

Importance of NO₃ radical in particulate nitrate formation in a southeast Chinese urban city: New constraints by δ¹⁵N-δ¹⁸O space of NO₃

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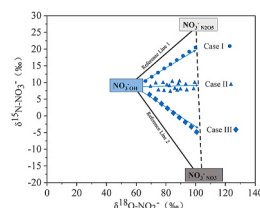
HIGHLIGHTS

The δ¹⁵N-NO₃ and δ¹⁸O-NO₃ ranged from -1.9‰ to +12.1‰ and +69.1‰ to +95.5‰, respectively.

The daily δ¹⁵N-NO₃ was independent with the corresponding δ¹⁸O-NO₃ (R² = 0.06, *p* > 0.05).

Results highlighted the importance of NO₃ radical in NO_x oxidation in polluted urban environment.

GRAPHICAL ABSTRACT



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ABSTRACT

Stable nitrogen and oxygen isotopic signatures of nitrate in atmospheric fine-mode particulate (as δ¹⁵N-NO₃ and δ¹⁸O-NO₃ in PM_{2.5}) was proposed to be useful in distinguishing the sources and oxidation chemistry of NO_x (NO_x = NO + NO₂). In the present study, the chemical oxidation processes of atmospheric NO_x in urban Nanchang, the capital of Jiangxi province, southeast of China were estimated based on the δ¹⁵N-δ¹⁸O space of NO₃. Daily PM_{2.5} samples (n = 91) were collected during wintertime of 2017–2018 (1 November to 31 January), the major water-soluble inorganic ions and the dual isotopic signatures of NO₃ were measured. During the observations, the NO₃ concentrations in PM_{2.5} varied widely from 0.8 μg/m³ to 57.7 μg/m³, on average of 15.5 ± 6.7 μg/m³. The δ¹⁵N-NO₃ and δ¹⁸O-NO₃ also ranged widely, from -1.9‰ to +12.1‰ (+6.5 ± 3.7‰) and +69.1‰ to +95.5‰ (+85.9 ± 17.7‰), respectively. The daily δ¹⁵N-NO₃ was observed to be independent with the corresponding δ¹⁸O-NO₃ (R² = 0.06, *p* > 0.05), which contrasted with many previous reports. By linking the δ¹⁵N-NO₃ to the NO_x oxidation chemistry, we tried to explore the environmental significance of the δ¹⁵N-δ¹⁸O space of NO₃. Our results suggested that the nocturnal pathways (e.g., N₂O₅ hydrolysis and NO₃ radical reacted with hydrocarbons: NO₃+HCs) dominated the chemical conversion of NO_x to NO₃ in the wintertime of Nanchang, with an average fractional contribution of 60%. Interestingly, results also indicated the importance of the NO₃+HCs channel in NO_x oxidation (on average of 33%), which can reach 45% during extreme nitrate aerosol polluted days. Our observations highlighted the importance of NO₃ radical in NO_x oxidation and particulate nitrate formation in polluted urban environment.

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1. Introduction

Anthropogenic emission of nitrogen oxides (NO_x, e.g., NO, NO₂) has significantly altered the natural nitrogen cycle and linked to numerous serious environmental and health problems (An et al., 2019; Fang et al., 2011b; Geng et al., 2019). Global investigations pointed out that fossil-fuel combustion and agricultural activity were responsible for the NO_x emission to the atmosphere (Fang et al., 2015; Geng et al., 2014; Vasilakos et al., 2018; Song et al., 2021). Once emitted to the atmosphere, the majority of NO_x was oxidized into inorganic or organic nitrate, which was the dominant sink for global NO_x (Alexander et al., 2019; Womack et al., 2019). As depicted in Fig. 1, the NO_x converted into inorganic nitrate mainly through three pathways in inland cities, which included: i) NO₂ oxidation by OH radicals in the daytime (termed as NO₂+OH); ii) hydrolysis of dinitrogen pentoxide (N₂O₅) on the surface of wet aerosols; iii) nitrate radicals (NO₃) reacting with saturated organic hydrocarbon (HC) compounds (as NO₃+HC; Seinfeld and Pandis, 2016). Generally, the channel of NO₂+OH dominated during the daytime and summer season, while the N₂O₅ hydrolysis and NO₃+HC channels were typically prevalent during winter and at night since the N₂O₅ and NO₃ radical were thermally unstable and easily photolyzed (Chang et al., 2019a; He et al., 2018).

The oxygen isotopic composition of nitrate in fine-mode particulate (PM_{2.5}), e.g., δ¹⁸O–NO₃⁻ and Δ¹⁷O–NO₃⁻ were suggested useful in inferring the relative importance of different channels of NO₃ oxidation pathways (Alexander et al., 2009; Hastings, 2004; Hastings et al., 2003; He et al., 2020; Michalski, 2005; Michalski et al., 2003; Savarino et al., 2016; Walters and Michalski, 2016; Wang et al., 2019; Fan et al., 2020). Usually, NO₃ produced via nocturnal channels of N₂O₅ hydrolysis and NO₃+HC were characterized by ¹⁸O-enriched signatures (e.g., 97% ~109%), as the N₂O₅ and NO₃ radical can reach equilibrium with the oxidant of O₃ (δ¹⁸O–O₃ was predicted to be +117 ± 5‰ relative to international standard) during the chemical oxidation reactions (Fig. 2; Chang et al., 2019a; Fang et al., 2011a; He et al., 2020; Zong et al., 2020; Fan et al., 2020). In contrast, NO₃ generated through the NO₂+OH channel was registered as relatively lower δ¹⁸O values. For example,

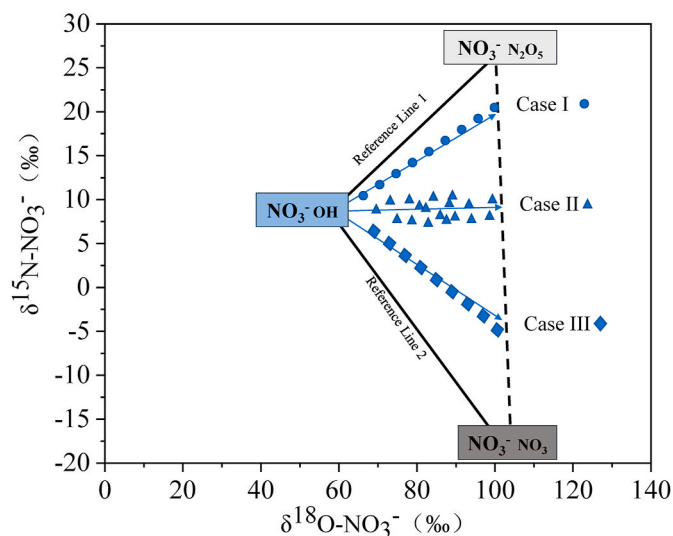


Fig. 2. Theoretical δ¹⁸O–δ¹⁵N space of particulate nitrate generate through daytime channel (assuming $f_{\text{NO}_2} = 0.8$) and nighttime channels under the typical wintertime conditions of urban Nanchang. δ¹⁵N–NO₂ was assumed to be 0. The reference line 1 and 2 represented that nocturnal NO₃ formed completely through N₂O₅ and NO₃+HC channels, respectively. In the real environment, δ¹⁸O of nitrate may correlated with the corresponding δ¹⁵N values positively (case I) or negatively (case III) or not (case II).

Fang et al. (2011a) suggested the endmember value of δ¹⁸O–NO₃⁻ via NO₂+OH channel was approximately +55‰, which was mainly attributed to the involvement of significantly ¹⁸O-depleted OH during NO₂ transformation (Fang et al., 2011a). Therefore, the δ¹⁸O–NO₃⁻ was extensively used to distinguish the relative contribution of daytime and nocturnal channels; however, it is difficult to differentiate the relative importance between N₂O₅ hydrolysis and NO₃+HC channels based only on the δ¹⁸O–NO₃⁻ approach.

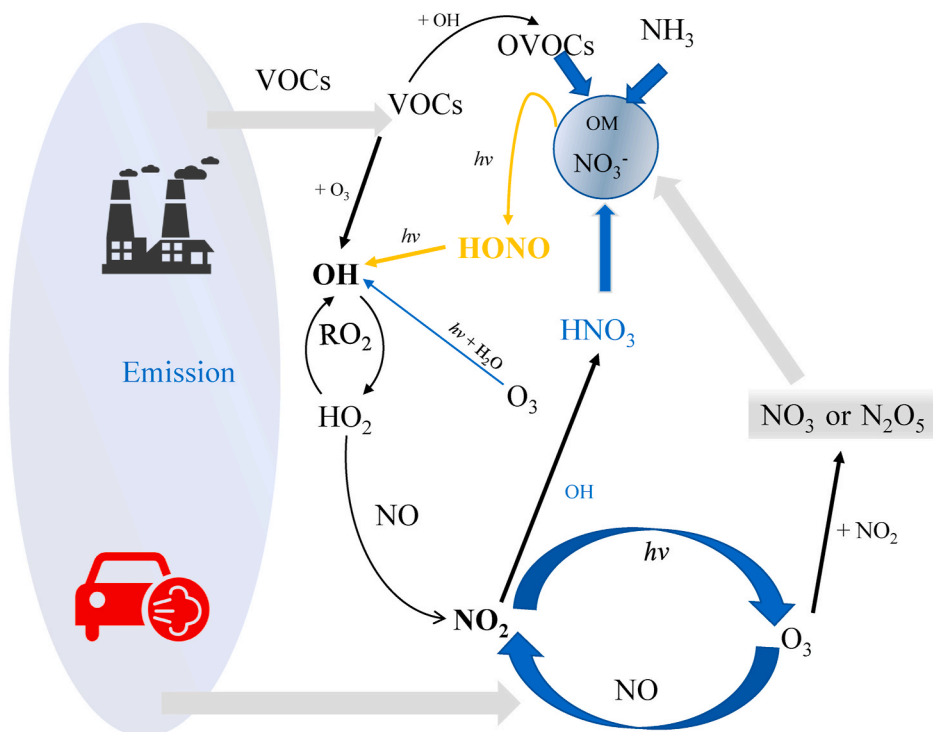


Fig. 1. Simplified schematic of the relevant particulate nitrate formation channels in polluted urban atmosphere (Lu et al., 2019). Blue marked chemical species dominant in daytime reaction, grey in nocturnal chemistry and the HONO triggered reactions are in debates.

While $\delta^{15}\text{N}\text{-NO}_3$ has been widely used to fingerprint the NO_x sources (Hastings, 2004; Morin et al., 2008; Song et al., 2019; Walters et al., 2015c), previous theoretical researches suggested it may also provide valuable insights into the NO_x oxidation chemistry at a process level (Walters and Michalski, 2015b, 2016). This was because the different oxidation mechanisms responsible for the formation of aerosol NO₃ may lead to distinctive nitrogen isotopic signatures that can be recorded in the environment media (e.g., PM_{2.5}, dry/wet deposition). As shown in Fig. 2, NO₃ generated through NO₂+OH pathway was registered as middle-ranking $\delta^{15}\text{N}$ values but most ^{18}O -depleted signatures, while that through the NO₃+HC or N₂O₅ hydrolysis pathways was characterized by similar ^{18}O -enriched signatures; however, the NO₃+HC channel was associated with ^{15}N -depleted signatures, as contrasted to the N₂O₅ hydrolysis pathway which was significantly ^{15}N -enriched (Walters and Michalski, 2016; Song et al., 2020; Liu et al., 2020; Zong et al., 2017). Therefore, the $\delta^{15}\text{N}\text{-}\delta^{18}\text{O}$ space of NO₃ may help to elucidate the fractional contribution of the three pathways in NO_x oxidation processes. However, the imprinted signals of $\delta^{15}\text{N}\text{-NO}_3$ during the oxidation pathways can be impacted substantially by the changes in the emission sources of NO_x and the underlying N fractionation effects. Taking the urban environment as an example, the dominant anthropogenic sources of NO_x from fossil fuel combustion were characterized by relatively ^{15}N -enriched values, while NO_x emitted through nitrification and denitrification processes was usually ^{15}N -depleted (Felix and Elliott, 2014; Felix et al., 2012; Walters et al., 2015a; Walters and Michalski, 2015b; Yu and Elliott, 2017). In brief, the environmental $\delta^{15}\text{N}\text{-NO}_3$ signatures were a hybrid of NO_x source information and oxidation processes-mediated fractionation effect (Zong et al., 2017; Fan et al., 2020; Liu et al., 2020). This may be the reason that hindered the deep understanding of the environmental significance of $\delta^{15}\text{N}\text{-}\delta^{18}\text{O}$ space of NO₃ in NO_x oxidation chemistry.

In the present study, the daily PM_{2.5} samples (23.5h) were collected from 1 November 2017 to 31 January 2018 (n = 91) at the urban city of Nanchang, the capital of Jiangxi, located in the southeast of China (Fig. S1). The urban city of Nanchang still experienced severe haze pollution in recent years, in which the particulate nitrate was suggested as the predominant inorganic component in PM_{2.5} (Xiao et al., 2020). The aerosol inorganic chemical species and properties (e.g., aerosol acidity), as well as the dual isotopic composition of nitrate were quantified, and those data were used to investigate the oxidation processes and factors that controlled the wintertime particulate nitrate formation in urban Nanchang. Our results highlighted the relative importance of nocturnal NO₃ radical chemistry, particularly the reaction of NO₃ radical with organic compounds in wintertime nitrate pollution in the urban environment.

2. Material and method

2.1. Site description and sampling

The sampling campaign was conducted on the Qingshanhu campus of Nanchang University (28°41'N, 115°56'E, Fig. S1) from 1 November 2017 to 31 January 2018. The campus is located in the old urban area of Nanchang city, about 3 km away from the Nanchang railway station and near the first-ring viaduct and Metro line No.1. The site is registered as a typical urban area for business and residents, with no industry and other point sources (Xiao et al., 2020). A total of 91 samples (89 p.m._{2.5} samples and 2 blank samples) were collected in the sampling campaign. The instruments were set on the roof of a six-story teaching building, approximately 25m above the ground. During the sampling campaign, the prevailing wind direction was northeast, with an average speed of 1.7 ± 0.7 m/s.

The PM_{2.5} sampling procedure has been well-described elsewhere (Xiao et al., 2020; Zhang et al., 2020c). In brief, quartz fiber filters were pre-combusted at 450 °C for 3h to remove impurity then used for PM_{2.5} collection with a high-volume air sampler (KC-1000, equipped with a

PM_{2.5} impactor). The duration of the daily sample was 23.5h at a flow rate of 1.05 m³/h. PM_{2.5} samples were immediately delivered to the laboratory and stored in the refrigerator once collected. Gaseous HNO₃ and particulate nitrate may be sampled simultaneously during the campaign, and therefore described as atmospheric inorganic nitrate in this study. The meteorological and air quality parameters (e.g., temperature, wind speed, and direction, relative humidity, hourly-resolution of NO, NO₂, SO₂, O₃) were acquired from the nearby station of Jiangxi Provincial Center for Environment Monitoring (approximately 500m).

2.2. Isotopic and chemical analysis

The laboratory analysis of water-soluble inorganic ions (including Cl⁻, NO₃⁻, SO₄²⁻, Na⁺, NH₄⁺, K⁺, Mg²⁺, etc.) and the dual isotopic composition of NO₃ ($\delta^{18}\text{O}$ and $\delta^{15}\text{N}$) was conducted in Jiangxi Province Key Laboratory of the Causes and Control of Atmospheric Pollution, East China University of Technology and the analytical procedure can be found in our previous work (Guo et al., 2020; Zhang et al., 2020a). Briefly, the major inorganic chemical species in the PM_{2.5} samples were extracted using ultrapure water (Millipore, 18.2MΩ) and the extract solutions were filtered. Inorganic ions in the extract solutions were determined using routine chemical methods (Guo et al., 2020). The bacterial approach was used to quantify the $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of NO₃ (Casciotti et al., 2002; Fang et al., 2011a; Sigman et al., 2001; Liu et al., 2018). The bacteria *Pseudomonas aureofaciens* (ATCC 13985#) without N₂O reductase was able to turn the NO₃ into gaseous N₂O quantitatively with very little impact on the $\delta^{18}\text{O}$ values of substrate NO₃ (Hastings, 2004; Hastings et al., 2003). The gaseous N₂O was then introduced to a continuous-flow isotope ratio mass spectrometer coupled with Gasbench-II for $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ determination. The reported data was first calibrated using international recognized nitrate standards (USGS32, USGS34, USGS35, IAEA-NO3) with certified $\delta^{18}\text{O}$ and $\delta^{15}\text{N}$ values and presented using the δ notation (in units of per mil, ‰) as follows in this study:

$$\delta^{18}\text{O}\text{-NO}_3 = [(^{18}\text{O}/^{16}\text{O})_{\text{sample}} / (^{18}\text{O}/^{16}\text{O})_{\text{VSMOW}} - 1] \times 1000$$

$$\delta^{15}\text{N}\text{-NO}_3 = [(^{15}\text{N}/^{14}\text{N})_{\text{sample}} / (^{15}\text{N}/^{14}\text{N})_{\text{N}_2 \text{ in air}} - 1] \times 1000$$

The presented $\delta^{15}\text{N}\text{-NO}_3$ values were also corrected for the contribution of mass-independent $^{14}\text{N}\text{-}^{14}\text{N}\text{-}^{17}\text{O}$ to the analyst N₂O, because the aerosol NO₃ has substantial mass-independent ^{17}O anomaly (Hastings, 2004; Hastings et al., 2003). Due to the unavailability of $\delta^{17}\text{O}$ determinations in our laboratory at present, published relationships between $\delta^{18}\text{O}$ and $\Delta^{17}\text{O}$ were used in this study (Hastings, 2004; Hastings et al., 2003; Michalski et al., 2003; Wang et al., 2019). The standard deviations of 20 replicate injections of international standards were better than $\pm 0.2\text{‰}$ for $\delta^{15}\text{N}$ and $\pm 0.5\text{‰}$ for $\delta^{18}\text{O}$. Concentrations of NO₂ were usually smaller than the NO₃ by two orders of magnitude and therefore did not impact the isotopic composition of NO₃. More details were included in the Supporting information.

2.3. Model simulation of NO₃ radical, N₂O₅, and aerosol properties

The aerosol acidity and liquid water content (ALWC) were predicted using the thermodynamic model of ISORROPIA-II, which was widely used to describe the aerosol thermodynamic equilibrium, e.g., the gas-to-particle partition of atmospheric nitrate (nitrate partition ratio: $\epsilon_{\text{NO}_3} = \text{NO}_3 / (\text{NO}_3 + \text{HNO}_3)$) (Fountoukis and Nenes, 2007; Guo et al., 2015, 2017; Weber et al., 2016). To improve the simulation of aerosol acidity and ALWC, the ISORROPIA-II model was run iteratively until the output particle NH₄⁺ concentration was very close to the measured one (<1% in this study, Zhang et al., 2020a). Usually, the “S” curve of ϵ_{NO_3} against the predicted pH was used to assess the degree of the gas-to-particle partition of atmospheric nitrate. The performance of ISORROPIA-II has been evaluated elsewhere (Guo et al., 2015, 2017;

Weber et al., 2016). The mixing ratios of NO_3 radical and N_2O_5 near the sampling sites were simulated using the standard Master Chemical Mechanism (MCM in website <http://mcm.leeds.ac.uk/>, last access: 3 June 2020), which have been well-presented by Xiao et al. (2020) (e.g., model performance and results). Detailed information on the two models was supported in Text S1.

2.4. The theoretical basis of $\delta^{15}\text{N}$ - $\delta^{18}\text{O}$ space of NO_3^- in revealing the different oxidation pathways

As abovementioned (and in Fig. 2), the three nitrate formation pathways were characterized by relatively distinct $\delta^{15}\text{N}$ - $\delta^{18}\text{O}$ signatures (Walters and Michalski, 2016; Liu et al., 2020). The isotopic fractionation associated with the three pathways has been theoretically determined and supported in Text S2 in detail. Assuming $\delta^{15}\text{N}\text{-NO}_x = 0$ and $T = 273.2\text{K}$, the endmember value of $\delta^{15}\text{N}\text{-NO}_3^-$ through OH channel was approximately $7.9 \pm 3.1\text{‰}$ (e.g., kinetic fractionation factor of OH oxidation was ignored, isotope equilibrium fractionation factor between NO and NO_2 ranged from 42‰ to 45‰, isotope equilibrium fractionation factor from gaseous HNO_3 to particle NO_3^- was minor since the ϵ_{NO_3} approximately 100‰ for the most of time, the fraction of NO_2 to NO_x (f_{NO_2} , Fig. S3) was from 0.70 to 0.95, see Text S2 for detailed explanations), the endmember value of $\delta^{15}\text{N}\text{-NO}_3^-$ through N_2O_5 hydrolysis channel was approximately $+29.4 \pm 1.1\text{‰}$ and that through NO_3+HC pathway was roughly $-18.7 \pm 1.1\text{‰}$; in comparison, the endmember value of $\delta^{18}\text{O}\text{-NO}_3^-$ generated through NO_3+HC and N_2O_5 pathway compared well, on average of $+92 \pm 9.5\text{‰}$, while that through NO_2+OH channel approximated $+55\text{‰}$ (Fang et al., 2011a). Therefore, the $\delta^{18}\text{O}\text{-NO}_3^-$ signatures can be used to distinguish the contribution of daytime and nocturnal channels (NO_3+HC and N_2O_5 hydrolysis), despite with nonnegligible uncertainty. Furthermore, the daily $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of NO_3^- should be within the triangle region restricted by the theoretical endmember values of $\delta^{15}\text{N}\text{-}\delta^{18}\text{O}$ of NO_3^- (Fig. 2). To assess the relative contribution of the three different pathways, we further assumed the complete mixing of NO_3^- formed through daytime and nocturnal pathways. Consequently, a linear regression was anticipated between the $\delta^{15}\text{N}\text{-NO}_3^-$ against the corresponding $\delta^{18}\text{O}\text{-NO}_3^-$ with the slope determined by the relative importance of NO_3+HC channel in the

overall nocturnal pathways (Fig. 2). Simply, the slope of $\delta^{15}\text{N}\text{-NO}_3^-$ against the $\delta^{18}\text{O}\text{-NO}_3^-$ was approximately 0.54 with the absence of NO_3+HC channel, while the slope approximated -0.65 with the absence of N_2O_5 hydrolysis pathway (Fig. 2). The slope was, therefore, regarded as the additional strong constraint on apportioning the relative contribution of the two nocturnal pathways.

3. Results and discussion

3.1. Temporal evolution of $\text{PM}_{2.5}$ NO_3^- and relevant parameters

Fig. 3 depicted the time series of NO_3^- in $\text{PM}_{2.5}$ and the relevant parameters, especially those impacted the accumulation of NO_3^- (e.g., T, RH, and particle pH). The temperature was gradually decreased with time, while the RH varied widely, from 30% to 95%. Consistent with previous reports, the concentrations of NO_3^- exhibited large variability during the campaign, ranged from $2.8 \mu\text{g}/\text{m}^3$ to $57.7 \mu\text{g}/\text{m}^3$ (on average of $20.5 \pm 10.4 \mu\text{g}/\text{m}^3$). The mean concentration of NO_3^- in the wintertime of Nanchang city was observed to be in line with that in nearby urban cities, e.g., Changsha ($19.4 \mu\text{g}/\text{m}^3$ from September 2013 to August 2014, Zhang et al., 2020), Wuhan ($23.9 \mu\text{g}/\text{m}^3$ in winter 2019, Zheng et al., 2020) and Nanjing ($16.7 \mu\text{g}/\text{m}^3$ from March 2016 to August 2017, Lin et al., 2020). Furthermore, the loadings of particulate NO_3^- were higher than that of SO_4^{2-} (range: $2.7\text{--}30.2 \mu\text{g}/\text{m}^3$; mean: $11.9 \pm 5.2 \mu\text{g}/\text{m}^3$), especially during haze episodes (Fig. 3). In fact, the ratio of NO_3^- to SO_4^{2-} ($\text{NO}_3^-/\text{SO}_4^{2-}$) increased from clean periods (<1.0) to polluted days (>2.0 , Fig. S2), on average of 1.8 during the whole observations, which was even higher than that in polluted northern city environments (e.g., 1.4 in Beijing of winter 2016 (Xu et al., 2019), 1.6 in several cities surrounded North China Plain in winter 2017 (Fu et al., 2020)). These results may suggest the increasing importance of NO_3^- in the haze development in southern urban cities of China, which has been recorded as a distinct feature over China. Other major ions, such as NH_4^+ presented parallel trends to the evolution of NO_3^- (Fig. 3).

Theoretically, the buildup of particulate NO_3^- highly depends on two subprocesses: the oxidation of NO_x to HNO_3 and the thermodynamic equilibrium partition between HNO_3 and NO_3^- (Seinfeld and Pandis, 2016; Zhang et al., 2020c). Usually, the gas-to-particle partition of

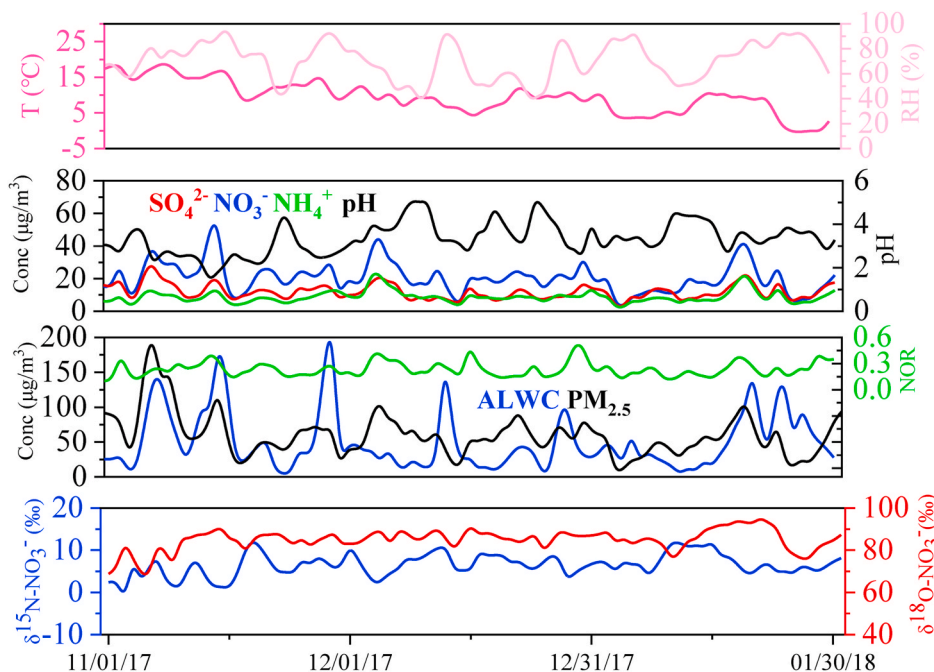


Fig. 3. Temporal evolution of meteorological variables including RH and T, mass concentrations of major inorganic chemical species in $\text{PM}_{2.5}$, NOR (NO_x oxidation ratio), the aerosol properties (acidity and aerosol water content) and the dual isotopic compositions of nitrate in $\text{PM}_{2.5}$ in winters of 2017–2018 in urban Nanchang.

atmospheric nitrate is closely related to the availability of gaseous NH_3 , the variations in RH and temperature (Chen et al., 2020; Guo et al., 2015; Li et al., 2018; Sun et al., 2018). In the present study, the degree of nitrate partition ratio ($\epsilon_{\text{NO}_3^-}$) was estimated using the aerosol pH, which was predicted by the thermodynamic model of ISORROPIA-II, since the ambient measurements of gaseous NH_3 and HNO_3 were unavailable during the campaign. As indicated by the S curve of $\epsilon_{\text{NO}_3^-}$ (Fig. 4), the atmospheric nitrate overwhelming existed in the particle phase, reflected by the relatively moderate pH of $\text{PM}_{2.5}$ (mean of 3.3, mode of 2.5). This was consistent with the observations that abundant NH_3 in the urban environment facilitated the partition of atmospheric nitrate into the particle phase (Chang et al., 2019b; Guo et al., 2018). The aerosol pH generally exhibited declining trends with the accumulation of NO_3^- , with a pH of approximately 1.5 during extreme nitrate polluted period ($\epsilon_{\text{NO}_3^-}$ corresponded to be 0.70, Figs. 3 and 4), suggesting a high ambient gaseous HNO_3 concentration when haze occurred. Although, these results indicated that the HNO_3 formation was the key process that determined the accumulation of particle NO_3^- in the wintertime of Nanchang city, which resembled that in northern urban environments of China (Fu et al., 2020; Wen et al., 2018; Yan et al., 2019). For instance, Fu et al. (2020) confirmed that the increased conversion ratio of NO_x to HNO_3 was the cause of the subdued response of NO_3^- to NO_x emission reduction in urban environments around the North China Plain (Fu et al., 2020). During the observations, the NOR (nitrogen oxidation ratio) varied significantly (from 0.10 to 0.52, mean of 0.24 ± 0.10), and generally increased with the haze development (Fig. 3). In contrast, the gaseous precursors of NO_2 and O_3 generally declined when haze pollution occurred, while the simulated concentrations of NO_3^- radical and N_2O_5 elevated with the haze development (Fig. S4). Together, results identified that the efficient secondary transformation was the driving factor that regulating nitrate aerosol pollution in the winter of urban Nanchang.

3.2. Signatures of $\delta^{18}\text{O}\text{-NO}_3^-$ and $\delta^{15}\text{N}\text{-NO}_3^-$ under different nitrate regimes

During the observations, the daily $\delta^{15}\text{N}\text{-NO}_3^-$ values varied significantly, ranged from -1.9‰ to $+12.5\text{‰}$ with a mean value of $+6.5 \pm 3.7\text{‰}$ (Fig. 3). The $\delta^{15}\text{N}\text{-NO}_3^-$ values of $\text{PM}_{2.5}$ in winter Nanchang were in the range of that reported in winter Shanghai (approximately $+7.5\text{‰}$ in 2013, $+8.9\text{‰}$ in 2016) and Guangzhou (approximately $+7.5\text{‰}$ in

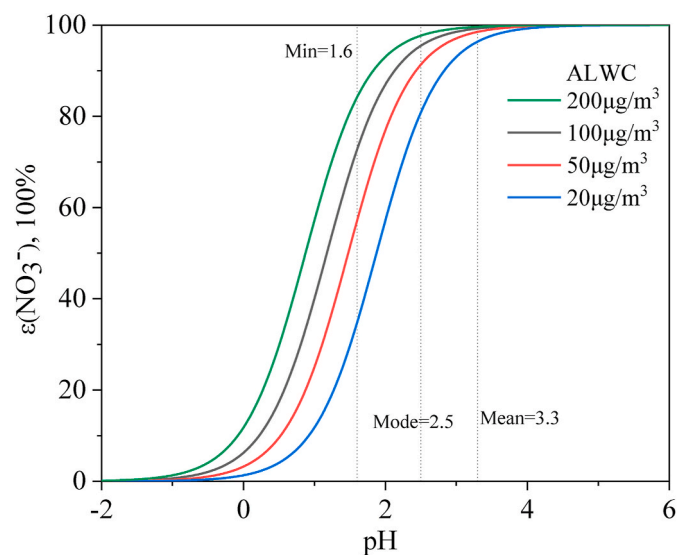


Fig. 4. S curves of $\epsilon(\text{NO}_3^-)$ plotted against ISORROPIA-predicted pH. The aerosol water content (ALWC) was set as 20, 50, 100 and $200 \mu\text{g}/\text{m}^3$, respectively.

2013; Zong et al., 2020), while lower than that in a nearby urban city of Wuhan (approximately $+13\text{‰}$ in 2013; Zong et al., 2020), the northern megacity of Beijing ($+11.9\text{‰}$ in 2014; Song et al., 2019) and Tianjin ($+14.1\text{‰}$ in 201; Feng et al., 2020). Meanwhile, the values of $\delta^{18}\text{O}\text{-NO}_3^-$ ($+85.9 \pm 17.7\text{‰}$) in $\text{PM}_{2.5}$ during our campaign ranged widely from $+69.1\text{‰}$ to $+95.5\text{‰}$, which were also well within the broad ranges reported in previous studies conducted in the northern hemisphere (Feng et al., 2020; He et al., 2020; Michalski, 2005; Michalski et al., 2012; Savarino et al., 2016; Wankel et al., 2010; Xiao et al., 2020; Yang et al., 2014; Zong et al., 2020). However, the mean value of $\delta^{18}\text{O}\text{-NO}_3^-$ in winter Nanchang was generally smaller than that in Shanghai and Wuhan (both approximately $+100\text{‰}$ in winter 2013; Zong et al., 2020), but compared well with that in Beijing ($+84.8\text{‰}$ in winter 2013; Wang et al., 2019) and Tianjin ($+85.1\text{‰}$ and $+83.4\text{‰}$ in January 2017 and 2018, respectively; Feng et al., 2020). Monthly, the $\delta^{15}\text{N}\text{-NO}_3^-$ values gradually increased from November to January, with values of $+5.2 \pm 3.3\text{‰}$, $+6.9 \pm 2.5\text{‰}$ and $+7.3 \pm 2.5\text{‰}$, respectively, which was consistent with previous reports that NO_3^- in $\text{PM}_{2.5}$ was usually ^{15}N -enriched in cold seasons than in warm periods (Gobel et al., 2013; Zong et al., 2017). In comparison, no obvious temporal trend was observed for $\delta^{18}\text{O}\text{-NO}_3^-$.

Simple comparison of $\delta^{15}\text{N}\text{-NO}_3^-$ in $\text{PM}_{2.5}$ among different seasons or urban environment traits are maybe of finite in extending our understandings of the diversity of atmospheric NO_3^- nitrogen signatures, since the $\delta^{15}\text{N}\text{-NO}_3^-$ in $\text{PM}_{2.5}$ was suggested as a comprehensive result of numerous drivers, e.g., source characteristics of NO_x , the isotopic exchange between NO and NO_2 , kinetic isotopic fractionation during NO_x conversion to HNO_3 , Rayleigh-controlled isotopic fractionation with NO_3^- accumulation and the potential equilibrium exchange between HNO_3 and NO_3^- , to the best of our knowledge. Interestingly, the $\delta^{15}\text{N}\text{-NO}_3^-$ presented a negative correlation with the corresponding NO_3^- concentrations, though not significant (Fig. S5), which implies that particulate NO_3^- may originate from different sources or generate from different pathways during the severe nitrate aerosol polluted episode (Fan et al., 2020). Meanwhile, the $\delta^{18}\text{O}\text{-NO}_3^-$ increased significantly with the increasing NO_3^- concentrations (Fig. S5). This indicated both the $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ were related to the nitrate accumulation, to a large extent, and can be used to infer the relative importance of different channels. For clarity, we compared the $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ under different nitrate regimes (e.g., $<5 \mu\text{g}/\text{m}^3$, $5\text{--}10 \mu\text{g}/\text{m}^3$, $10\text{--}15 \mu\text{g}/\text{m}^3$, etc., as presented in Fig. 5) to assess the ability of the dual-isotope signatures in revealing the nitrate formation pathways. As stated, the $\delta^{18}\text{O}\text{-NO}_3^-$ increased with the buildup of NO_3^- (except for the nitrate regime of $25\text{--}30 \mu\text{g}/\text{m}^3$), while the $\delta^{15}\text{N}\text{-NO}_3^-$ increased until NO_3^-

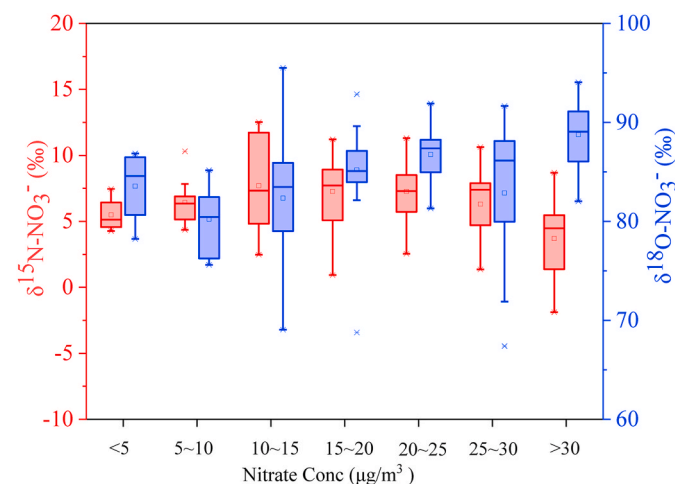


Fig. 5. The dual isotopic signatures of nitrate under different nitrate regimes. Apparently, the $\delta^{15}\text{N}\text{-NO}_3^-$ values declined significantly when nitrate concentrations higher than $20 \mu\text{g}/\text{m}^3$.

reached up to $25 \mu\text{g}/\text{m}^3$ ($+5.5\%$ to $+7.3\%$) and then decreased considerably ($+7.3\%$ to $+4.0\%$) as NO_3^- continued to ascend ($25\text{--}57.6 \mu\text{g}/\text{m}^3$). The parallel increments of $\delta^{18}\text{O}\text{-NO}_3^-$ and $\delta^{15}\text{N}\text{-NO}_3^-$ before NO_3^- accumulated to $25 \mu\text{g}/\text{m}^3$ may indicate the dominance of N_2O_5 hydrolysis in NO_3^- generation, since this channel was associated with ^{15}N and ^{18}O enriched signatures (Walters and Michalski, 2016). The importance of N_2O_5 heterogeneous uptake on the aerosol surface in the fast development of nitrate aerosol across China has been well confirmed in previous researches (Chen et al., 2018; Mitroo et al., 2019; Wang et al., 2017; Yan et al., 2019). However, the opposite variation tendency between $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ under the extreme nitrate aerosol pollution regimes (e.g., $25\text{--}57.6 \mu\text{g}/\text{m}^3$) may suggest the potential contribution of $\text{NO}_3\text{+HC}$ channel in the nitrate formation. In addition, the $\delta^{15}\text{N}\text{-NO}_3^-$ values exhibited negative correlations with the ratio of $\text{NO}_3/\text{N}_2\text{O}_5$ (or NO_3/NO_2) during the observations, further implying the contribution of $\text{NO}_3\text{+HC}$ pathways in nitrate production, to some extent (Fig. S6). Overall, these results preliminarily suggested the potential capacity of $\delta^{15}\text{N}\text{-NO}_3^-$ in elucidating the NO_x oxidation chemistry.

3.3. Relative contribution of specific-pathway in nitrate production of wintertime Nanchang

Throughout the winter campaign, no obvious correlation was observed between $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ in urban Nanchang (Fig. 6 A). A similar phenomenon between $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ or $\Delta^{17}\text{O}\text{-NO}_3^-$ has been reported in previous studies in the coastal megacity of Shanghai and Guangzhou (He et al., 2018; Su et al., 2020), an island in Bohai (Zong et al., 2017) and marine boundary layer (Morin et al., 2009; Savarino et al., 2013), while other reports suggested aforementioned correlations may be positive (for $\text{PM}_{2.5}$ in Beijing, Song et al., 2019; dry deposition in several CASTNET sites of USA, Elliott et al., 2009) or negative (e.g., polar sites, Morin et al., 2008; Savarino et al., 2016; marine island, Altieri et al., 2013; Yang et al., 2014). It appeared that the $\delta^{15}\text{N}\text{-NO}_3^-$ increased with the increasing of $\delta^{18}\text{O}\text{-NO}_3^-$ or $\Delta^{17}\text{O}\text{-NO}_3^-$ in inland environments, while $\delta^{15}\text{N}\text{-NO}_3^-$ changed little or even decreased with the increment of $\delta^{18}\text{O}\text{-NO}_3^-$ or $\Delta^{17}\text{O}\text{-NO}_3^-$ in marine or coastal environments, despite with a few exceptions. We inferred that the N_2O_5 hydrolysis channel was responsible for the positive correlations between the N and O isotopic composition of NO_3^- in inland environments, while the participant of $\text{NO}_3\text{+HC}$ or $\text{NO}_3\text{+DMS}$ channels, of which the endmember was registered as ^{15}N -depleted but ^{18}O -enriched values resulted in relatively negative $\delta^{15}\text{N}\text{-NO}_3^-$ values when $\delta^{18}\text{O}\text{-NO}_3^-$ values (or NO_3^- concentrations) increased. Therefore, results that the distribution tendency corresponds to case II in Fig. 2 may indicate that the $\text{NO}_3\text{+HC}$ channel contributed to the NO_3^- accumulation in the inland urban city of Nanchang, to a large extent.

However, we also realized that the Rayleigh-controlled fractionation effect was also responsible for the ^{15}N -depleted NO_3^- in the regime of

NO_3^- concentrations higher than $30 \mu\text{g}/\text{m}^3$ (corresponding to rapid nitrate accumulation stage). In the rapid-growth stage of particulate NO_3^- , the $\delta^{15}\text{N}\text{-NO}_3^-$ should be gradually decreased with time (Fig. S7). Assuming the overall nitrogen isotopic fractionation effect (ϵ) approximated 10.0% and initial $\delta^{15}\text{N}\text{-NO}_x = 0\%$, values of $\delta^{15}\text{N}\text{-NO}_3^-$ should become less positive from $+10\%$ initially to $+7.7\%$ when NOR reached up to 0.4 (the remaining fraction of NO_x was 0.6, corresponding to NO_3^- concentration of $30 \mu\text{g}/\text{m}^3$, Fig. 3). To alleviate the Rayleigh-driven ^{15}N -depleted effect on $\delta^{15}\text{N}\text{-NO}_3^-$, strategies of either data with NO_3^- concentration higher than $30 \mu\text{g}/\text{m}^3$ ($n = 15$) were removed or all data was rectified for Rayleigh fractionation effect were applied in the following discussion. As presented in Fig. 6 (A, B, and C), no obvious correlations and resemble distribution in biplot of dual-isotopic mixing space were observed under each scenario, which further indicated the importance of the $\text{NO}_3\text{+HC}$ channel in nitrate accumulation in urban Nanchang.

To quantify the relative contribution of individual oxidation pathways, an alternative approach based on the $\delta^{15}\text{N}\text{-}\delta^{18}\text{O}$ space of NO_3^- was presented in this study (see Material and Method). The fractional contribution of OH oxidation channel to NO_3^- production based on the $\delta^{18}\text{O}$ approach fluctuated widely during the campaign, from less than 10% in heavily nitrate polluted days to approximately 70% in clean periods, with a mean of $30.4 \pm 14.3\%$ (Fig. 7). The broad range can be attributed to the distinct daily atmospheric conditions during the observation period. The possible fractional contribution of nocturnal channels correspondingly accounted for $69.6 \pm 14.3\%$. Results directly suggested the dominance of nocturnal chemistry in nitrate production in winter seasons of urban environments, consistent with previous reports

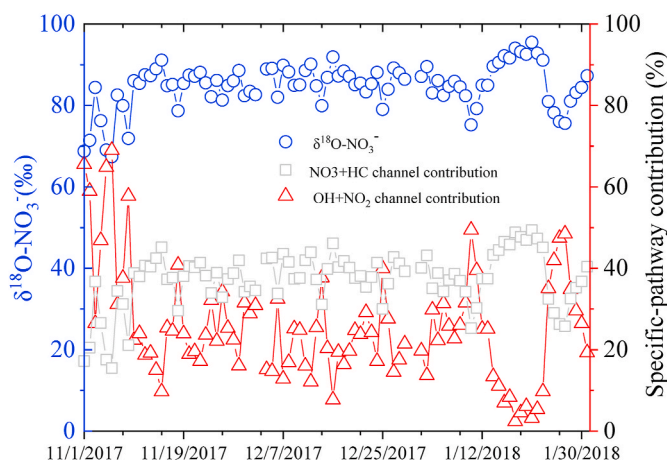


Fig. 7. Specific pathway contribution to nitrate formation during the campaign.

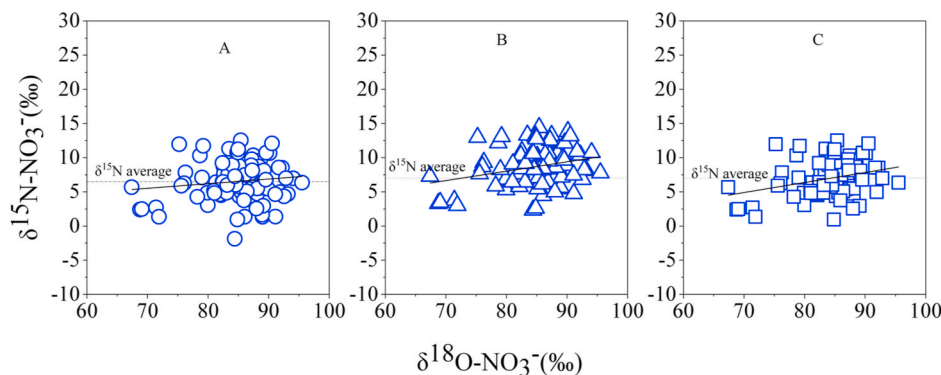


Fig. 6. Correlations between $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ values under different scenarios: A) original data; B) data corrected for Rayleigh effect; C) data with nitrate concentration higher than $30 \mu\text{g}/\text{m}^3$ were removed. Clearly, no correlations were observed between $\delta^{15}\text{N}\text{-NO}_3^-$ and $\delta^{18}\text{O}\text{-NO}_3^-$ values.

(He et al., 2018; Luo et al., 2019; Wang et al., 2017; Yan et al., 2019). The relative importance of N_2O_5 hydrolysis and NO_3+HC channel was further quantified according to the regression slopes of $\delta^{15}\text{N}-\text{NO}_3$ against $\delta^{18}\text{O}-\text{NO}_3$. As elucidated in Fig. 2, the NO_3+HC channel may contribute to the nocturnal nitrate accumulation in urban Nanchang in much the same level that N_2O_5 hydrolysis contributed, as $\delta^{15}\text{N}-\text{NO}_3$ did not correlate with $\delta^{18}\text{O}-\text{NO}_3$ (slope, therefore, assumed to be 0). The fractional contribution of the NO_3+HC channel was calculated to range from 14.2% to 45.5%, on average of 32.8% during the campaign, which emphasized the unignorable role of NO_3 radical chemistry in nitrate formation in winter Nanchang. In contrast, numerous researches suggested the minor role of the NO_3+HC channel in nitrate accumulation in the inland city (He et al., 2018; Michalski et al., 2003). For instance, Alexander et al. (2020) predicted that the NO_3+HC channel only accounted for 4–5% of the NO_3 accumulation globally (Alexander et al., 2019). Further, particulate NO_3 was generated only when NO_3 radical reacted with saturated hydrocarbons but the rate constants were relatively slow (e.g., in the order of $10^{-16} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$), while organic nitrate was produced when NO_3 radical reacted with unsaturated hydrocarbons (i.e., isoprene, rate constants was in the order of $10^{-13} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$), which may account for 20% of the total nitrates (Ng et al., 2008). Although, our calculation of the contribution of NO_3+HC channel (32.8%) was consistent with the corresponding results of specific-pathways contribution in winter Beijing ($34 \pm 10\%$, in winter 2014) estimated using $\Delta^{17}\text{O}-\text{NO}_3$ (Wang et al., 2019). Another possible reason for the high contribution of NO_3+HC channel in nitrate production was that the organic nitrate produced via NO_3 radical reacted with alkanes may be transformed into inorganic nitrate during the pretreatment processes or the organic nitrate was also converted into gaseous N_2O by the denitrifier bacteria (Lockwood et al., 2008). The hypothesis was based on the fact that the atmospheric organic oxidized nitrogen can be absorbed by foliage and incorporated into amino acids (Lockwood et al., 2008), however, no such test of whether the organic nitrogen can be chemically (or biologically) converted into inorganic one during routine laboratory pretreatment was conducted at present, to the best of our knowledge. Therefore, a solid conclusion on the specific-pathways contribution to NO_3 formation was hard to draw; however, our results may infer the non-negligible contribution of the NO_3+HC pathway in nitrate generation in the wintertime of urban Nanchang.

4. Conclusion

Particulate nitrate has been observed as a major component of $\text{PM}_{2.5}$ in the urban environment in China. A better understanding regarding the contribution of different chemical pathways to the production of NO_3 is therefore crucial for mitigation $\text{PM}_{2.5}$ pollution in the future. In the present study, daily $\text{PM}_{2.5}$ samples were collected in the wintertime of urban Nanchang and the related parameters (e.g., NO_3 concentration, dual isotopic signature, precursors concentrations) were determined to quantify the chemical conversion of NO_3 . Concentrations of nitrate ranged from $2.8 \mu\text{g}/\text{m}^3$ to $57.7 \mu\text{g}/\text{m}^3$, with a mean of $20.5 \pm 10.4 \mu\text{g}/\text{m}^3$, making the nitrate an important component that driven the haze appearance during the observation period. To distinguish the relative contribution of three major oxidation pathways, a simple but robust approach was used in this study. Briefly, the environmental significance of $\delta^{15}\text{N}-\delta^{18}\text{O}$ space of NO_3 was defined, which performed excellently in partition the relative importance of NO_3+HC and N_2O_5 hydrolysis channels. Taking full advantage of this approach, our results suggested the considerable contribution of NO_3+HC in nitrate production (average of 32.8%), even reached up to 45.5% on extremely polluted days. Although with a certain defect, our result provided an alternative perspective to better understanding nitrate chemistry in the polluted urban environment.

To date, explanations on the variability of $\delta^{15}\text{N}-\text{NO}_3$ is still less well-established. The ambient $\delta^{15}\text{N}-\text{NO}_3$ was usually regarded as a hybrid of

$\delta^{15}\text{N}$ of the source NO_x , the isotopic effect of the Leighton cycle, the fractionation effect associated with the oxidation reaction ($\text{NO}_x \rightarrow \text{HNO}_3$) and partition process ($\text{HNO}_3 \rightleftharpoons \text{NO}_3$), and possible the deposition influence. Our analysis suggested that certain neglected processes also significantly affected the signatures of $\delta^{15}\text{N}-\text{NO}_3$, e.g., the Rayleigh type fractionation during the rapid haze development period and the potential gas-to-particle exchange of atmospheric nitrate. Both the fractionation effects have been reported previously and may affect the source appointment of NO_x , to a certain extent. Furthermore, the interference of organic nitrates (RONO_2) to the isotopic composition of inorganic nitrate during laboratory analysis should be tested in the future. Nevertheless, $\delta^{15}\text{N}-\text{NO}_3$ was still a robust regional indicator for revealing the NO_x source.

CRediT author contribution statement

Zhongyi Zhang: Conceptualization, Software, Investigation, Writing – original draft. **Lin Cao:** and. **Yue Liang:** Methodology, Resources, Writing – review & editing, Data curation. **Hui Guan:** Software, Investigation, Writing – original draft. **Nengjian Zheng:** Conceptualization, Resources, Writing – review & editing, Data curation.

Declaration of competing interest

The authors declare that they have no conflicts of interest.

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Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.atmosenv.2021.118387>.

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