0.6	1403	0.910	0.32
BJS-1-2	223.8	0.282507	0.000009	0.019328	0.000199	0.000791	0.000007	-4.6	0.6	1379	0.937	0.20
BJS-1-3	223.2	0.282500	0.000009	0.027605	0.000672	0.001132	0.000026	-4.9	0.6	1397	0.906	0.43
BJS-1-4	219.9	0.282521	0.000010	0.020444	0.000324	0.000841	0.000013	-4.2	0.6	1354	0.896	0.28
BJS-1-5	220.9	0.282532	0.000010	0.024769	0.000606	0.001023	0.000024	-3.8	0.6	1334	0.962	0.20
BJS-1-6	219.1	0.282558	0.000011	0.024579	0.001142	0.001018	0.000046	-2.9	0.7	1283	0.859	0.27
BJS-1-7	222.9	0.282539	0.000009	0.027212	0.000138	0.001101	0.000003	-3.5	0.6	1321	0.879	0.34
BJS-1-8	225.1	0.282532	0.000012	0.027327	0.000424	0.001107	0.000016	-3.7	0.7	1333	0.885	0.25
BJS-1-9	227.6	0.282544	0.000011	0.024410	0.000400	0.000994	0.000017	-3.2	0.7	1307	0.964	0.2
BJS-1-10	229.7	0.282540	0.000010	0.021444	0.000811	0.000912	0.000033	-3.3	0.6	1313	0.893	0.41
BJS-1-11	228.9	0.282541	0.000010	0.023967	0.000636	0.000977	0.000023	-3.3	0.6	1313	0.879	0.3
BJS-1-12	223.7	0.282729	0.000012	0.072467	0.003883	0.002795	0.000146	3.0	0.7	961	0.824	0.20
BJS-1-13	225.6	0.282535	0.000010	0.023423	0.001205	0.000976	0.000047	-3.6	0.6	1326	0.903	0.2
BJS-1-14	232.5	0.282495	0.000009	0.023008	0.000551	0.000980	0.000021	-4.8	0.6	1401	0.851	0.28
BJS-1-15	213.7	0.282494	0.000010	0.027248	0.000643	0.001096	0.000024	-5.3	0.6	1411	0.886	0.33
BJS-1-16	228.6	0.282604	0.000011	0.037229	0.001713	0.001463	0.000065	-1.1	0.7	1193	0.912	0.28
3JS-1-17	224.4	0.282528	0.000010	0.033934	0.001877	0.001351	0.000072	-3.9	0.7	1343	0.925	0.31
BJS-1-18	241.9	0.282567	0.000010	0.034057	0.001626	0.001343	0.000061	-2.1	0.6	1259	0.853	0.23
BJS-1-19	225.0	0.282528	0.000009	0.022449	0.000215	0.000920	0.000008	-3.8	0.6	1339	0.969	0.33
BJS-1-20	222.2	0.282515	0.000010	0.019815	0.000697	0.000794	0.000027	-4.3	0.6	1364	0.973	0.32
BJS-1-21	228	0.282558	0.000010	0.026795	0.001427	0.001094	0.000054	-2.7	0.6	1281	0.861	0.42
BJS-1-22	225.4	0.282510	0.000010	0.033483	0.000362	0.001349	0.000014	-4.5	0.6	1378	0.957	0.30
BJS-1-23	224.9	0.282526	0.000010	0.024462	0.000193	0.001001	0.000007	-3.9	0.6	1343	0.913	0.39
BJS-1-24	223.1	0.282510	0.000010	0.022526	0.000835	0.000912	0.000032	-4.5	0.6	1374	0.959	0.41
BJS-1-25	232.1	0.282543	0.000012	0.025537	0.000606	0.001014	0.000024	-3.2	0.7	1308	0.921	0.33

阶段的产物,形成时间应略晚于后者。

包金山金矿产于紫云山岩体入侵形成的环状断裂 内,那么其成矿时间不会早于紫云山岩体的侵位时间; 而花岗闪长斑岩脉穿插包金山矿脉,表明包金山成矿 不会晚于花岗闪长斑岩的形成时间。因此,包金山成 矿时间应介于紫云山岩体((227.0±2.2)~(225.2±1.7) Ma^[38])和包金山花岗闪长斑岩((225.1±1.5) Ma 和 (223.3±1.4) Ma)侵位时间之间,即 227~223 Ma。

5.2 岩浆源区及意义

对花岗闪长斑岩锆石原位 Hf-O 同位素数据统 计(见图 6),其 Hf-O 同位素的组成基本相似,基本 可以重叠。初始¹⁷⁶Hf/¹⁷⁷Hf 值均集中分布在 0.2825 附近,对应的 $\varepsilon_{\rm Hf}(t)$ 值在-5.9~3.0之间,平均为-3.98; 两阶段模式年龄 $T_{\rm DM2}$ 为0.96~1.45Ga,平均为1.35Ga; δ^{18} O值为0.824~0.973%之间,均大于地幔 δ^{18} O值((0.53±0.03)%)^[47];在 $\varepsilon_{\rm Hf}(t)$ -U/PbAge图解中(见图7(a)),样品点主要落于下地壳区域,表明花岗闪长斑岩为中元古代下地壳岩石重熔形成。而在 δ^{18} O- $\varepsilon_{\rm Hf}(t)$ 图解中(见图7(b)),Hf-O同位素完全落在以大容山过铝质花岗岩为代表的地壳端元区域以外,表明源区有幔源岩浆的加入,且加入地幔物质的比例至少可达20%^[48-49]。样品BJS-1的12号测点,初始¹⁷⁶Hf/¹⁷⁷Hf比值较大,对应其年龄计算的 $\varepsilon_{\rm Hf}(t)$ 值为3.0(见表3和4),其可能为幔源物质参与的结果。与临近的紫云山岩体相比,两者的Hf同位素组成基 第27卷第7期

表 4 包金山金矿床花岗闪长斑岩样品 BJS-2 锆石 Hf-O 同位素分析结果

Table 4Zirons Hf-O isotopic analytical results of No. BJS-2 sample from the granodiorite-porphyry in the Baojinshan golddeposit

Analytical position	Age/ Ma	¹⁷⁶ Hf/ ¹⁷⁷ Hf	1σ	¹⁷⁶ Yb/ ¹⁷⁷ Hf	1σ	¹⁷⁶ Lu/ ¹⁷⁷ Hf	1σ	$arepsilon_{\mathrm{Hf}}$	1σ	T _{DM2} / Ma	$\delta^{18}\mathrm{O}/\%$	2σ
BJS-2-1	223.3	0.282487	0.000011	0.016702	0.000381	0.000700	0.000016	-3.8	0.6	1337	9.19	0.30
BJS-2-2	220.5	0.282511	0.000010	0.030778	0.001362	0.001219	0.000051	-3.8	0.7	1335	9.03	0.34
BJS-2-3	221.3	0.282544	0.000011	0.025340	0.001246	0.001009	0.000047	-4.8	0.7	1392	9.31	0.25
BJS-2-4	221.9	0.282494	0.000010	0.023034	0.000061	0.000956	0.000002	-5.9	0.7	1449	8.85	0.23
BJS-2-5	218.6	0.282539	0.000010	0.015681	0.000263	0.000663	0.000010	-4.8	0.6	1389	8.86	0.22
BJS-2-6	218.3	0.282537	0.000011	0.019808	0.000212	0.000817	0.000009	-4.8	0.6	1389	9.16	0.25
BJS-2-7	228.4	0.282471	0.000011	0.032466	0.000760	0.001312	0.000029	-5.2	0.7	1419	8.66	0.29
BJS-2-8	226.9	0.282592	0.000012	0.021845	0.000197	0.000917	0.000009	-5.2	0.6	1418	9.26	0.21
BJS-2-9	223.3	0.282539	0.000011	0.027870	0.000758	0.001145	0.000029	-5.7	0.7	1439	9.30	0.21
BJS-2-10	228.5	0.282544	0.000013	0.034676	0.000578	0.001407	0.000026	-3.9	0.7	1344	9.12	0.28
BJS-2-11	226.7	0.282545	0.00001	0.023038	0.000109	0.000938	0.000004	-5.2	0.7	1419	9.38	0.29
BJS-2-12	230.6	0.282514	0.000011	0.022905	0.000110	0.000941	0.000004	-4.5	0.6	1380	9.13	0.23
BJS-2-13	222.0	0.282433	0.000010	0.019175	0.000180	0.000793	0.000007	-4.5	0.7	1375	9.48	0.24
BJS-2-14	220.5	0.282525	0.000010	0.021314	0.000102	0.000873	0.000004	-4.3	0.6	1363	9.31	0.33
BJS-2-15	224.6	0.282509	0.000010	0.022285	0.000183	0.000915	0.000007	-4.0	0.6	1346	9.40	0.27
BJS-2-16	227.9	0.282425	0.000009	0.015069	0.000386	0.000637	0.000015	-4.1	0.7	1354	9.39	0.25
BJS-2-17	224.8	0.282411	0.000010	0.039812	0.001206	0.001568	0.000045	-4.2	0.7	1362	9.21	0.20
BJS-2-18	229.9	0.282453	0.000014	0.023493	0.000288	0.000955	0.000012	-4.2	0.7	1366	9.11	0.30
BJS-2-19	221.9	0.282525	0.000012	0.020411	0.000192	0.000841	0.000007	-4.7	0.7	1385	9.23	0.35
BJS-2-20	217.8	0.282556	0.000010	0.029296	0.001034	0.001186	0.000041	-5.6	0.7	1432	8.83	0.39
BJS-2-21	219.0	0.282524	0.000012	0.032111	0.001688	0.001288	0.000064	-2.3	0.7	1252	9.35	0.27
BJS-2-22	221.1	0.282286	0.000010	0.022139	0.000119	0.000908	0.000005	-3.9	0.6	1343	9.25	0.28
BJS-2-23	225.9	0.282538	0.000011	0.013333	0.000223	0.000569	0.000009	-5.1	0.6	1411	9.11	0.24
BJS-2-24	226.9	0.282529	0.000010	0.027524	0.000113	0.001133	0.000004	-3.6	0.6	1327	9.15	0.25
BJS-2-25	222.3	0.282428	0.000012	0.021802	0.000321	0.000903	0.000012	-3.8	0.6	1338	8.90	0.25

Zirons Hf isotopic: LA-MC-ICP-MS, State Key Laboratory of Geological Processes and Mineral Resources, China University of Geosciences (Wuhan) Airons Hf Isotopic; Beijing nano SIMS Lab, Institute of Geology and Geophysics, Chinese Academy of Sciences.

本相同,分布范围和峰值基本一致(见图 6(a));两者的 O 同位素组成也基本相似,不过花岗闪长斑岩 δ^{18} O 值 的峰值略高一些(见图 6(b)),这可能是由于花岗闪长 斑岩后期发生蚀变的结果。在图7中,花岗闪长斑岩 与紫云山花岗岩落于同一区域,表明两者为同源的 特征。

因此,本文作者认为区内花岗闪长斑岩由中元古 代下地壳岩石熔融,与幔源岩浆形成的壳、幔混合岩 浆源区演化而来,其与紫云山花岗岩体应属于同源岩 浆不同演化阶段的产物。

5.3 动力学背景探讨

在华南地块内印支晚期花岗岩的逐渐确认,尤其 是一系列印支晚期 A 型花岗岩的发现,如湖南的锡 田岩体、浙江的大爽岩体、江西蔡江岩体和福建高溪 岩体等^[39,50-55],对于华南印支晚期花岗岩形成于地壳 伸展构造体制的认识已经成为共识。但对于华南印支 期花岗岩形成的动力学背景还存在分歧,部分学 者^[44,53-60]认为印支运动期间(约 260 Ma),印支地块向 北挤压,与华南地块发生碰撞,导致地层加厚,之后 10~20 Ma 发生热-应力松弛作用,进入伸展阶段,



图 7 包金山金矿床花岗闪长斑岩 $\varepsilon_{Hf}(t)$ -U/Pb 图解(a)和 $\delta^{18}O-\varepsilon_{Hf}(t)$ 图解(b)

Fig. 7 Plot of $\varepsilon_{\text{Hf}}(t)$ –U/Pb (a) and δ^{18} O– $\varepsilon_{\text{Hf}}(t)$ (b) for granodiorite-porphyry of Baojinshan gold deposit

地壳减压熔融,进而形成花岗岩。LI 等^[61]认为,在中 二叠世(约 280 Ma)太平洋板块西缘转为活动大陆边 缘,开始向华南地块俯冲,到三叠世华南地块发生造 山运动,之后进入伸展--减薄构造背景,地壳熔融形 成花岗岩^[62]。从印支期华南地块内的构造变形来分 析,其以 NE-NNE 向的褶皱及逆冲推覆构造为主,单 纯印支地块引起的南北向的挤压应力很难解释这些构 造的形成。这些构造的形成应与同期东部太平洋板块 的俯冲有着密切的关系。

所以,本文作者认为华南印支期应该属于多板块 汇聚的动力学背景,其构造变形和同期花岗岩的形成 应为印支地块向北挤压和太平洋板块向西俯冲共同作 用的结果。

6 结论

1) 包金山矿区两条花岗闪长斑岩锆石 SIMS
 U-Pb 年龄分别为(225.1±1.5) Ma 和(223.3±1.4) Ma,
 在误差范围内一致,其与印支晚期(230~200 Ma)大规

模的岩浆活动时限一致,并间接的指示了包金山成矿 时间为 227~223 Ma,属典型印支晚期成矿。

2)花岗闪长斑岩由中元古代下地壳岩石重熔与
 部分幔源岩浆组成的壳、幔混合岩浆源区演化而来,
 其与紫云山岩体为同源不同演化阶段的产物。

 3)花岗闪长斑岩形成于多板块汇聚的动力学背景,为印支地块向北挤压和太平洋板块向西俯冲共同 作用的结果。

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Zircon U-Pb ages and Hf-O isotopes of granodiorite -porphyry in Baojinshan mining area and their geological significance

LU Yu-long^{1, 3}, PENG Jian-tang^{1, 2}, YANG Jie-hua², LI Yu-kun¹, CHEN Xian-jia¹, ZHOU Xi^{1, 3}, LI Gan-long³

(1. Key Laboratory of Non-ferrous Metals Metallogenic Prediction, Ministry of Education,

School of Geosciences and Info-physics, Central South University, Changsha 410083, China;

2. State Key Laboratory of Ore Deposit Geochemistry, Institute of Geochemistry,

Chinese Academy of Sciences, Guiyang 550002, China;

3. The Second Corps of Nonferrous Geological Prospecting Bureau in Hunan Province,

Geological Exploration Institute, Xiangtan 411102, China)

Abstract: Baojinshan gold deposit is Au-dominated polymetallic deposit discovered recently in central Hunan, its mineralization is considered to be closely related to the magmatic activities in late Indosinian. High accuracy SIMS zircon U-Pb age determination was used in this study to obtain the ages of two granodiorite-porphyry dykes at (225.1 ± 1.5) Ma and (223.3 ± 1.4) Ma, indicating that this deposit was the product of extensive magma activity in late Indosinian. The ¹⁷⁶Hf/¹⁷⁷Hf values of two granodiorite-porphyry dykes intensively distributed around 0.2825, with the corresponding $\varepsilon_{\rm Hf}(t)$ ranged from -5.9 to 3, averaging -3.98. The T_{DM2} model age was estimated from 0.96 GPa to 1.45 Ga, averaging 1.35 Ga; the δ^{18} O value of granodiorite-porphyry ranged from 0.824% to 0.973%. Hf-O isotope analysis shows that the granodiorite-porphyry is mainly formed by the rock remelting of Mesoproterozoic lower crust, and mantle-sourced materials are considered to participate the diagenesis. Contrastive analysis on the granodiorite-porphyry in Baojinshan with Ziyunshan granite suggests that both of them are the products of comagmatic evolution, which indirectly indicates that the metallogenetic time of Baojinshan gold deposit ranges in 227–223 Ma. The granodiorite-porphyry is formed under the multi-plates convergent dynamics background, and affected by the northward squeeze of Indochina Block and westward subduction of Pacific Plate.

Key words: SIMS U-Pb zircon dating; Hf-O isotopes tracing; granodiorite-porphyry; Baojinshan gold deposit; central Hunan Province

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Corresponding author: PENG Jian-tang; Tel: +86-13787787177; E-mail: jtpeng@126.com

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