



In situ U-Pb dating of calcite indicates a Miocene Sb-Pb mineralization event in the Sanjiang base metal metallogenic belt, SW China

Xiangyuan Sheng^{a,b}, Yongyong Tang^{b,*}, Xianwu Bi^b, Ruizhong Hu^b, Leiluo Xu^b, Juan Li^{b,c}, Yanwen Tang^b

^a College of Resources and Environmental Engineering, Mianyang Normal University, Mianyang 621000, Sichuan, China

^b State Key Laboratory of Ore Deposit Geochemistry, Institute of Geochemistry, Chinese Academy of Sciences, Guiyang 550081, China

^c University of Chinese Academy of Sciences, Beijing 100049, China

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ABSTRACT

The Sanjiang base metal metallogenic belt (SMB) in southwestern China is characteristic of composite mineralization during the India-Asia continental collision. However, the geodynamic settings of the composite processes of this belt are poorly understood. The Lanuoma deposit, located in the Changdu Basin in the central segment of the SMB, is a Zn-Pb-Sb polymetallic deposit. The cross-cutting relationships and paragenetic mineral associations suggest the superimposition of a late Sb-Pb mineralization event over an early Zn-dominant mineralization event. Calcite is the main gangue mineral, intergrown with late Sb-Pb minerals (e.g., boulangierite, zinkenite, sorbyite and plagionite). In situ U-Pb dating by laser ablation-inductively coupled plasma mass spectrometry (LA-ICPMS) was successfully carried out on calcite from this deposit, and restricts the Sb-Pb mineralization to 19.7 ± 1.6 Ma, with a mean square weighted deviation (MSWD) of 1.2. This age is approximately synchronous with that of strike-slip shearing-related leucogranites in the marginal orogens; together with previous data, this result suggests that the contemporaneous magmatism significantly contributed to the ore-forming fluids associated with the Sb-Pb mineralization. Overprinting of the Miocene Sb-Pb event in a trans-tensional regime on the Oligocene transpression-driven Zn mineralization resulted in the characteristic Zn-Pb-Sb assemblage at Lanuoma. Moreover, this study indicates that the SMB holds great exploration potential for Sb-(Pb) resources.

1. Introduction

The Sanjiang base metal metallogenic belt (SMB; Fig. 1) is one of the most important producers of base metals (e.g., Pb, Zn, Cu, Sb and Co) in China (Hou et al., 2008; Liu et al., 2017b; Song et al., 2017; Bi et al., 2019; Leach and Song, 2019). This belt contains several economically significant large to giant Pb-Zn polymetallic deposits (e.g., Huoshaoyun, Jinding, Baiyangping, Mohailaheng and Chaqapacha) and numerous small-medium deposits, which are mainly distributed in the Tianshuihai, Tuotuohe, Yushu, Changdu and Lanping areas. The deposits are hosted in Mesozoic limestones and sandstones and are spatially controlled by Cenozoic fault systems associated with the India-Asia continental collision (Hou et al., 2008; He et al., 2009; Hou and Zhang, 2015). Although many studies have considered the SMB to be a Mississippi Valley-type (MVT) belt (Leach et al., 2017; Song et al., 2017; Leach and Song, 2019), characteristic polymetallic assemblages (e.g., Pb-Zn-(Ag-Cu-Co,

Sb-Hg and Cd-Tl)) distinguish it from the typical polymetallic assemblages (e.g., Pb-Zn-(Ag) of MVT deposits). Most recent studies have suggested that the composite mineralization of basinal brine, meteoric water and magmatic fluid can be used to interpret the SMB (Tao et al., 2011; Wang et al., 2011b; Deng et al., 2014; Bi et al., 2019). However, the detailed geodynamic settings and processes remain poorly constrained.

Geochronological advances show a peak of 38–28 Ma for the Pb-Zn mineralization in the SMB (Wang et al., 2011a; Zou et al., 2015; Liu et al., 2016; Feng et al., 2017; Yalikhun et al., 2018; Bi et al., 2019; Zhang et al., 2019), corresponding to the late-collisional stage of tectonic transformation (from compression to extension) of the India-Asia collision (Hou and Cook, 2009). Because of possible inclusions (e.g., carbonate and clay minerals) in sphalerite, Bradley et al. (2004) and Kesler et al. (2004) expressed their concern about the reliability of sphalerite Rb-Sr isotope isochron ages for MVT Pb-Zn deposits. As most of the

* Corresponding author.

E-mail address: tangyongyong@vip.gyig.ac.cn (Y. Tang).

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sphalerite Rb-Sr isotope dates for the SMB Pb-Zn deposits are consistent (with an age cluster of 33–28 Ma) and agree with the Sm-Nd isotope dates of the intergrown calcites (Wang et al., 2011a; Zou et al., 2015; Feng et al., 2017; Bi et al., 2019), these sphalerite Rb-Sr dating data likely represent the timing of the early Pb-Zn ore formation. Very few data are available to constrain the late base metal mineralization timing for two possible reasons: 1) a lack of minerals suitable for conventional radioisotopic dating and 2) the influence by earlier ores.

The Lanuoma deposit, located in the Changdu Basin in the central SMB, has a metal assemblage of Pb-Zn-Sb. Field observations show two episodes of hydrothermal mineralization: an early Zn-dominated mineralization episode and a late Sb-Pb mineralization episode, represented by an early sphalerite-pyrite assemblage and a late Sb-Pb-sulfosalt-sphalerite-calcite-orpiment-realgar assemblage, respectively (Sheng et al., 2019). Sphalerite Rb-Sr dating produces an isochron age of 30 Ma (Bi et al., 2019), which likely represents the age of the early Zn mineralization episode. The results of this study provide a constraint on the timing of the late Sb-Pb mineralization episode based on in situ laser ablation-inductively coupled plasma mass spectrometry (LA-ICPMS) U-Pb dating of calcite, which indicates that a Miocene Sb-Pb mineralization event occurred in the transtensional setting of the India-Asia collision. Moreover, a model of fluid mixing between shallow fluid (e.g., basal brine) and magmatic fluid is proposed for Sb-Pb mineralization of the Lanuoma deposit, which provides significant insights into the metallogenesis within the SMB.

2. Geological setting and ore deposit characteristics

The SMB, restricted between the Jinshajiang and Nujiang sutures within the North Qiangtang and Changdu-Simao blocks (Fig. 1; Pan

et al., 1997; Zhong, 1998; Metcalfe, 2002), is in the eastern India-Asia collision zone along the northern to northeastern margin of the Tibetan Plateau. The belt extends over a thousand kilometres and is characterized by a series of large-scale Cenozoic thrust and strike-slip fault systems, Tertiary foreland basins and Cenozoic potassic-ultrapotassic igneous rocks (Hou et al., 2008). The Tertiary basins, developed on a Precambrian to Palaeozoic metamorphic basement (Zhai and Cong, 1993; Yin et al., 1999; Wang et al., 2000), underwent late Triassic rifting and Jurassic to Cretaceous depression (Yin and Harrison, 2000; Pan et al., 2003; Spurlin et al., 2005), filled with shallow marine clastic and carbonate rocks and marine-terrestrial to terrestrial red clastic rocks (Mu et al., 1999; Zhong et al., 2000; Liao and Chen, 2005). Associated with the crustal contraction caused by the India-Asia collision in the early Eocene (prior to 51 Ma; Spurlin et al., 2005), the Mesozoic depression basins evolved into isolated foreland basins filled with lake-facies siliciclastic rock sequences (Mu et al., 1999; Spurlin et al., 2005). These basins underwent four stages of deformation in the Cenozoic: Palaeocene-Eocene compression (~65–42 Ma), Eocene-Oligocene transpression (~40–26 Ma), early mid-Miocene transtension (~25–18 Ma) and late Neocene extension (<18 Ma; Wang et al., 2001; Hou et al., 2006). The south-north-trending and/or northwest-southeast-trending thrust faults defined the major tectonic framework of the SMB and remained episodically active because of the crustal shortening caused by the India-Asia collision during the Eocene (50–37 Ma; Hou et al., 2006). These thrusts were overprinted by strike-slip faults that initially occurred at 42 Ma and lasted to 23 Ma with early-stage sinistral and late-stage dextral movements (Tapponnier et al., 1990; Hou et al., 2003; Spurlin et al., 2005; Liu et al., 2006).

Cenozoic magmatism in the Sanjiang region clustered into three episodes: Palaeocene to early Eocene, middle Eocene to late Oligocene

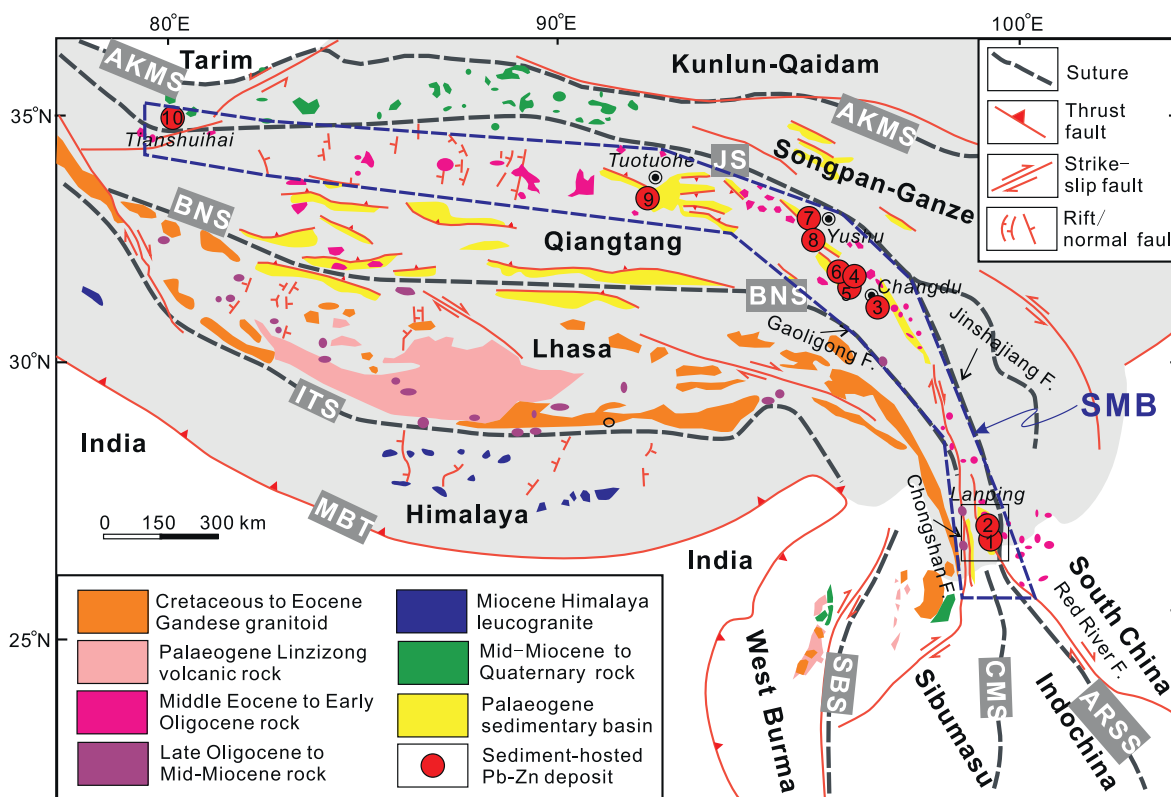


Fig. 1. Simplified geological map showing the location of the SMB along the northern to northeastern margin of the Tibetan Plateau (modified from Chung et al., 2005; Wang et al., 2010). The numbers in the solid red circles indicate major Pb-Zn deposits: 1, Jinding; 2, Baiyangping; 3, Lanuoma; 4, Zhaoafayong; 5, Lalongla; 6, Jiamoshan; 7, Dongmohazhuzha; 8, Mohailaheng; 9, Chaqupacha and 10, Huoshayun. Abbreviations: AKMS, Animaqing-Kunlun-Muztagh suture; JS, Jinshajiang suture; ARSS, Ailaoshan-Red River shear zone; BNS, Bangong-Nujiang suture; ITS, Indus-Tsangpo suture; MBT, Main Boundary thrust and CMS, Changning-Menglian suture. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

and late Miocene to Holocene (Fig. 1). Evidence of Palaeocene to early Eocene magmatic activity is mainly distributed in the Tengchong block, with ages from 62 to 47 Ma, resulting from Neotethyan oceanic slab subduction (Xu et al., 2012; Chen et al., 2014). The middle Eocene to early Oligocene magmatism comprises the following: 1) late Eocene intraplate magmatism, expressed as mafic dykes (42–39 Ma) in the Gaoligongshan shear zone and eastern margin of the Tengchong block and generated by partial melting of metasomatized lithospheric mantle in response to asthenospheric upwelling following the breakoff of the subducted Neotethyan oceanic slab (Xu et al., 2008); and 2) middle Eocene to early Oligocene potassic-ultrapotassic igneous rocks along the Jinshajiang-Ailaoshan tectonic belt, which display increasingly younger ages from the northern segment (41–36 Ma) to the southern segment (~35 Ma) and likely formed as a magmatic expression of lower lithospheric mantle delamination (Hou et al., 2003; Lu et al., 2012; Deng et al., 2014). The late Miocene to Holocene magmatism is mainly Pleistocene-Holocene volcanic rocks in the northern and southern segments of the Sanjiang domain, which are genetically ascribed to decompression melting of the asthenosphere associated with subduction-induced metasomatism (Wang et al., 2001; Zhou et al., 2012; Flower et al., 2013).

Above the metamorphic base of the Jitang and Youxi Groups, the Changdu Basin was filled with Triassic to Palaeogene marine carbonate and terrigenous siliciclastic sedimentary rocks (Fig. 2a). Indosinian magmatism occurred as granodiorite and biotite granites in the southern basin and as rhyolites and dacites in the northern basin (Tao et al., 2014; Bi et al., 2019). Cenozoic felsic rocks (e.g., monzogranite, quartz monzonite, granite porphyry and syenite porphyry) are mainly exposed around Yulong in the northern basin and locally in the southern basin (Hou et al., 2003; Guo et al., 2006; Jiang et al., 2006). After the India-Asia continental collision, the tectonic activities of the Changdu Basin were characterized by thrust and strike-slip faults during the Eocene to Oligocene, which controlled the distribution of major ore deposits in the basin (Tang et al., 2006). Pb and Zn deposits, e.g., Jiamoshan, Zhaofayong, Lalongla, Lanuoma and Cuona, are mainly localized within the Lancangjiang thrust system, which constitutes a northwest-southeast-trending metallogenic belt across Leiwuqi and Zuogong Counties.

The Lanuoma deposit is located in the central Changdu Basin. The

strata exposed in the mining area contain the Upper Triassic Jiapila, Bolila, Adula and Duogaila Formations and Quaternary sediments (Fig. 2b). The Duogaila, Adula and Jiapila Formations consist mainly of greywacke, mudstone, shales and sandstones, and the Bolila Formation consists of clastic rocks and microcrystalline limestone. The deposit consists of two orebodies (I and II) hosted by the limestone of the Bolila Formation (Fig. 2b). Orebody I is predominantly composed of Zn, whereas orebody II is predominantly composed of Sb (inferred resource: 253 hundred tons; Feng et al., 2006) and Pb. Two-stage mineralizations consisting of early Zn-dominated mineralization and late Sb-Pb mineralization exist in the Lanuoma deposit (Fig. 3a, b; Sheng et al., 2019). The primary ore minerals are mainly sphalerites, S-Sb-Pb minerals (Sb-Pb sulfosalts), pyrite, orpiment and realgar (Fig. 3c, d; Sheng et al., 2019). The Sb-Pb sulfosalts mainly consist of boulangerite, zinkenite, sorbyite and plagiomite (Sheng et al., 2019). Generally, sphalerite formed earlier than the Sb-Pb minerals and coexist with minor pyrite (Fig. 3b, d, e); the Sb-Pb minerals are mainly intergrown with calcite (Fig. 3b, e), and minor sphalerite formed in association with Sb-Pb minerals (Fig. 3f).

3. Sample preparation and analysis methods

In situ U-Pb dating was performed on calcite intergrown with S-Sb-Pb minerals (Fig. 4). Samples were polished into thin sections before examination by optical microscopy to confirm generations. U and Pb isotope ratios were measured using an LA-ICPMS instrument equipped with an Agilent 7900 quadruple mass spectrometer and a GeolASpRO 193 nm ArF excimer laser, at the State Key Laboratory of Ore Deposit Geochemistry, Institute of Geochemistry, Chinese Academy of Sciences. The parameters for ablation were as follows: spot size, 120 μm ; repetition rate, 10 Hz; and energy density, 8 J/cm². The samples and standards were ablated for 3–5 pulses to clean the surfaces before formal ablation. Before entering the torch, the aerosols were carried by pure helium (450 ml/min) and mixed with argon via a T-connector. Pure nitrogen (~3.0 ml/min) was added to the helium carrier gas to increase the sensitivity via a Y junction before entering the ICP (Hu et al., 2008). The standard-sample bracketing approach was used to measure ²⁰²Hg, ²⁰⁴Pb, ²⁰⁶Pb, ²⁰⁷Pb, ²⁰⁸Pb, ²³²U and ²³⁸U. Each analysis involved 20 s for background

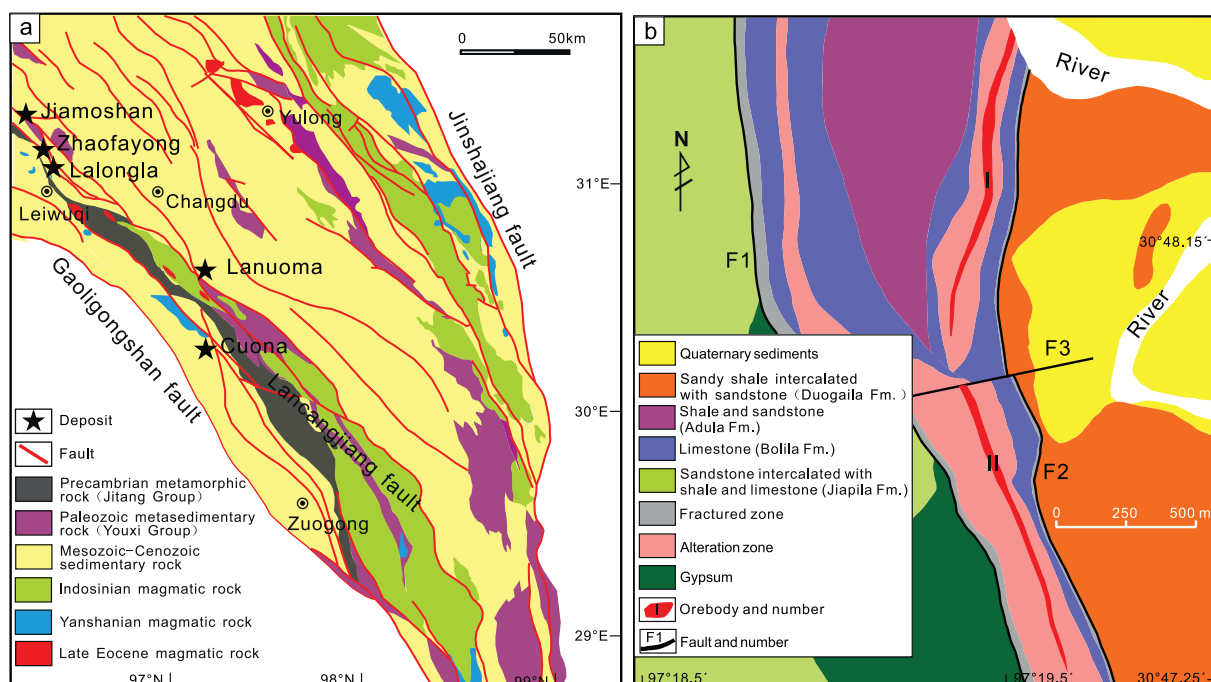


Fig. 2. Geological maps of the Changdu Basin (a; modified from Pan et al., 2004) and the Lanuoma deposit (b; modified from Feng et al., 2006).

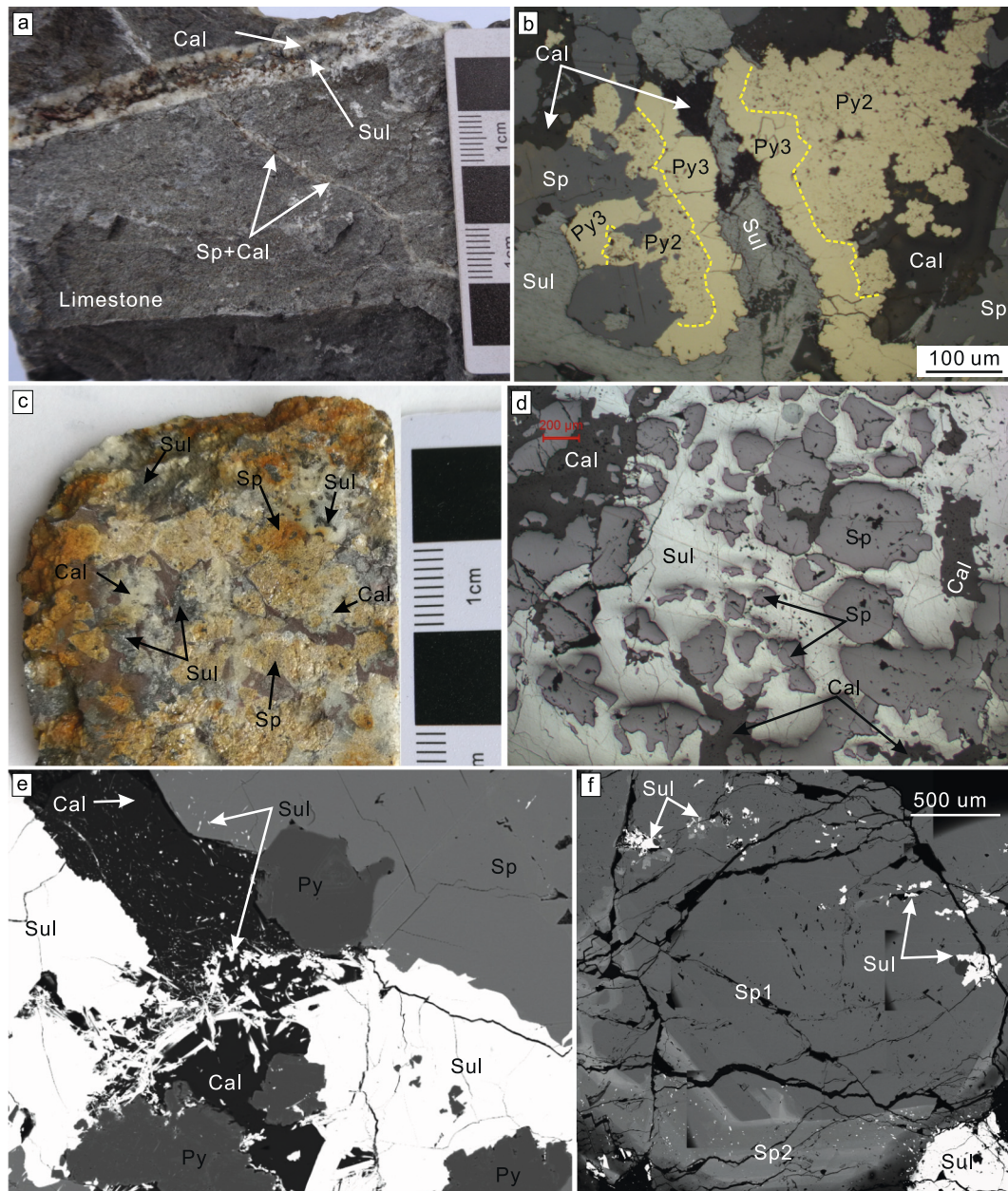


Fig. 3. Mineral association shows two mineralization stages of the Lanuoma deposit. a and b (b; modified from Sheng et al., 2019) reveal that late S-Sb-Pb minerals-calcite veins cut early sphalerite-calcite veins. The hand specimen in c shows the major primary ore minerals. d and e show that S-Sb-Pb minerals coexist with calcite but cut sphalerites. f shows that Sp1 is overgrown by Sp2, and Sp2 is associated with S-Sb-Pb minerals. Abbreviations: Py, pyrite; Sp, sphalerite; Sul, S-Sb-Pb minerals; Cal, calcite.

acquisition, 30 s for data acquisition and 40 s for the elimination of memory effects. NIST 614 was used to correct the $^{207}\text{Pb}/^{206}\text{Pb}$ ratios, and a matrix-matched carbonate reference (WC-1) was used for fractionation correction of the $^{238}\text{U}/^{206}\text{Pb}$ ratios. Data reduction was performed using the program ICPMSDataCal (Liu et al., 2008). Other details can be found in Roberts and Walker (2016) and Roberts et al. (2017).

4. Results

In total, 90 analyses of calcite associated with S-Sb-Pb minerals were conducted, which yielded Pb concentrations of 0.09 to 9.85 ppm, ^{232}Th values of 0.04 to 1.97 ppm and ^{238}U values of 0.10 to 2.58 ppm. The $^{238}\text{U}/^{206}\text{Pb}$ ratios ranged from 0.05 to 35.08, and the $^{207}\text{Pb}/^{206}\text{Pb}$ ratios ranged from 0.730 to 0.843 (Suppl. A1), revealing a lower intercept age of 19.7 ± 1.6 Ma with an initial $^{207}\text{Pb}/^{206}\text{Pb}$ value of 0.820 and a mean

square weighted deviation (MSWD) of 1.2 in the Tera-Wasserburg diagram (Fig. 4).

5. Discussion

5.1. Timing of the Sb-Pb mineralization episode

The $^{207}\text{Pb}/^{206}\text{Pb}$ ratios of the S-Sb-Pb minerals range from 0.817 to 0.821, with an average value of 0.820 (our unpublished data). Because the amounts of massed radiogenic Pb are negligible in the Sb-Pb minerals due to their extremely low U contents, the $^{207}\text{Pb}/^{206}\text{Pb}$ compositions of the Sb-Pb minerals are approximately identical to those of the precipitating fluids. The initial $^{207}\text{Pb}/^{206}\text{Pb}$ value of the dated calcite approximates that of the precipitating fluid, which implies that the influence of common Pb and post-ore modification are negligible. In

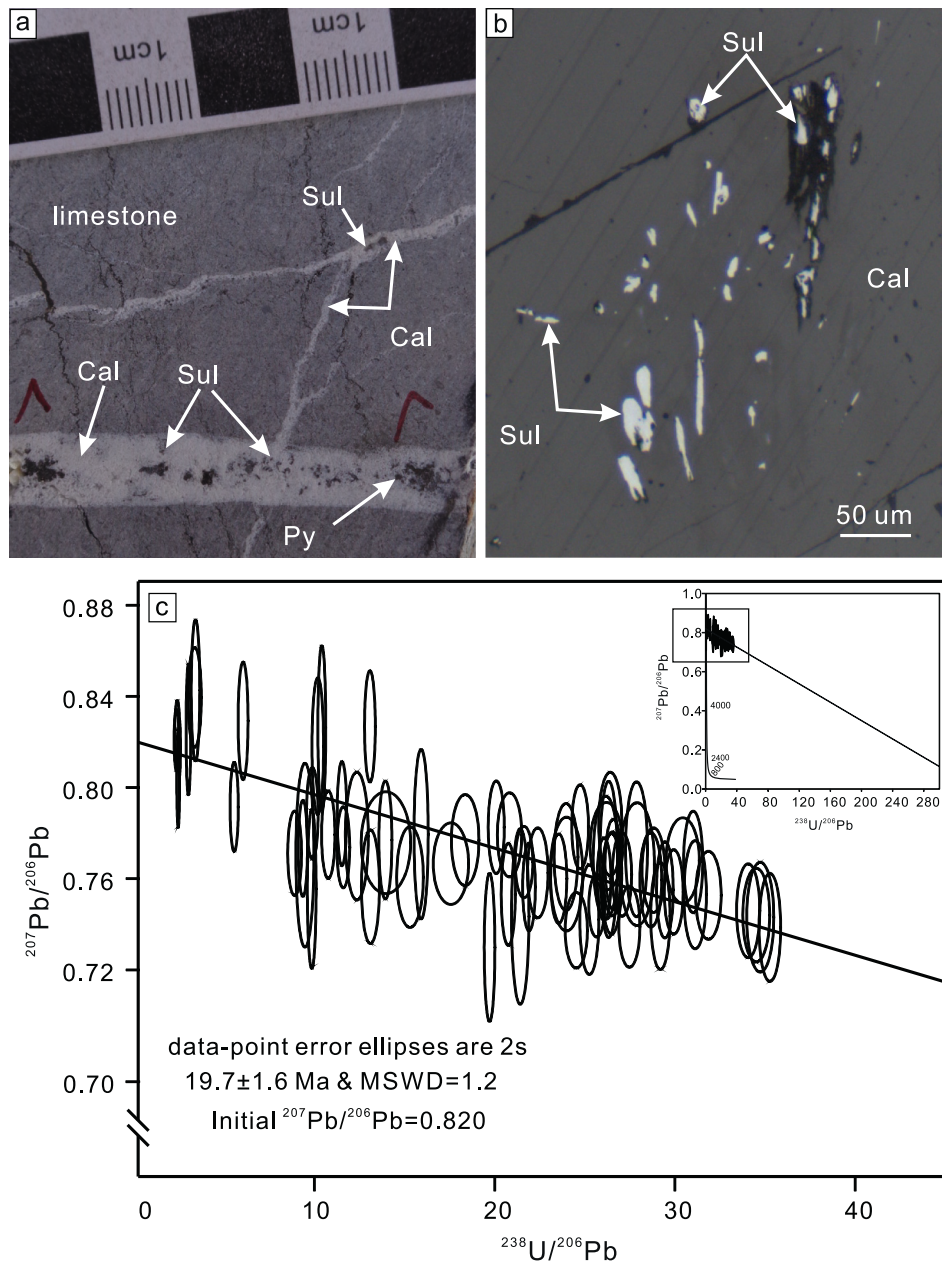


Fig. 4. In situ U-Pb dating of calcite from the Lanuoma deposit. a and b show the dated calcite, intergrown with Sb-Pb minerals; c shows the U-Pb age of calcite in the Tera-Wasserburg diagram. Data are from Suppl. A1. Abbreviations: Cal, calcite; Sul, Sb-Pb minerals; Sp, sphalerite; Py, pyrite.

addition, the Lanuoma deposit is an epithermal deposit, as indicated by the ore textures, which should have formed after the deposition of the host rocks (Bolila Formation >200 Ma) and the ore-controlling thrust faults (initially occurring in the Eocene) (Tang et al., 2006). Moreover, the paragenetic mineral associations suggest that the Lanuoma deposit consists of early Zn-dominant mineralization and late Sb-Pb mineralization (Sheng et al., 2019), and Bi et al. (2019) reported an early Zn mineralization age of 30 Ma. Thus, the calcite age agrees with these geological constraints, and the calcite U-Pb age (19.7 ± 1.6 Ma) likely represents the Sb-Pb mineralization age at Lanuoma.

5.2. Genesis of the Lanuoma deposit and regional significance

Wang et al. (2001) suggested that the interval of 24–17 Ma represents a geodynamic transition from processes controlled mainly by crustal deformation to those dominated by mantle tectonics in the

evolution of the eastern India-Asia collision zone. Temporally, the age of the Lanuoma Sb-Pb mineralization agrees well with the timing of leucogranite intrusions related to ductile left-lateral shearing along three main shear zones (i.e., Chongshan, Gaoligongshan and Ailaoshan-Red River; Figs. 1 and 5) in the southern Sanjiang region (Deng et al., 2014). Studies have shown that regional sinistral strike-slip shearing was almost contemporaneously initiated ca. 32 Ma and terminated ca. 20–17 Ma (Searle et al., 2010; Cao et al., 2011; Liu et al., 2012; Tang et al., 2013); it subsequently transitioned to dextral movements (Zhang et al., 2010; Deng et al., 2014). Such a transition could have resulted in a rapid release of pressure on the planes of faults, upwelling of the underlying fluids along the faults and their discharge upon mixing with local fluids. Xu (2017) suggested that deep fluids with higher $^3\text{He}/^4\text{He}$ ratios (0.09–0.14 Ra for fluid inclusions in S-Sb-Pb minerals) contributed to the late Sb-Pb mineralization episode. Based on the $\delta^{34}\text{S}_{\text{CDT}}$ values (-1.6 to +6.0‰) of S-Sb-Pb minerals (boulangerite), orpiment and

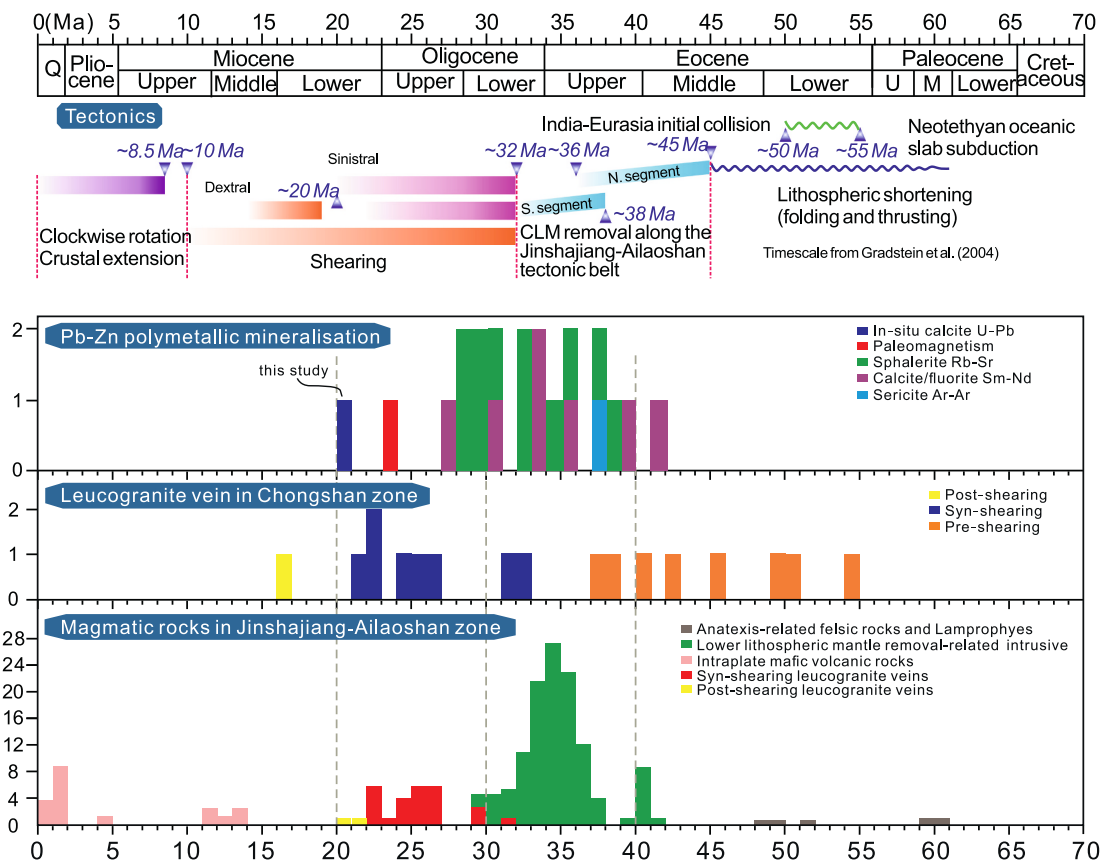


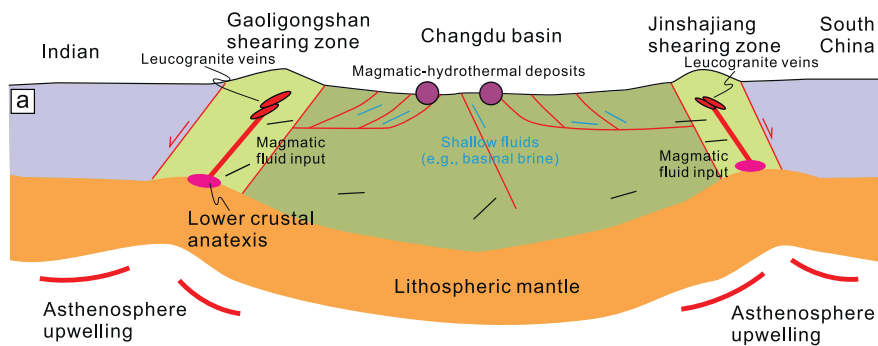
Fig. 5. Tectono-magmatic-metallogenic events in and around the SMB. The data are from Suppl. A2, Deng et al., 2014, and references therein. Abbreviations: CLM-continental lithospheric mantle.

pyrite from the late Sb-Pb mineralization episode, magmatic sulfur and sulfate-reduced sulfur (e.g., sulfur originating from gypsum via thermochemical sulfate reduction) have been suggested to have jointly participated in the Sb-Pb mineralization (Tao et al., 2011; Yang, 2011; Xu, 2017; Sheng et al., 2019). Furthermore, $\delta^{13}\text{C}_{\text{PDB}}$ (-4.7 to +0.4‰; Xu, 2017) and $\delta^{18}\text{O}_{\text{SMOW}}$ (+14.2 to +20.4‰; Xu, 2017) data of calcite from the Sb-Pb mineralization also display composite signals between the deep fluid ($\delta^{13}\text{C}_{\text{PDB}}$ of mantle: -5 to -2‰; $\delta^{13}\text{C}_{\text{PDB}}$ of magma: -9 to -3‰; Taylor, 1986) and sedimentary carbonates ($\delta^{13}\text{C}_{\text{PDB}}$: -2 to +3‰, approximately 0‰; Veizer et al., 1980). As a result, the ore-forming fluid of Sb-Pb mineralization is characterized by moderate-low homogenization temperatures (208–297 °C) and a wide range of salinities (3.2–17.1 wt% NaCl_{eq}) measured from fluid inclusions in orpiment from the late Sb-Pb mineralization (Xu, 2017). Therefore, fluid inclusions and S, C-O and He isotope data suggest the participation of magmatic components in the Sb-Pb mineralization of the Lanuoma deposit (Tao et al., 2011; Xu, 2017; Sheng et al., 2019), which implies that the magmatic components were most likely sourced from contemporaneous magmatism related to strike-slip shearing and resulting from the remelting of crustal materials (Wang et al., 2006; Zhang et al., 2010; Liu et al., 2015). Some analogues may refer to the Zhaxikang, Rujevac and Čumavići Pb-Zn-Sb deposits in the Tethyan domain, which have been shown to share a genetic relationship with coeval felsic magmatism (Janković et al., 1977; Radosavljević et al., 2014, 2016; Xie et al., 2017; Zhou et al., 2018; Wang et al., 2019). It is postulated that the ~20 Ma magmatic fluid induced by coeval shearing strike-slip movements along the margins of the Cenozoic basins was transported along the detachment zone of large-scale thrust systems and then mingled with shallow fluid (e.g., basal fluid); metals subsequently precipitated due to fluid mixing and/or cooling (e.g., at Lanuoma) during the transition from transpression to transtension (Fig. 6a).

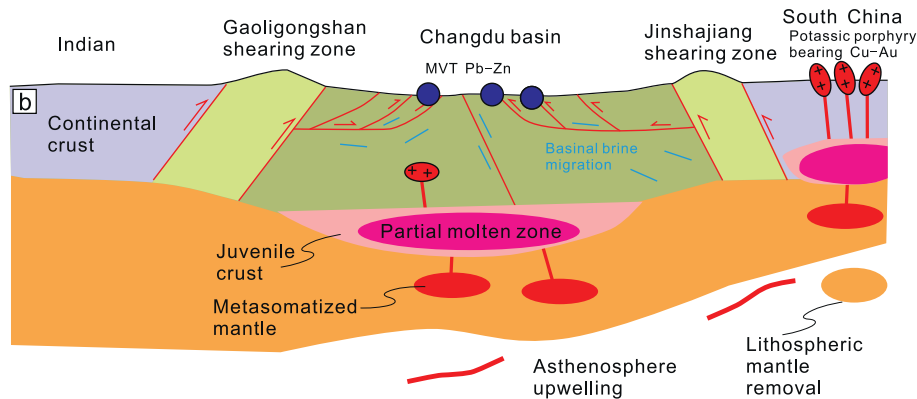
Bi et al. (2019) reported an isochron age of 30 Ma using sphalerite Rb-Sr dating at Lanuoma, consistent with the ages (28–38 Ma) of most Pb-Zn deposits in the SMB (Fig. 5), likely reflecting a regional prolonged hydrothermal Pb-Zn event. Many studies have considered that this episode of Pb-Zn mineralization was closely associated with the circulation of basinal brines (Liu et al., 2017b; Tang et al., 2017; Wang et al., 2017; Bi et al., 2019; Mu et al., 2021). Notably, the 38–28 Ma Pb-Zn deposits display synchronization with the potassic-ultrapotassic igneous rocks (~40–26 Ma; Deng et al., 2014) along the Jinshajiang-Ailaoshan tectonic belt. In particular, the Pb-Zn deposits in the northern segment of the SMB appear slightly older than those in the southern SMB. For example, calcite Sm-Nd dating shows ages of 41–38 Ma for the Jiamoshan, Zhaofayong and Lalongla deposits in the Changdu Basin (Liu et al., 2016), in contrast to the ages of the deposits (approximately 30 Ma or younger) within the Lanping Basin; these differences are likely related to the distinct tectonic locations of these deposits (Deng et al., 2014). This trend is in line with the increasingly younger ages of the potassic-ultrapotassic igneous rocks from north to south (Hou et al., 2003; Lu et al., 2012; Deng et al., 2014), indicating a likely genetic connection between them. Although most of these Pb-Zn deposits show little evidence of direct magmatic inputs, the fluid inclusion homogenization temperatures (~150–250 °C; Tang et al., 2011; Tao et al., 2011; Liu et al., 2016; Xu, 2017), which are higher than values expected for geological reasonable geothermal gradients and estimated stratigraphic burial temperature at the time of mineralization, are most likely explained by anomalous heat sources related to coeval magmatism. The late Eocene to early Oligocene potassic-ultrapotassic igneous activities had clear impacts on the interior of basins, as evidenced by the exposures of the Lianhuashan and Zhuopan intrusions with ages of 35–34 Ma in the Lanping Basin (Liu et al., 2017a; Du et al., 2018).

After the removal of the lithospheric mantle that generated the

Transension along the marginal shearing zones in the end Oligocene to Early Miocene (24–16 Ma)



Transpression along the marginal shearing zones in the Late Eocene to Early Oligocene (41–28 Ma)



potassic-ultrapotassic igneous rocks, the crust of the South China block was split with a partial thrust westward underneath the Changdu - Lanping - Simao blocks, resulting in transpressional faults along the margins of those basins (Fig. 6b). In response to regional compression, voluminous basinal fluid was released and transported along the contemporaneous thrust faults, producing Pb-Zn deposits similar to MVT ore deposits elsewhere. Therefore, the base metal mineralizations of the SMB are concluded to be composed of late Eocene to early Oligocene Pb-Zn ore mineralization (e.g., Baiyangping Pb-Zn deposits) that formed in a transpressional setting and Miocene Sb(-Pb) mineralization that occurred in a transensional setting during the India-Asia continental collision. Cenozoic Sb mineralized deposits are gradually being reported, such as the Bijiashan Sb deposits (Sb: 52664 tons; Tong et al., 2016) and the Shangnuluo and Baiji Pb-Zn-Sb deposits (Liu et al., 2021) in the Lanping Basin. Moreover, the age of Sb(-Pb) mineralization from the Baiji deposit also suggests a transensional setting during the India-Asia continental collision (Liu et al., 2021). The similarity of the tectonic evolution between these basins in the SMB suggests that this region has significant exploration potential for Sb(-Pb) resources.

6. Conclusions

In situ U-Pb dating of calcite indicates that the Sb-Pb mineralization in the Lanuoma deposit occurred at 19.7 ± 1.6 Ma, which was likely related to contemporaneous Miocene tectono-magmatic activity in the Sanjiang Tethys region. The overprinting of the ~20 Ma Sb-Pb mineralization on early Oligocene (30 Ma) Zn mineralization during the India-Asia collision generated the characteristic Zn-Pb-Sb assemblage at

Fig. 6. Tectonic model for the formation of sediment-hosted Pb-Zn polymetallic deposits in the SMB. a During the transensional regime, asthenospheric upwelling beneath the marginal orogens induced anatexis of lower crustal materials, resulting in leucogranite veins that contributed magmatic fluid for base metal mineralization. b During the transpressional regime associated with the removal of the lithospheric mantle under the South China block, the underplating of the asthenosphere in the basin led to partial melting of juvenile crust, which heated circulating basinal brines and drove fluid transport upwards along thrust faults to form the MVT Pb-Zn deposits (modified from Deng et al., 2014).

Lanuoma. The discovery of Miocene Sb-Pb mineralization suggests great exploration potential for Sb(-Pb) resources within the SMB.

Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have influenced the work reported in this submission.

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Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.gexplo.2022.107004>.

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