15.5 Sample Digestion Methods

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15.5.1	Introduction	87
15.5.2	General Considerations	88
15.5.2.1	Mineral Acids	88
15.5.2.1.1	Hydrofluoric acid	88
15.5.2.1.2	Nitric acid	88
15.5.2.1.3	Agua regia	88
15.5.2.1.4	Hydrochloric acid	89
15.5.2.1.5	Perchloric acid	89
15.5.2.1.6	Sulfuric acid	89
15.5.2.1.7	Phosphoric acid	89
15.5.2.2	Digestion Vessel Materials	89
15.5.2.3	Contamination from the Digestion Process	89
15.5.2.4	Assessing a Digestion Procedure	90
15.5.3	Sample Digestion Methods	91
15.5.3.1	Open Vessel Acid Digestions	91
15.5.3.2	Closed Vessel Digestions	92
15.5.3.3	Microwave Digestions	93
15.5.3.4	Partial Dissolutions	95
15.5.3.5	Dry Ashing Techniques	95
15.5.3.6	Alkali Fusions	97
15.5.3.6.1	Fusion with NaOH and KOH	97
15.5.3.6.2	Sinter or fusion with Na ₂ O ₂	97
15.5.3.6.3	Fusion with LiBO ₂ and $Li_2B_4O_7$	98
15.5.3.6.4	Fusion with Na ₂ CO ₃	98
15.5.3.7	Fire Assays	99
15.5.3.7.1	Lead fire assay	99
15.5.3.7.2	Nickel fire assay	99
15.5.3.8	Carius Tube and High-Pressure Asher (HPA-S)	100
15.5.3.9	Direct Fusion of Rock Powder	101
15.5.3.9.1	Flux-free fusion glasses	102
15.5.3.9.2	Lithium-borate fusion glasses	102
15.5.4 References	Summary and Overview	103 103

15.5.1 Introduction

Knowledge of the chemical and isotopic composition of samples is a prerequisite for understanding their characteristics. The major task of sample digestion is to convert the form of sample into one suitable for chemical analysis. Generally, after the sample is digested, the component of interest is in the solution as a soluble salt. The major advantage of the solution is its excellent homogeneity. It represents the original solution composition even in microliter volume. Despite a considerable amount of new research on sample introduction techniques, solution sampling remains the preferred method for most modern instrumental techniques, such as inductively coupled plasma mass spectrometry (ICP-MS)/inductively coupled plasma atomic emission spectrometry (ICP-AES). If the sample is a liquid or soluble in a liquid, then the sample dissolution is relatively simple. For example, water samples for which the further normal treatment necessary is to stabilize dissolved elements by the addition of small amounts of acid. However, most solid geologic samples cannot be dissolved in water, and sample digestion methods, such as acid digestion, fusion, or ashing, are required. Sample digestion is thus a fundamental and critical stage in the process of sample analysis, and it is often the limiting factor for sample throughput, especially with the rapid development of modern multielement measurement instrumentation (Chao and Sanzolone, 1992; Eggins et al., 1997; Jarvis, 1988; Longerich et al., 1990; Navarro et al., 2008).

The complexity of sample materials makes it necessary to choose a sample digestion method that is compatible with the specific objective of the analysis (Chao and Sanzolone, 1992). In reality, the choice of a digestion method should consider and be consistent with the chemical and physical composition and properties of the sample, the elements to be analyzed, the sample size required, the precision and accuracy desired, the

sample throughput needed, the suitability of the resulting digestion matrix for the analysis method, the apparatus and laboratory facilities available, the economic aspects including the reagent and labor consumption, and the safety considerations. There is an extensive literature on the various digestion methods for geologic and environmental samples (Chao and Sanzolone, 1992; Jarvis, 1992; Jeffery, 1975; Johnson and Maxwell, 1981; Matusiewicz, 2003; Nóbrega et al., 2006; Potts, 1987; Potts and Robinson, 2003; Šulcek and Povondra, 1989; Walsh et al., 1997). Despite numerous studies, the solution chemistry involved in many digestion methods is not well understood, and the knowledge in this area lags far behind developments in analytical instrumentation (Jarvis, 1992; Makishima et al., 2009; Tanaka et al., 2003; Yokoyama et al., 1999; Yu et al., 2001). Sample digestion thus remains a popular research theme for analytical geochemists.

This chapter provides a brief overview of digestion methods for the analysis of geologic samples. The authors do not attempt to discuss all dissolution methods; rather, many methods that are routinely used in the laboratory are presented.

15.5.2 General Considerations

15.5.2.1 Mineral Acids

Strong mineral acids (HNO₃, HCl, HF, HClO₄, H_2SO_4 , and H_3PO_4) are generally used for sample dissolution, and appropriate combinations of acids have been used successfully to decompose various geologic samples, such as rocks, minerals, soils, sediments, and ores (Chao and Sanzolone, 1992; Jarvis, 1992; Johnson and Maxwell, 1981; Potts, 1987; Potts and Robinson, 2003; Šulcek and Povondra, 1989). All of these acids are corrosive in nature, especially when heated and concentrated, and should be handled with extreme caution to prevent injury and accidents. Concentrated acids with the requisite high degree of purity are commercially available, and they can be further purified using subboiling distillation methods if needed (Yuan et al., 2000). The physical properties of the common mineral acids used in sample preparation are summarized in Table 1.

15.5.2.1.1 Hydrofluoric acid

Hydrofluoric acid (HF) is the most effective mineral acid for breaking up strong Si–O bonds to form SiF_6^{-2} ions in acidic solution. Silicates are converted to volatile SiF_4 , which will be lost in open vessel digestion procedures. HF by itself is more effective in the digestion of silicate rock minerals than when mixed with another acid. However, HF is rarely used as the sole reagent because some salts are poorly soluble in this acid

 Table 1
 General physical properties of common mineral acids

(Potts, 1987). HF is almost always mixed with other oxidizing acids such as HNO3 and/or HClO4 to ensure complete dissolution and to produce uniformly high oxidation states in the final solutions. Even diluted HF solutions will etch glass, so plastic labware (preferably PTFE or Teflon) is essential. It should also be noted that HF is one of the most hazardous mineral acids used in the laboratory, and it is both highly corrosive and toxic. Any HF spills on the skin should be immediately washed with copious cold water, and the affected area should be treated with a gel containing monosodium glutamate (Potts, 1987). Contact with HF does not cause an immediate burning sensation or pain but readily penetrates deep tissue and causes intense pain after an hour or more. HF will cause irreparable damage to the skin and eyes and should never be used without full safety precautions. Recently, NH₄F has been proposed to replace HF for the acid digestion of geologic and environmental samples (Hu et al., 2010; Mariet et al., 2008). A clear advantage of NH₄F-assisted acid digestion is that it does not require handling the very corrosive and toxic HF. Notably, the conventional subboiling purification procedure is not effective at removing As impurities in HF. This limit may be related to the presence of volatile As species such as AsF₃ (boiling point of 63 °C). To remove this volatile As species, it is recommended to employ a boiling procedure for hydrofluoric acid prior to the conventional subboiling purification procedure (Hu et al., 2005).

15.5.2.1.2 Nitric acid

Nitric acid (HNO₃) is one of the most widely used digestion reagents and the most widely used primary oxidant for the decomposition of organic matter. Hot and concentrated HNO₃ (16 M and 68%) is a strong oxidizing agent that will liberate trace elements from many materials as highly soluble nitrate salts. The oxidizing properties of nitric acid are lost when it is diluted below approximately 2 M. The most important application of HNO3 in rock analysis is to decompose both carbonate and sulfide minerals (usually in association with HCl). Nitric acid matrices are the best acid medium for ICP-MS analysis. Its constituents (H₂, N₂, and O₂) are already present in air entrained by the plasma, and the range of polyatomic ions are not increased significantly by the addition of an HNO3 matrix (Gray, 1986; Tan and Horlick, 1986), and it also does not interfere with most determinations. Additionally, nitric acid is available commercially in sufficient purity.

15.5.2.1.3 Aqua regia

A fresh mixture of concentrated HNO_3 (16 M) and HCl (12 M) in a volume ratio of 1:3 produces a particularly useful reagent known as aqua regia. This reagent has much stronger oxidizing

Mineral acids	Formula	Concentration (%)	Molarity (M)	Density (kg I^{-1})	Boiling point (°C)	Comments		
Nitric acid	HNO ₃	68	16	1.42	122	68% HNO ₃ , azeotrope		
Hydrochloric acid	HCI	36	12	1.19	110	20.4% HCl, azeotrope		
Hydrofluoric acid	HF	48	29	1.16	112	38.3% HF, azeotrope		
Perchloric acid	HCIO₄	70	12	1.67	203	72.4% HClO ₄ , azeotrope		
Sulfuric acid	H_2SO_4	98	18	1.84	338	98.3% H ₂ SO ₄		
Phosphoric acid	H ₃ PO ₄	85	15	1.71	213	Decomposes to HPO ₃		

and dissolving powers than HNO_3 alone. The effectiveness of this reagent is most likely due to the complexing power of the Cl^- combined with the catalytic effects of free Cl_2 and NOCl (Jarvis, 1992). This reagent should be freshly prepared before use. Aqua regia is mainly used for the decomposition of metals, alloys, sulfides, silicates (in association with HF), and other ores and is well known for its ability to dissolve Au, Pt, and Pd.

15.5.2.1.4 Hydrochloric acid

Concentrated hydrochloric acid (HCl) is the most frequently used halogen acid for the dissolution of geologic samples. Unlike HNO3, HCl is a weak reducing acid and is not generally used to digest organic materials. It is an excellent solvent for carbonates, phosphates, many metal oxides, and metals. For example, due to its reducing properties and the complexing ability of Cl⁻, HCl is a better solvent for dissolving iron and manganese oxides than HNO3. For silicate analysis, HCl is generally used in combination with other acids, such as HF and HNO3, although some basic silicate minerals can be completely or partially decomposed by HCl alone. At elevated temperatures and pressures, many silicates and other refractory oxides, sulfates, and fluorides are attacked by HCl to produce soluble salts. HCl is the preferred acid medium to dissolve residues that remain after acid digestion or melts of alkali fusion for later analysis using atomic absorption spectrometry (AAS) (Potts, 1987). Unlike atomic absorption techniques, HCl is not a suitable sample matrix for ICP-MS analysis because chloride-bearing polyatomic ions cause major interferences (e.g., ArCl, ClO, and ClOH) with As and V (75As and ⁵¹V) and with many other trace elements (Cr, Fe, Ga, Ge, Se, Ti, and Zn) to a lesser extent (Jarvis, 1992). Hydrochloric acid can be effectively removed from sample solutions by repeated evaporation to incipient dryness with HNO3 because the boiling point of the HCl azeotrope (110 °C) is below that of the HNO₃ azeotrope (122 °C). However, it should be noted that there may be potential losses of the volatile metal chlorides (As, Sb, Sn, Se, Ge, and Hg) if HCl is used in acid digestion procedures.

15.5.2.1.5 Perchloric acid

Perchloric acid (HClO₄) is one of the strongest mineral acids. Hot and concentrated perchloric acid has powerful oxidizing and dehydrating properties, and it will react explosively with organic compounds. For this reason, it is best to pretreat geologic samples containing organic material or organic samples with HNO3 or an HNO3-HClO4 mixture. Perchlorate salts are generally highly soluble and stable in aqueous solutions, but some alkali (K, Rb, and Cs) perchlorates are exceptions to this general rule. The high boiling point of the acid ensures a more efficient attack of refractory minerals by improving the efficiency of HF and the more complete removal of HF during evaporation stages. Unlike HCl, chlorine ions introduced during the digestion procedure in the form of HClO₄ are difficult to be removed by evaporation and will be harmful to the determination of low levels of As and V in ICP-MS (Jarvis, 1992). Some salts of HClO₄ are spontaneously flammable in the anhydrous form. Therefore, HClO₄ must be used in a specially designed hood that is equipped with washdown facilities for cleaning after use (Chao and Sanzolone, 1992; Potts, 1987).

15.5.2.1.6 Sulfuric acid

Concentrated sulfuric acid (H₂SO₄) has dehydrating and mildly oxidizing abilities and the highest boiling point (338 °C for the 98.3% acid) of the mineral acids. Cold concentrated sulfuric acid has an extremely high affinity for water, and it will produce a significant amount of heat when diluted. Therefore, this acid must always be added slowly to excess water and not vice versa. Sulfuric acid is a highly effective reagent in combination with HF for the decomposition of most resistant minerals, such as zircon, chromite, monazite, cryolite, and many naturally occurring fluorides (Potts and Robinson, 2003; Yu et al., 2001). Unfortunately, some inorganic sulfates have low solubilities (e.g., Ba, Ca, Pb, and Sr), and volatilization of trace elements (Ag, As, Ge, Hg, Re, and Se) may occur during the digestion of some samples (Šulcek and Povondra, 1989). Furthermore, H₂SO₄ is very difficult to remove by evaporation (several days) due to its high boiling point (338 °C). The viscosity of this acid results in transport effects during sample introduction in ICP-MS. This acid also causes severe sulfur polyatomic ion interferences and attacks the nickel sampler cones in the ICP-MS instrument. For these reasons, H₂SO₄ has not been widely used for the decomposition of geologic samples.

15.5.2.1.7 Phosphoric acid

Phosphoric acid (H_3PO_4) is somewhat limited in its use in analytical chemistry because phosphate ions can cause interferences by complexing or precipitating some of the elements to be analyzed (Chao and Sanzolone, 1992). Possible difficulties encountered in ICP-MS analysis are the presence of polyatomic species of P, transport effects induced by the high viscosity of the acid, and rapid erosion of the nickel sampler cone. H_3PO_4 undergoes a series of condensation reactions to form condensed phosphoric acid upon heating. Condensed phosphoric acid alone or with HClO₄ can decompose 70 natural minerals among sulfides, oxides, silicates, and carbonates (Hannaker and Hou, 1984). Further investigations are needed to explore the full potential of H_3PO_4 as an acid decomposition agent.

15.5.2.2 Digestion Vessel Materials

Many different vessel materials are used to handle samples during sample digestion in laboratories. The vessel material must be chosen carefully according to its nature (e.g., resistance to acids and alkalies, heat resistance and conductance, surface properties, reactivity, mechanical strength, and contamination), the sample components to be analyzed, and the analytical requirements. **Table 2** lists the preferred vessel materials for sample digestion. These materials have gained considerable popularity for use in the handling of samples.

15.5.2.3 Contamination from the Digestion Process

The influence of contamination on the analytical results becomes increasingly important with decreasing concentrations of the analyte. Modern analytical methods and instrumentation make possible the measurement of extremely low concentrations of elements in complex matrices. In many

Materials	Maximum temperature (°C)	Comments
Borosilicate glass	800	Resistant to most acids, but should not be used with HF or boiling H_3PO_4 or alkaline solutions
Porcelain	1100	Popular material used for ashing purpose
Quartz	1200	The most suitable material for the wet digestion of organic materials
Platinum	1500	Resistant to attack by most acids and fusion reagents. Heats up and cools down rapidly, making it excellent for ash determinations
Glassy carbon	500	An inexpensive material for alkaline fusions with low melting point agents
Polytetrafluorethylene (PTFE)	250	Generally used for closed digestion vessels
Perfluoroalkoxy (PFA)	250	Commonly used in microwave digestion vessels
High-density polyethylene (HDPE)	120	Typically used for containment of the diluted sample digestion solution
Low-density polyethylene (LDPE)	80	
Polypropylene (PP)	135	

Table 2	Preferred	vessel	materials	for	sample	digestion
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cases, the blank level determines the lower limit of detection, which has increased the importance of digesting samples in a manner that keeps them as contamination-free as possible.

Contamination can occur from the reagents, the vessel materials, and the environment during the digestion process and is often an unforeseen barrier in sample analyses that can lead to false results. Liquid reagents, such as water and acids, are most important for sample preparation, and they are generally available in high-purity grades. If required, these reagents also can be further purified using subboiling distillation. In contrast, solid reagents are difficult to purify and result in comparatively high blank levels. Many different vessel materials have been used to handle samples during sample preparation. Contaminants can be desorbed from impurities on the surfaces of the vessels or leached out from the vessel materials. It has been reported that the degree of contamination from the commonly used vessel materials is in the following order: polyethylene (low density) < fluorocarbons (e.g., Teflon, PTFE, and Tefzel) < quartz (synthetic) < polyethylene (high density) < quartz (natural) < platinum < borosilicate (Gaines, 2012; Moody and Lindstrom, 1977; Murphy, 1974). Although containers made of linear polyethylene or Teflon introduce the least amount of contamination for most trace element analysis, these containers should be carefully cleaned with HCl and HNO₃ before use (Moody and Lindstrom, 1977). Simple acid washing of the bottles prior to use does not solve the contamination problem for all elements. For example, Ta contamination from perfluoroalkoxy (PFA) Teflon® vessels (Makishima et al., 1999), Ba, Zn, and Cr contamination from high-density polyethylene bottles (Reimann et al., 1999, 2007) and Sb contamination from polyethylene terephthalate bottles (Shotyk et al., 2006) occur via leaching of the contaminants from the containers. Reimann et al. (2010) reported that dark-colored containers leach more materials than clear containers do for most elements, and this observation is independent of the container material. Therefore, great care is required when choosing containers, reagents, and the apparatuses used for sample digestion. Environmental contamination is caused by airborne particles and gaseous matter. To keep contamination risks low during chemical treatment in open systems, work should be performed in clean rooms equipped with laminar-flow clean benches (Gaines, 2012; Knapp and Schramel, 2003; Tschoepel and Toelg, 1982; this volume, Chapter 15.6).

15.5.2.4 Assessing a Digestion Procedure

Quality assurance, which is aimed at obtaining the correct result, is extremely important in analytical chemistry. Long-term experience has shown that precise sample digestion is the most critical part of an analysis because it is responsible for the largest and most often hidden sources of errors (Hoenig, 2001). For quality control in sample digestion, it is necessary to measure and record certain parameters exactly to be able to subsequently trace the course of the digestion process (Matusiewicz, 2003). Precision and accuracy are the two most important parameters for the quality control of sample digestion. Precision is defined as an estimation of the reproducibility of a sample digestion method, which is normally calculated from the standard deviation of the analyses of duplicate digestions of the same sample. Accuracy refers to a measure of how closely the analytical data agree with the 'true' composition of the sample. For the sample digestion procedures, the use of an adequate certified reference material (CRM) is strongly recommended for quality assurance and for method development. A CRM is a substance for which one or more analytes have certified values, which are produced by a technically valid procedure and are accompanied with a traceable certificate and issued by an appropriate certifying agency (see Chapter 15.3 in this volume). CRMs are available from many distributors and for a wide range of matrices (see Chapter 15.3 in this volume). Selecting an appropriate test portion of the sample materials before digestion is also an important part of quality assurance. Since the 1980s, when using a particle size level of 200 mesh (74 µm), it has been recommended to use 100 mg as a minimum test portion size to assure representative subsampling, which is a compromise between sample size and detection limits. However, for modern high-sensitivity analytical techniques, such as ICP-MS, the practical sampling mass is much less than 100 mg. Thus, a significant challenge to use the minimum sampling mass of sample materials has emerged with the development of modern analytical techniques (Wang et al., 2004). Similarly, the added amount of digestion reagents should be appropriated for the selected test portion of the sample materials. For example, 1-2 ml of HF and 0.5-1 ml of HNO3 are typically used for the digestion of 50-100 mg silicate materials in closed vessels. In addition, procedural blanks should always be prepared for each batch of samples, which are very helpful for tracing the sources of errors.

15.5.3 Sample Digestion Methods

15.5.3.1 Open Vessel Acid Digestions

Open vessel acid digestions have long been a popular and simple method for the digestion of inorganic and organic sample materials in chemical laboratories. It refers to acid attack in open containers or screw-top vials (low pressure) placed on a hot plate. This method is an extremely important value for routine analysis because of its flexibility in the control of digestion parameters, such as temperature, time, and the addition of reagents. However, the maximum digestion temperatures are limited by the ambient-pressure boiling point of the corresponding acid or acid mixture. Other disadvantages of open vessel digestions are the potential risk of contamination from the laboratory air and the large amounts of reagents that are required and the potential loss of trace elements. Therefore, open vessel acid digestion has not been considered to be a state-of-the-art technology in trace and ultratrace sample preparation (Matusiewicz, 2003).

For the digestion of geologic samples, HF in combination with other mineral acids (most commonly nitric and perchloric acids) in an open vessel is routinely used (Chao and Sanzolone, 1992; Jarvis, 1990, 1992; Jenner et al., 1990; Johnson and Maxwell, 1981; Potts, 1987; Potts and Robinson, 2003; Robinson et al., 1999; Šulcek and Povondra, 1989). Silicon is quantitatively removed as the volatile silicon tetrafluoride SiF4, which reduces the total amount of dissolved solid in the solutions and makes the solution far more stable because Si in the solution tends to hydrolyze and precipitate to form undissolved polysilicic acid. Evaporation using a mineral acid in addition to hydrofluoric acid also facilitates the removal of fluorine ions and ensures that insoluble fluorides are converted to more soluble salts. For example, the high boiling point of HClO₄ (203 °C) compared with that of HF (112 °C) ensures the complete removal of fluorides and the more efficient attack of refractory minerals due to the increased boiling temperature of the reaction mixture and the high solubility of perchlorates (Potts and Robinson, 2003; Totland et al., 1992). The order of effectiveness in removing residual amounts of fluorine decreases in the order of sulfuric acid, perchloric acid, and nitric acid. Langmyhr (1967) reported that double evaporation with $HClO_4$ at 180 °C gives about the same trace residue of fluoride as one evaporation with H₂SO₄ at 250 °C. However, as mentioned in Section 15.5.2.1, HClO₄ is preferred to H₂SO₄ in analytical practice. The choice of the accompanying acid or the optimum combination of acid mixtures is determined by many factors, such as the nature of the sample, the elements to be analyzed, and the suitability of the resulting digest matrix for the analysis method. There is also some disagreement in the literature regarding the optimum combination of acids and the role of the second acid (Langmyhr and Sveen, 1965; Tang et al., 1992). For example, Langmyhr and Sveen (1965) reported that hydrofluoric acid alone is more effective as a decomposition agent for the digestion of minerals than a mixture containing HF with another mineral acid. Tang et al. (1992) analyzed 11 Chinese geologic reference materials and found no significant difference between HF/HNO₃/HClO₄, HF/aqua regia/HClO₄, and HF/ HClO₄ in an open acid digestion.

During sample decomposition using HF, stable insoluble Mg–Ca–Al fluorides are usually produced, which differentially incorporate trace elements into their lattices during

precipitation and lead to negative errors in the analysis (Boer et al., 1993; Croudace, 1980; Langmyhr and Kringstad, 1966; Makishima et al., 2009; Yokoyama et al., 1999). Not only do insoluble fluorides form with elements, such as rare earth elements (REEs) and Th, but some soluble elements, such as Rb, Cs, and U, also coprecipitate with these fluorides (Boer et al., 1993; Hu et al., 2010; Yokoyama et al., 1999). It is clear that the analysis of trace elements in the presence of insoluble fluorides will yield erroneous results and misleading conclusions, especially in terms of Rb-Sr, La-Ce, Sm-Nd, U-Th-Pb systematics, or REE patterns. To remove these fluorides, repeated evaporation of sample solutions with HClO₄ has been commonly used (Langmyhr, 1967; Yokoyama et al., 1999). Yokoyama et al. (1999) obtained 100% recovery of the incompatible elements Rb, Sr, Y, Cs, Ba, REEs, Pb, Th, and U by using larger amounts of HClO₄ and evaporating the sample to dryness in a stepwise fashion. However, the highfield strength elements (HFSE), such as Ti, Zr, Nb, Hf, and Ta, were lost as insoluble oxides. The complete removal of fluoride from sample solutions is not always straightforward and is often difficult to accomplish (Chao and Sanzolone, 1992; Croudace, 1980; Makishima et al., 2009; Takei et al., 2001; Tanaka et al., 2003). For example, the presence of ralstonite (Boer et al., 1993; Croudace, 1980), NaUZr₂F₁₂ (Boer et al., 1993), and AlF₃ (formed at 205 °C in the bombs) (Takei et al., 2001) remains a problem because these phases do not dissolve in acids. Therefore, it is very important to suppress the formation of insoluble fluorides. To prevent the formation of insoluble fluorides, the use of small quantities of a sample (not exceeding 100 mg) is recommended (Hu et al., 2010; Potts and Robinson, 2003; Yokoyama et al., 1999; Zhang et al., 2012). It has also been proposed that during sample digestion, treated powders, after acid attack, should not be allowed to dry completely, a step that apparently can minimize the formation of insoluble compounds (Croudace, 1980; Dulski, 2001; Navarro et al., 2008). Takei et al. (2001) and Tanaka et al. (2003) found that the bulk rock proportions of Ca:Al:Mg have a significant effect on the formation of insoluble fluorides. They have developed the novel 'Al-addition' and 'Mg-addition' methods to overcome these problems.

The main drawbacks of open vessel acid digestions for the digestions of geologic samples are the incomplete dissolution of refractory minerals, such as chromite, corundum, cassiterite, spinel, garnet, magnetite, monazite, rutile, and zircon, and the potential loss of some elements (e.g., As, B, Ge, Sb, Se, Hg, Te, Re, Os, and Ru) (Chao and Sanzolone, 1992; Hall and plant, 1992; Potts, 1987; Révillon and Hureau-Mazaudier, 2009; Roy et al., 2007; Šulcek and Povondra, 1989; Walsh et al., 1997; Yu et al., 2001). Elements such as Zr, Hf, Cr, Sn, Mo, Y, Ba, and HREEs are often present in these refractory minerals, and their recovery may be low (Hall and plant, 1992; Révillon and Hureau-Mazaudier, 2009; Roy et al., 2007; Sulcek and Povondra, 1989; Tang et al., 1992). To overcome this drawback, the residue that remains after open acid digestion is sometimes filtered and decomposed using alkali fusion or closed vessel digestion (Dupta, 1994; Jarvis and Jarvis, 1988; Pin and Joannon, 1997). However, the lack of ultrapure reagents precludes the use of alkali fusion methods for the analysis of geologic samples with low trace abundances. Recently, a new method using a Siliconit furnace was developed to accomplish the

flux-free fusion of silicate rock prior to HF–HClO₄ digestion in a Teflon beaker in preparation for bulk analysis using ICP-MS (Shimizu et al., 2011). This method is especially effective for felsic samples that contain refractory minerals, such as zircon and tourmaline.

Open vessel acid digestions are a well-established method incorporating the sample into the solution and are very successful for samples without refractory minerals. However, the digestion recipe used by each laboratory varies. **Table 3** lists a selection of typical sample treatment procedures for open vessel acid digestions.

15.5.3.2 Closed Vessel Digestions

An important advance in acid dissolution procedures was made with the introduction of the closed acid digestion bomb that was popularized by Langmyhr and Paus (1968) and Bernas (1968). The complete digestion of a sample is a prerequisite to achieve reproducible and accurate results in routine analysis. Digestions performed in closed vessel devices benefit from the synergic effects of temperature and pressure, in which digestions reach higher temperatures because the boiling point of the reagents is raised by the pressure generated within the vessel. The principal advantages of digestion in a closed system can be summarized as follows (Chao and Sanzolone, 1992; Jackwerth and Gomiscek, 1984; Jarvis, 1992; Šulcek and Povondra, 1989):

- This method enables the rapid digestion of refractory phases, which are not decomposed or are only partially decomposed in an open vessel.
- 2. The digestion time is significantly shortened.
- 3. Smaller reagent volumes are required.
- 4. Losses of volatile elements, such as As, B, Cr, Cd, Hg, Sb, Se, Sn, and Pb, are prevented.
- Contamination is reduced by lowering the reagent volumes and excluding the possible introduction of airborne particles during the decomposition.

Due to these advantages, closed vessel digestions are particularly suitable for trace and ultratrace analyses, especially when the supply of the sample is limited. However, if silicon is to be determined, boric acid is normally added to complex the excess

 Table 3
 A selection of typical sample treatment procedures for open vessel acid digestions

Material	Determined elements and instruments	Sample treatment	References
Acid, intermediate and basic rocks, sediments, shales, limestones	Y, Zr, Nb, Hf, Ta, Pb, Th, U, REEs (ICP-MS)	(1) A 0.2 g aliquot of the sample is placed in a beaker and moistened with 2 ml of water. (2) After the addition of 10 ml of HF and 8 ml of HClO ₄ , the sample is evaporated to dryness at 200 °C. (3) Step 2 is typically repeated 3–4 times. (4) An additional 5 ml of HClO ₄ is added, and the sample is evaporated to dryness. (5) The sample is heated for 10 min following the addition of 10 ml of 5N HNO ₃ . (6) The sample is transferred to 100 ml volumetric flasks	Jarvis (1990)
Serpentinite, limestone, dolomite, andesite, basalt, granite, shale, syenite	Ti, Al, Fe, Mg, Mn, Ca, Ba, Cr, Cu, Ni, Sr, V, Y, Zn, Zr (ICP-AES); Ba, Be, Co, Cr, Cs, Cu, Hf, Mo, Nb, Ni, Pb, Rb, Sb, Sc, Sn, Sr, Ta, Th, Tl, U, W, Y, Zn, Zr, REEs (ICP-MS)	(1) An aliquot of 0.500 g of the sample is placed in a beaker and moistened with a few milliliters of water. (2) Next, 10 ml of HF and 4 ml of HClO ₄ are added, and the sample is evaporated to dryness at 200 °C. (3) Step 2 is typically repeated 3–4 times. (4) The sample is evaporated to dryness after the addition of 4 ml of HClO ₄ . (5) A 10 ml portion of 5N HNO ₃ is added, and the sample is heated gently. (6) The sample is transferred to 50 ml volumetric flasks	Totland et al. (1992)
A range of basalt, granite, and ironstone geologic reference materials	Sc, Rb, Y, Zr, Nb, Mo, Sn, Cs, Ba, REEs, Hf, Ta, Pb, Th, U (ICP-MS)	(1) A 100 mg aliquot of the sample is placed in Teflon vials and wetted with a few drops of water. (2) The addition of 2 ml of HF and 0.5 ml of HNO ₃ is followed by digestion at 130–150 °C for 48 h and evaporation to dryness. During the digestion, the vial is removed from the hot plate twice and placed in an ultrasonic bath for a couple of minutes. (3) The evaporation is repeated twice more after adding 1 ml of HNO ₃ each time. (4) Next, 2 ml of HNO ₃ and 3–5 ml of water are added. (5) The sample is then diluted to 100 ml with water	Robinson et al. (1999) and Yu et al. (2001)
Stream and lake sediment	Sc, Ti, V, Cr, Mn, Co, Ni, Cu, Ga, Zn, Rb, Y, Zr, Nb, Cs, Ba, REEs, Hf, Ta, Pb, Th, U (HR-ICP-MS)	(1) A 100 mg aliquot of the sample is placed in PFA Teflon beakers. (2) After the addition of 3 ml of HF and 1 ml of HNO ₃ , the sample is agitated for 20 min in an ultrasonic bath. (3) The sample is then digested at 130 °C for 48 h and evaporated to dryness at 80–120 °C. (4) The sample is mixed with 1 ml of HCl, heated for 12 h, and evaporated to dryness. (5) The dried residue is then dissolved in another solution of 40 ml of HCl (3 mol l^{-1})	Révillon and Hureau-Mazaudier (2009)

fluoride. The addition of boric acid permits the retention of Si and other volatile fluorides and prevents the precipitation of insoluble fluorides (Bernas, 1968; Langmyhr and Paus, 1968; Van Eenbergen and Bruninx, 1978). Because the resulting solution still attacks glassware, it should not be aspirated into the conventional sample introduction systems of ICP-MS/ICP-AES. In addition, the increased amount of dissolved solids significantly hinders the analysis of trace elements. Thus, the application of this technique is limited in routine sample analysis. For the analysis of trace elements, it is better to remove silicate as the volatile SiF₄ on a hot plate after opening the vessel.

Many types of vessels are available for sample digestion in a closed system. The different vessels have different temperature limits. Polypropylene and polycarbonate bottles can be used up to 130 °C, and PTFE-lined pressure vessels can withstand temperatures in the range of 150-250 °C (Bernas, 1968; Langmyhr and Paus, 1968). Decompositions in closed systems at elevated temperatures started to develop successfully only after the introduction of fluorinated hydrocarbon polymers (PTFE and Teflon[®]) as the materials for the autoclave internal vessels (Sulcek and Povondra, 1989). Modern conventional acid digestion bombs consist of a PTFE beaker with a lid, which fits tightly into an outer stainless-steel pressure jacket. The outer jacket has a screw-top lid, which, when tightened, forms a gas-tight highpressure seal between the beaker and its lid. These bombs produce very high pressures (7-12 MPa) when the sample and acids are subjected to high temperatures (110-250 °C) (Jarvis, 1992; Potts and Robinson, 2003). At any temperature, the pressure developed within the closed vessel depends on the partial pressure of the solvents under nonequilibrium conditions and the gaseous decomposition products formed during the decomposition. To avoid explosive rupturing of the vessels, organic material should not be mixed with strong oxidizing agents in closed vessels, and samples and reagents should never constitute more than 20% of the total vessel volume. It is also essential that the closed vessel is not opened until it has completely returned to room temperature.

Closed vessel acid digestion may result in the decomposition of many refractory minerals and compounds that are not decomposed or are only partially decomposed in an open vessel (Bernas, 1968; Hu and Gao, 2008; Langmyhr and sveen, 1965; Maqueda and Rodriquez, 1986; Olive et al., 2001; Pretorius et al., 2006; Qi and Grégoire, 2000; Qi et al., 2000). Beryl, chalcopyrite, kyanite, pyrite, pyrrhotite, and staurolite are only partially decomposed at 95 °C in an open vessel by HF/HClO₄, but they are completely dissolved in a Teflonlined bomb at 250 °C (Langmyhr and Sveen, 1965). Topaz is a significant exception, which may require alkali fusion to be brought into the solution. Pyrophyllite is not attacked by HF/ HClO₄/HNO₃ in an open vessel, but it is dissolved completely in PTFE-lined sealed vessel (Maqueda and Rodriquez, 1986). Refractory compounds, such as titanium dioxide, niobium dioxide, and silicate nitride, were also successfully decomposed at 170 °C using closed vessel digestion (Bernas, 1968). Felsic rocks are known to be very difficult to dissolve because of the presence of refractory minerals, such as zircon, tourmaline, magnetite, and monazite. Recent results show that felsic rocks, such as granites and granodiorites, can be completely decomposed using closed vessel acid digestion at the temperatures of 190-200 °C for 12 h to several days (Hu et al., 2010; Pretorius et al., 2006; Qi et al., 2000; Yu et al., 2001). Currently, closed vessel acid digestion is one of the most successful and popular sample preparation techniques, and it has been widely used to decompose various geologic samples (e.g., felsic rocks, ultramafic rocks, sediments, and soils) for multielement determinations, such as REEs and HFSE (Begum et al., 2007; Diegor et al., 2001; Dulski, 2001; Hu and Gao, 2008; Hu et al., 2010; Navarro et al., 2008; Olive et al., 2001; Pretorius et al., 2006; Qi and Grégoire, 2000; Qi et al., 2000; Révillon and Hureau-Mazaudier, 2009; Yu et al., 2000, 2001; Zhang et al., 2012).

93

To decompose silicate materials, acid mixtures such as HF/HNO₃ (Diegor et al., 2001; Hu and Gao, 2008; Navarro et al., 2008; Qi and Grégoire, 2000; Qi et al., 2000; Zhang et al., 2012), HF/HClO₄ (Dulski, 2001; Olive et al., 2001; Yu et al., 2000, 2001), HF/H₂SO₄ (Münker, 1998; Yu et al., 2001), HF/ HNO₃/HCl (Roy et al., 2007), and HF/HClO₄/HNO₃ (Pretorius et al., 2006; Révillon and Hureau-Mazaudier, 2009; Xie and Kerrich, 1995) have been used extensively in closed vessel digestions for multielement analysis. Different reliable digestion procedures have been described by these authors. Regardless of the different acid mixture used, the final sample uptake is routinely performed with HNO3 in most geochemical laboratories. Sometimes, in order to stabilize HFSE in the solution, a final sample uptake procedure using HCl or HNO₃/HCl has also been proposed (Dulski, 2001; Münker, 1998; Yu et al., 2000, 2001). The choice of acid mixtures is critical for sample digestion in some circumstance. For example, the use of HF/ HClO₄ is preferred to the use of HF/HNO₃ or HF/HNO₃/HCl for trace elements in ultramafic rocks in which spinel-group minerals are effectively decomposed (Olive et al., 2001; Potts and Robinson, 2003). In contrast to HF/HNO3 digestion in high-pressure bombs, the use of HF/HClO₄/HNO₃ has resulted in superior recoveries and better precision for the bulk of the trace elements analyzed in felsic rocks (Pretorius et al., 2006). Yu et al. (2001) reported that HF/H_2SO_4 is more powerful for digesting granites and ironstones than HF/HClO₄. However, the high boiling point of H₂SO₄ (338 °C) dictates a minimum of 4 days of evaporation time (Yu et al., 2001). Ito (1962) reported that HF is more powerful for digesting zircon than HF/H₂SO₄. Zhang et al. (2012) reported that the addition of HNO₃ inhibited the digestion capabilities of HF for zircon. Table 4 lists a selection of typical sample treatment procedures for closed vessel acid digestion.

15.5.3.3 Microwave Digestions

In 1975, Abu-Samra et al. (1975) described one of the first uses of microwave heating for the rapid wet acid digestion of biological materials. This discovery stimulated the long-term development of microwave technology for the preparation of all types of samples for analysis (Bettinelli et al., 1989; Chao and Sanzolone, 1992; Chen et al., 2008; Kingston and Haswell, 1997; Kingston and Jassie, 1988; Kuss, 1992; Lamble and Hill, 1998; Matusiewicz and Sturgeon, 1989; Nadkarni, 1984; Smith and Arsenault, 1996; Srogi, 2006; Suzuki and Sensui, 1991). Currently, microwave technology is being applied not only in analytical chemistry but also in organic synthesis, inorganic reactions, preparation of catalysts, and other fields (Chen et al., 2008; Kingston and Haswell, 1997; Kingston and Jassie, 1988). This technology has now advanced to the point where it is

Material	Determined elements and instruments	Sample treatment	References	
Granite, andesite, basalt, sediments, sandstone, shale, limestone, and soil	Li, Sc, V, Cr, Co, Ni, Cu, Zn, Ga, Rb, Sr, Y, Zr, Nb, Mo, Cs, Ba, REEs, Hf, Ta, W, Pb, Th, U (HR-ICP-MS)	(1) A 100 mg aliquot of the sample is placed in a PTFE bomb. (2) Next, 1 ml of HF and 0.5 ml of HNO ₃ are added, and the sample is evaporated to dryness on a hot plate. (3) The addition of 1 ml of HF and 0.5 ml of HNO ₃ is followed by heating to 200 °C for 12 h under pressure. (4) After cooling, 1 ml of a 1 μ g ml ⁻¹ Rh solution is added as an internal standard, and the sample is evaporated to dryness at 150 °C. (5) Next, 1 ml of HNO ₃ is added, and the sample is evaporated to dryness at 150 °C. (5) Next, 1 ml of HNO ₃ is added, and the sample is evaporated to dryness. (6) Repeat step 5 one time. (7) Approximately 6–8 ml of 40% v/v HNO ₃ is added, and the bombs are resealed and heated to 110–140 °C for 3 h. (8) The sample is then diluted to 100 ml with ultranure water	Qi et al. (2000) and Qi and Grégoire (2000)	
Ninety reference materials of different matrix compositions	Rb, Sr, Y, Zr, Cs, Ba, REEs, Hf, Pb, Th, U (ICP-MS)	(1) A 100 mg aliquot of the sample is placed in a PTFE bomb and wetted with a few drops of water. (2) Next, 3 ml of HF and 3 ml of HClO ₄ are added, and the sample is heated to 180 °C for 16 h under pressure. (3) After cooling, the sample is evaporated to near dryness at 180 °C. (4) A 5 ml portion of HCl is then added, and the sample is evaporated to incipient dryness. (5) After the addition of 5 ml of HCl, the sample is heated to 130 °C for 12 h under pressure. (6) After cooling, the sample is evaporated to near dryness and redissolved in 2 ml of HCl and 10 ml of water. (7) The sample is then diluted to 50 ml with ultranure water	Dulski (2001)	
Granite, granodiorite	Li, Sc, V, Co, Ni, Cu, Zn, Ga, Rb, Sr, Y, Zr, Nb, Mo, Cd, Sn, Sb, Cs, Ba, REEs, Hf, Ta, W, Pb, Bi, Th, U (HR-ICP-MS)	(1) A 100 mg aliquot of the sample is placed in a PTFE bomb. (2) Next, 5 ml of HF, 0.714 ml of HClO ₄ , and 0.714 ml of HNO ₃ are added, and the sample is heated to 190 °C for 5 days under pressure. (3) After cooling, the sample is evaporated to dryness, redissolved in 6 ml of 6 mol I^{-1} HCl, and heated to 190 °C for 24 h. (4) The sample is then transferred to a Savillex [®] vial and evaporated to dryness. (5) The sample is placed on a hot plate for 24 h after the addition of 1 ml of HNO ₃ . (6) The sample is then evaporated to dryness and stored until dilution for analysis	Pretorius et al. (2006)	
Leucogranite, diorite, granodiorite, granite, andesite, basalt, sandstone, shale, limestone, graywacke, pelite, and loess	Li, Be, B, Sc, V, Cr, Co, Ni, Cu, Zn, Ga, Ge, As, Rb, Sr, Y, Zr, Nb, Mo, Cd, In, Sn, Sb, Te, Cs, Ba, REEs, Hf, Ta, W, TI, Pb, Bi, Th, U (ICP-MS)	(1) A 50 mg aliquot of the sample is placed in a PTFE bomb. (2) Next, 1 ml of HNO ₃ and 1 ml of HF are added, and the sample is heated to 190 °C for 48 h under pressure. (3) After cooling, the bomb is opened, and the sample is evaporated to incipient dryness at 120– 150 °C. (4) After the addition of 1 ml of HNO ₃ , the sample is evaporated to dryness. (5) Step 4 is then repeated. (6) An additional 1.5 ml of HNO ₃ and 2.50 ml of ultrapure water are added. (7) The bomb is then resealed and heated to 150 °C overnight. (8) The sample is then diluted to 50 ml with ultrapure water	Hu and Gao (2008)	

Table	4	A selection o	t typica	l sample '	treatment	procedures [•]	tor close	d vesse	acid	digestio	'n
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revolutionizing chemical sample preparation and chemical synthesis (Chen et al., 2008; Kingston and Haswell, 1997; Smith and Arsenault, 1996).

Microwave digestion makes use of electromagnetic radiation with a typical frequency of 2450 MHz to generate heat (Matusiewicz and Sturgeon, 1989). Compared with classical heating, microwave heating is many times more efficient. When irradiated by microwave energy, polar molecules and ions are energized via mechanisms of dipole rotation and ion conductance, respectively (Gilman and Engelhart, 1989; Neas and Collins, 1988). The radiating energy is absorbed by both the digestion medium and the sample molecules, which enhances the chemical reaction that completes the decomposition of the sample. Furthermore, localized internal heating of individual sample particles can cause these particles to burst, thus allowing new surfaces to come into contact with reagents increasing the dissolution rate (Nadkarni, 1984). The high dielectric constant of the reagent facilitates the absorption of the radiating energy. Aqueous solutions of acids absorb microwave radiation to a lesser extent than water alone does. It has been shown that the efficiency of absorption of MW radiation decreases according to the following series: nitric, hydrofluoric, sulfuric, and hydrochloric acids (Kingston and Jassie, 1986; Šulcek and Povondra, 1989). Microwaves only heat the liquid phase, while vapors do not absorb microwave energy. Thus, very high temperatures can be reached at relatively low pressures, which is a key advantage of microwave technology (Matusiewicz, 2003). The vessel material also plays a significant role in microwave dissolution techniques. Glassy carbon and platinum vessels cannot be used with this method, but quartz glass and plastics, such as Teflon PFA and polycarbonate, are suitable because they cause no loss of energy to the vessel and allow for quick and efficient heating. Modern microwave digestion systems monitor both the pressure and temperature in the container. The electronic controls of these systems allow for very reproducible digestion conditions, which also reduces the need for operator attention.

The microwave digestion system has been used with both open and closed vessels. Moreover, it is also suitable for online digestion in continuous-flow systems (Chen et al., 2008; Pichler et al., 1999; Smith and Arsenault, 1996). Open vessel microwave-assisted digestion is a suitable and rapid alternative to conventional hot plate digestion for mafic geologic, environmental, and biological materials (Bettinelli et al., 1989; Lamble and Hill, 1998; Nna-Mvondo et al., 2008; Taylor et al., 2002). Microwave heating is also particularly applicable to closed vessel digestion because it offers the advantages of a more vigorous digestion at elevated temperatures and pressures, a reduced likelihood of contamination, the retention of volatile elements, very rapid heating and cooling compared with conventional PTFE-lined stainless-steel bombs, and improved safety measures (Matusiewicz, 1994; Matusiewicz and Sturgeon, 1989; Pichler et al., 1999; Smith and Arsenault, 1996).

The microwave digestion technique has been widely used in the digestion of soils, sediments, coals, airborne particulates, sludge, and organic environmental and biological samples for which easily dissolved elements are required (Bettinelli et al., 1989, 2000; McGrath, 1998; Nandy et al., 2008; Sen Gupta and Bouvier, 1995; Srogi, 2007; Sun et al., 2001a,b; Tuzen et al., 2004; Wang et al., 1996, 2006; Xu et al., 2005). However, it should be noted that the relatively short digestion times (within 1 h) in a microwave system are insufficient to decompose all minerals, which is in contrast to the long digestion times (up to several days) used in conventional PTFE-lined stainless-steel bombs (Navarro et al., 2008; Potts and Robinson, 2003; Yu et al., 2001). Many authors found that microwave digestion did not always allow the measurement of elements, such as Cr, Zr, Hf, and HREEs, especially when refractory mineral phases (e.g., zircon, chromite, rutile, corundum, and cassiterite) were present (Lamothe et al., 1986; Navarro et al., 2008; Sen Gupta and Bertrand, 1995a,b; Totland et al., 1992; Wu et al., 1996; Yoshida et al., 1996; Yu et al., 2001).

15.5.3.4 Partial Dissolutions

In some special geochemical studies, such as mineral exploration and environmental monitoring programs, only the elements of interest need to be analyzed in the rocks, soils, sediments, sludges, oils, and other samples. The elements of interest are not thought to be present in the undissolved fraction (or the part present in the insoluble substrate is not of interest) and can be totally dissolved by an acid or combination of acids without HF. Partial dissolution may also be useful for selectively removing unwanted components of rock samples to eliminate bias (Walsh et al., 1997). For these purposes, complete digestion is often unnecessary. Due to partial acid attack without HF for sample digestion, the resulting solution can be used directly for measurements without evaporating to dryness on a hot plate, and such methods are simple, fast, and economical.

The purpose of partial dissolution is to save time and to simplify the sample digestion procedure. The procedure for different sample types is variable. To leach the total contents of elements of interest in samples, the results of a partial dissolution method should be compared with those of the total dissolution with HF to evaluate the validity of the proposed method (Sastre et al., 2002). To leach the elements of interest in one phase of the samples, the elements in the residue of the samples should not be leached. For example, 5% HNO₃ was used to leach off the trace elements in the carbonate mineral phase from the residue phase (Chen et al., 2005; Feng et al., 2009).

HCl, HNO₃, and HClO₄ are generally used when searching ore deposits in which base metals and sulfides are of interest. Partial dissolution using HNO3 and HClO4 is useful for analyzing heavy metals in soils and sediments for environmental geochemistry studies (Duzgoren-Aydin et al., 2006; Lee et al., 2006; Luo et al., 2011). Partial leach procedures using aqua regia and cyanide are often used for gold exploration because a very large sample mass can be digested to measure very low concentrations of gold. Aqua regia is often used for the rapid dissolution sulfide phases (Church et al., 1987), soils, and sediments (Balaram et al., 1999; Bettinelli et al., 2000). In comparison with HF and fusion techniques, aqua regia dissolution invariably results in the retention of volatile elements (As, Hg, S, Sb, Se, etc.) in the dissolved solution (Potts and Robinson, 2003). The use of aqua regia and HNO₃ leaches in a high-pressure microwave system on sediments, soils, and sludges, and environmental samples for the determination of Cd, Co, Cr, Cu, Hg, Ni, Pb, and Zn were compared by Florian et al. (1998). The element recoveries obtained using the aqua regia procedures were superior to those obtained using the nitric acid procedures. When a nitric acid procedure was applied, the extraction efficiency strongly depended on the applied leaching parameters and varied for certain elements among different materials. For samples with a high content of organic matter, digestion with HNO3 can substitute for total digestion with HF (Sastre et al., 2002).

An important problem encountered in analyses based on the partial dissolution of the material is that there are few recognized reference materials available that can be used for method validation (Walsh et al., 1997).

15.5.3.5 Dry Ashing Techniques

Dry ashing techniques are especially suitable for samples with a high organic matter content, such as biological tissues, plants, or food. The organic matter can be completely eliminated using a muffle furnace equipped with temperature control programs and with atmospheric oxygen serving as the oxidizing agent, and all the associated elements in the samples are typically transformed to carbonate or oxide forms in the residue. The characteristics of the carbonate or oxide forms in ash after dry ashing differ: some will be stable to subsequent acid attack, whereas others will be digested readily with no loss of the analytes. The selection of an appropriate ashing temperature (typical ashing temperatures are 450–550 °C) and ashing vessels is important to ensure quantitative decomposition of the organic matter without loss of interest by volatilization, the formation of refractory oxides that are difficult to dissolve by normal digestion, or reactions with the crucible materials.

Porcelain, vitreous silica, quartz, and platinum crucibles are commonly used as ashing vessels. Porcelain and vitreous silica contain some K, Na, Al, Fe, Ca, Mg, and Ti oxide impurities, which may react with the sample ash to produce a complex silicate and result in contamination and loss of the analytes. A quartz crucible has a relatively higher melt point than a vitreous silica crucible and much lower impurity levels. The best vessel for dry ashing is the platinum crucible, which is unaffected by any of the usual acids, except concentrated, high-temperature phosphoric acid and aqua regia. Although the cost of platinum is higher than that of porcelain and quartz crucibles, its lifetime is unlimited, and it does not induce contamination.

The typical procedure for dry ashing is as follows (Hoenig, 2003): 0.2-2.0 g of dried sample (105 °C) is accurately weighed into a platinum crucible and then placed into a cold muffle furnace and progressively heated to 450 °C over 8 h; the sample is maintained at this temperature for 5 h. After cooling, the resulting ash, if it is of white or light gray color and without silicates in its matrix (e.g., animal tissues, milk, fruit juices, or blood), is dissolved in 1 ml of HNO3 and 20 ml of water and heated to a gentle boil. The solution is subsequently transferred to a calibrated flask or tube and made up to an appreciate volume for measurement. If the sample contains silicates (e.g., plants, sludges, soils, or sediments), 0.5-2.0 ml of water is added at the wall of the crucible to moisten the residue and avoid loss of the fly ash. Then, 1.0-3.0 ml of HNO3 and 0.5-2.0 ml of HF are added, and the sample is evaporated slowly to dryness on a hot plate. This step is repeated two times, and the residue is then dissolved, as described previously. Good recoveries for most major and trace elements (except volatile elements, such as As, Se, Hg, and S) are obtained for environmental, biological, and food samples. The addition of HF is necessary for quantitative release of the analytes into the solution when the sample contains silicates.

The main advantage of dry ashing is the possibility to treat large masses of samples and to dissolve the resulting ash in a small volume of acid to preconcentrate the trace elements in the final solution. Because of the heterogeneity of many biological materials, the ability to process larger masses of sample using dry ashing helps minimize errors in the results. Compared to wet digestion methods, the sample matrix is simplified, and the resulting ash is completely free of organic matter; consequently, the influence of residual carbon or some undigested organic molecules on ICP-MS or ICP-AES measurements is minimized. However, the charring process may have both oxidizing and reducing conditions during the combustion process, and the temperature, in some cases, may be several hundred degrees greater than that of a muffle furnace because of the combustion of the organic materials, which may result in poor recoveries of some volatile elements. It is therefore recommended that the whole procedure must first be

validated using CRMs with compositions similar to those of the samples to be analyzed (Hoenig, 2003).

Numerous studies about the determination of major and trace elements in biological tissues, plants, and food samples using dry ashing techniques have been published in the past several decades. Heanes (1981) measured Co and Mo in plants. Dixon et al. (1980), Harju et al. (1997), and Lee et al. (1986) used the dry ashing technique to determine trace metals in proteins, oily foodstuffs, and trunk wood. Arnaud et al. (1992) used dry ashing to determine Zn in human milk. An interlaboratory study of a method for the determination of Pb, Cd, Zn, Cu, Fe, Cr, and Ni in foodstuffs after dry ashing at 450 °C that involved 16 laboratories was reported by Jorhem (1993). Kucera and Soukal (1995) reported that only 7% of Cd losses were observed after dry ashing at temperatures less than 500 °C and 30% when the temperature was increased to 900 °C. The most appreciable Cd losses (up to 35%) occurred when sulfuric acid was used on wet ashing due to Cd retention in the CaSO₄ precipitate formed. Mader et al. (1997) reported that, in the absence of an ashing aid, a temperature of approximately 500 °C was necessary for complete decomposition of plant tissues, whereas the more resistant animal tissues required significantly higher temperatures. Ming and Bing (1998) used dry ashing, microwave-oven digestion, and traditional open vessel acid digestion methods to determine REEs in human hair, and the results for the three digestion methods were comparable. Quantitative recoveries (>95%) were obtained for the determination of Fe, Cu, Zn, and Mn in pigfeed sample by dry ashing (Van Paemel et al., 2005). A dry ashing temperature in the 500-550 °C range was found to produce ashes with an organic carbon content of less than 1 wt%, which allowed the recovery of a significant proportion of the elements present in the original biomass (Llorente and Garcia, 2006). Fish, baby food, and honey were decomposed by dry ashing at 450 °C for 4-16 h, with subsequent recoveries of approximately 95% for Al, Cd, Co, Cr, Cu, Fe, Mn, Se, and Zn (Saracoglu et al., 2007; Tuzen, 2003; Tuzen et al., 2007). Khuder et al. (2010) used an improved dry ashing method to determine K, Ca, Ti, Cr, Mn, Fe, Ni, Cu, Zn, Rb, and Sr in bee honey samples. Borosova et al. (2010) used the dry ashing technique to measure Ni in the hair. Wang and Li (2010) used microwave-assisted decomposition and dry ashing methods to measure 27 elements in 16 Chinese rice samples and 11 Japanese rice samples. The dry ashing technique was also used to determine boron isotopes in plant tissues (Rosner et al., 2011; Vogl et al., 2011).

Although the dry ashing technique has been successfully used for various samples and different elements, as previously described, it is still a problematic technique for sample decomposition, especially with respect to the determination of trace contents of heavy metals, halogenides, S, As, Sb, Se, Te, and Bi. Saarela et al. (1995) reported that dry ashing is a sensitive and reliable technique for the determination of elements with atomic numbers greater than 20 in biological materials with a low ash content; however, they also reported that volatile elements, such as halogenides and sulfur, can be partly lost depending on the composition of the material studied. Mader et al. (1998) studied the retention of Cd, Cr, Cu, Mn, Pb, and Zn in solid residue when 1.5% HNO₃ was used as a leaching medium after the classical dry ashing of ten materials. Their

results showed that the addition of HF was necessary for quantitative release of the analytes into the solution. Koh et al. (1999) investigated the losses of trace elements in plant samples after dry ashing at 105-600 °C for 0.5-24 h in a muffle furnace. The amounts of Cl and Br decreased over the temperature range of 200-400 °C in all samples, especially in stem samples. Although arsenic is very problematic, Ybanez et al. (1992) successfully used dry ashing and hydride generation AAS to measure As in mussel from seafood products. Vassileva et al. (2001) reported results obtained by dry ashing of plants of terrestrial origin without As and Se losses that did not require a separate preparation methodology for As and Se analysis; however, they observed significant losses of As and Se for aquatic plants. Sahayam et al. (2010) added palladium nitrate $(Pd(NO_3)_2)$ to avoid the loss of volatile As during dry ashing. The recovery of both As(III) and As(V) was found to be nearly quantitative; the results were in good agreement with the certified values and were comparable with results obtained using closed microwave digestion. Oxidizing reagents, such as high-purity magnesium nitrate and magnesium oxide, are commonly used as ashing aids to prevent the volatilization of analytes (e.g., As, Se, Sb, Te, and Bi) and also to speed up the ashing process (Gorsuch, 1970; Hoenig, 2001; Matos-Reves et al., 2010; Mindak and Dolan, 1999). However, the addition of an ashing aid significantly increases the content of inorganic salts, which might be a problem for the subsequent determination of trace elements and might also contribute to contamination (Korn et al., 2008)

Because numerous researchers doubt the validity of dry ashing decomposition, comparisons between dry ashing and wet digestion have been reported since the appearance of commercial wet digestion microwave heating devices for trace element analysis. The wet digestion methods were generally faster than the dry ashing methods but required the use of large amounts of reagents and, therefore, gave higher blank contributions for some elements (Adeloju, 1989). Some researchers have reported that microwave digestion methods give higher recoveries for many elements in comparison with dry ashing methods (Aydin, 2008; Bakircioglu et al., 2011; Demirel et al., 2008; Guldas, 2008; Kaya et al., 2008; Soylak et al., 2004; Yaman and Cokol, 2004; Zachariadis et al., 1995). However, other researchers have found no significant differences between dry ashing and microwave digestion for the determination of major and trace elements in organic matter (Moreno et al., 2008), pine needles (Vaisanen et al., 2008), plant samples (Masson et al., 2010), and human hair (Dogan and Kaya, 2010; Ming and Bing, 1998). Based on these observations, it may be stated that dry ashing techniques also permit good recoveries for most major, minor, and trace elements in various organic materials under suitable conditions. Dry ashing techniques ensure the best simplification of the matrix through the efficient elimination of organic matter, which is a prerequisite for ensuring accuracy with some analytical techniques. This technique will undoubtedly remain an important and unique alternative method for the decomposition of organic materials in the near future.

15.5.3.6 Alkali Fusions

Alkali fusion is generally used to decompose geologic samples that contain refractory minerals that are difficult to digest with acids, and this method ensures the complete quantitative attack and subsequent dissolution of all elements in the solution. NaOH, KOH, Na₂O₂, Na₂CO₃, LiBO₂, and Li₂B₄O₇ are typically used for this type of fusion. Alkali fusion is traditionally used to analyze the major elements in geologic samples. Due to its relatively high blank level and the large amount of total dissolved solids in the final solution, this method is rarely used for trace element analysis in geologic samples recently.

15.5.3.6.1 Fusion with NaOH and KOH

NaOH and KOH are well-known extremely efficient fluxes for the decomposition of silicate minerals. This decomposition occurs at relatively lower temperatures (500-700 °C) and requires less time (10-30 min) than fusion with Na₂CO₃ and LiBO₂ does. Molten alkalies are particularly corrosive, and the fusion procedures using these materials should be performed at the lowest temperature as possible. The amount of NaOH and KOH used is generally at least five times the sample mass for fusion in a muffle furnace or spirit lamp. The sample powder must be well mixed with the fluxes to avoid any incomplete decomposition during the fusion; otherwise, the fluxes will not be in contact with the sample, thus resulting in a partial decomposition and difficulty dissolving the alkali cake after fusion. The muffle furnace is normally heated to approximately 300 °C for 30 min to remove the water in the alkali fluxes and, thus, to prevent 'alkali creep' over the edge of the crucible. The temperature is then increased to 650-700 °C and maintained for approximately 15-30 min. The alkali cake can be rapidly dissolved with hot water and then acidified with HCl or HNO₃. A silver, nickel, iron, or graphite crucible can be used for NaOH and KOH fusion; the use of a silver crucible is quite common.

15.5.3.6.2 Sinter or fusion with Na_2O_2

Sodium peroxide is a strongly oxidizing alkali flux. Most of the minerals in silicate rocks can be rapidly decomposed by sinter with Na₂O₂ at a temperature of 480 ± 20 °C, and a sample-to-Na₂O₂ ratio of 1:4 is usually used (Rafter, 1950; Seelye and Rafter, 1950). Lower temperatures (below the melting point of Na₂O₂) result in less corrosion of the decomposition vessel. The sample powder must be fine and well mixed with Na₂O₂ before heating. Na₂O₂ is particularly useful in mineral analyses, and it is the only flux that can easily be used for the complete decomposition of zircon, tourmaline, cassiterites, or chromites.

Platinum, gold, zirconium, silver, nickel, and iron crucibles can be used for Na_2O_2 sinter experiments. Nickel or iron crucibles may need to be discarded after a few uses due to the strong corrosion of the crucible materials. To avoid excessive attack of the crucible materials, a thick layer of fused anhydrous sodium carbonate is used before adding and mixing the sample with the Na_2O_2 flux, and the duration of the subsequent fusion should be limited. Zirconium crucibles have been shown to have superior resistance to molten Na_2O_2 flux, and the temperature can be increased to 700 °C for the fusion process (Belcher, 1963). The temperature must be precisely controlled at 510–520 °C for sintering in silver, nickel, and iron crucibles to prevent excessive corrosion of the crucible materials. For platinum crucibles, the fusion can be conducted at 540 °C without introducing Pt into the solution. The reaction of Na₂O₂ cake with water is violent and may give rise to local overheating and spitting of the alkali solution. For silicon analyses, the use of glass beakers must be avoided because the beaker materials may be corroded. The use of a PTFE beaker is suggested for dissolution of the alkali cake. After cooling, water is carefully added to the crucible to leach the contents, and then, the alkali solution is dissolved with HCl or HNO₃. REEs, Sc, Y, Th, and Sr can be analyzed using this technique (Robinson et al., 1986). Yu et al. (2001) reported poor recoveries for Rb, Zr, Nb, Mo, Sn, Sb, Cs, Ba, Hf, Ta, Tl, Pb, and Bi in nine reference rocks, but Meisel et al. (2002) reported a reproducibility of 5–10% for REEs, Y, Zr, Hf, Ta, and Nb.

15.5.3.6.3 Fusion with LiBO₂ and $Li_2B_4O_7$

LiBO₂ is a nonoxidative alkali flux. Fusions can be performed at 950-1050 °C with sample-to-flux ratios of 1:3-1:5 in Pt crucibles for approximately 15 min. LiBO₂ is readily available as a highly pure reagent and can decompose all forming minerals and most accessory minerals in rocks (Ingamell, 1970). LiBO₂ is the flux that most commonly replaces NaOH and Na₂O₂ fusion. The major advantage of LiBO₂ fusion is that all of the major elements can be analyzed with one sample preparation. The disadvantages of this method are that loss of K and Na may occur at relatively high temperatures, a long time is required for sample fusion, and the alkali cake is difficult to leach off for dissolution. Ultrasonication has been suggested to accelerate the cake dissolution (Ohlweile et al., 1973).

 $Li_2B_4O_7$ has a higher melting point (930 °C) than that of LiBO₂ (845 °C) and is slower to dissolve. $Li_2B_4O_7$ is better for ultramafic rocks, whereas $LiBO_2$ is preferred for mafic and acidic rocks. A HF/HNO₃ mixture was used to digest the melt to remove some of the Si and B as volatile fluorides to reduce the matrix in the solution (Panteeva et al., 2003; Yu et al., 2001).

15.5.3.6.4 Fusion with Na₂CO₃

Anhydrous sodium carbonate can decompose many silicate materials in platinum crucibles when heated to 1000-1200 °C for approximately 1 h at a sample-to-flux ratio of 1:3 or 1:5 for acidic rocks and up to 1:15 for ultramafic rocks (less than 40% SiO₂) (Sulcek and Povondra, 1989). Volatile elements, such as As, Se, Tl, and Hg, are lost during this type of fusion. Some refractory minerals cannot be decomposed by Na_2CO_3 . The addition of an oxidizer, such as KNO₃, Na_2O_2 , or KClO₃, enhances the decomposition. Na₂CO₃ together with NaOH or Na2O2 decomposes accessory phases, such as zircon, cassiterite, corundum, ilmenite, and monazite, and melts at much lower temperatures than do Na₂CO₃ fluxes (Potts, 1987). A platinum crucible may become iron-stained after a few uses for the Na₂CO₃ fusion of rocks, indicating that some iron has been reduced to its metallic state and became alloyed with platinum. The iron can be removed by heating with 6N HCl or fusing with potassium pyrosulfate in the crucibles.

The alkali fusion technique is traditionally used for the analysis of major elements in silicate rocks (Jeffery, 1975). When NaOH, KOH, and Na_2O_2 have been used for this purpose, Na and K were unable to be detected, and large quantities of a given element may result in polyatomic interferences during ICP-MS analyses (Jarvis, 1992). Different combinations of the reagents for a flux mentioned earlier can be used for different types of samples. For example, the combinations of

Na₂CO₃ and H₃BO₃ (Watanabe, 2001; Watanabe and Otsuki, 2008), Na₂CO₃ and Na₂B₄O₇ (Mitsumata and Aoki, 2004), and KOH and H₃BO₃ (Mutsuga et al., 2011) are used for the decomposition of SiC, silicate, blue phosphor, and titanium dioxide, respectively. One advantage of the alkali fusion method is the rapid decomposition of geologic samples containing refractory minerals compared with acid digestion. Almost all minerals and accessory minerals can be completely dissolved. For example, platinum group elements (PGEs) in mafic and ultramafic rocks (Jarvis et al., 1997; Meisel et al., 2003b; Qi et al., 2003; Stone and Crocket, 1993; Totland et al., 1995), graphite and silicon carbide (Yamaguchi et al., 2000), barium sulfate (using ammonium hydrogen sulfate for the decomposition) (Matsumoto and Koura, 2001), ceramic materials (using a mixture of Na₂CO₃ and ZnO for the decomposition) (Mihaljevic et al., 2001), zircon (Jain et al., 2001; Yamagata et al., 2008), and chromitites (Potts et al., 1992) can all be dissolved.

The main disadvantage of using fusion for the decomposition of rocks and minerals is the relatively high blank level and the high level of total dissolved solids in the final solution, which affects modern instrumental methods of analysis (e.g., ICP-MS and ICP-AES). For example, when the alkali fusion technique is used, the dilution factor for ICP-MS is 5-10 times higher than that of acid digestion (at least 1000), resulting in a deterioration in the detection limits and the need for frequent cleaning of the cones, nebulizer, and spray chamber. When LiBO₂ has been used, it can be difficult to determine Li or B in subsequent acid digestion runs due to high background levels left in the ICP-MS. Thus, the relatively high blank level and high salt concentration in the solution after alkali fusion limit the application of the alkali fusion technique for trace element analysis, which has resulted in a limited number of papers reporting the use of alkali fusion for the detection of major to ultratrace elements in samples using ICP-AES or ICP-MS (Awaji et al., 2006; Fujimori et al., 2001; Sekimoto et al., 2002; Wei and Haraguchi, 1999). Wang et al. (2003) used LiBO₂ fusion to decompose geologic samples. After the fusion and the subsequent dissolution with HNO₃, the elements Ti, Mn, Co, Sr, Y, Nb, Ta, Zr, Hf, In, Ba, Th, and REEs can be quantitatively precipitated using 50% NaOH to adjust the solution to strong alkaline. This method can effectively decompose samples containing high concentrations of Al, Nb, Ta, Zr, and Hf, which are difficult to decompose using acid digestion. The matrix from the flux is removed, and the detection limits are improved with this method, but only 26 elements can be determined.

Volatile elements, such as Sn, Sb, Tl, Pb, and Zn, may be loss when $LiBO_2$ or Na_2CO_3 fusion is used at high temperatures (Totland et al., 1992; Yu et al., 2001). However, alkali fusion is often used to prepare samples of some volatile elements that can easily evaporate by acid digestion, such as fluorine and chlorine (Anazawa et al., 2001; Malde et al., 2001; Tarafder et al., 1997), boron (Terashima et al., 1998), boron isotope (Tonarini et al., 1997), and iodine (Yamada et al., 1996).

In summary, the alkali fusion technique is used primarily to analyze major elements, decompose refractory minerals, and prepare samples of some volatile elements that can easily evaporate by acid digestion. Unlike the PTFE high-pressure bomb technique, hazardous reagents, such as HF or HClO₄, are not used for alkali fusion. However, it is difficult to purify the alkali flux, which results in relatively high blank levels and larger dilution factors for modern instrumental analysis. Although the decomposition time in the fusion method is short compared with that of acid digestion, the fusion method is labor intensive and more difficult to apply with large numbers of geologic samples. An automated fusion system has been reported to simplify the alkali fusion procedure (Govindaraju and Mevelle, 1987).

15.5.3.7 Fire Assays

Fire assay (FA) technique is normally used for preconcentration and separation of the precious metals, including PGEs, Au, and Ag, from a large amount of complex matrices, such as Cu-Ni-PGE ores, mafic and ultramafic rocks, minerals, and soils. PGEs are known to be sensitive indicators for understanding several fundamental aspects of the origin and evolution of the Earth and mantle-derived magmas, such as the core segregation, core-mantle interaction, degree of partial melting, the extent of sulfide segregation and silicate fractionation, and melt-rock reactions, and thus are important for understanding the petrogenesis of mantle-derived rocks (Crocket and Paul, 2004; Ely and Neal, 2003; Lightfoot and Keays, 2005). Consequently, the FA technique is important for geochemical studies. A large amount of sample mass (5-100 g) is typically taken to obtain better accuracy, precision, and lower limits of detection and to avoid the 'nugget effect.' Two types of collector are currently used to collect precious metals from the matrix: lead FA (Pb-FA) and nickel sulfide FA (NiS-FA).

15.5.3.7.1 Lead fire assay

For Pb-FA, PGEs are extracted from sample matrix by reductive fusion and are collected in a lead button (Bugbee, 1957; Farago et al., 1996; Haffty et al., 1977; Hall and Pelchat, 1994; Van Loon and Barefoot, 1991). The fluxes comprise sodium carbonate, borax, litharge, and a reducing agent, such as flour or charcoal, and are fused at 1000 °C in a clay crucible. The exact composition of the flux largely depends on the nature of the sample. Silver is normally added to the flux as powder or as a solution to increase the recoveries of PGEs. The ratio of Ag to Au+PGE is approximately 20:1. The precious metals are partitioned into a Pb button that settles to the bottom of the crucible. The matrix elements react with the flux component to form a slag that is subsequently discarded. After cooling, this button is separated and cupelled (i.e., oxidized) in a furnace to give a noble-metal prill, which can be dissolved for the quantitative analysis of Au, Pt, and Pd. Rhodium, Ir, Ru, and Os are partially lost (Hall and Pelchat, 1994; Van Loon and Barefoot, 1991). The molten lead and slag are poured into iron molds and are allowed to solidify. The slag is removed, and the lead button is subsequently cupelled by being placed in a cupel made of bone ash or MgO and heated to approximately 960 °C in an oxidizing atmosphere. Most of the molten lead is absorbed (as PbO) by the cupel or is volatilized, which leaves a Ag prill that contains Au, Pt, Pd, and Rh. Osmium is completely lost during the cupellation procedure as a volatile oxide, and Ru and Ir are partially lost as oxides into the cupel. The silver prill can be measured by instrumental neutron activation analysis or dissolved in acid and measured by AAS, ICP-AES, or ICP-MS.

Pb-FA is normally used for the determination of Au, Pt, and Pd in geologic samples. Detection limits of 2, 0.1, and 0.5 ppb for Au, Pt, and Pd, respectively, have been reported. Further improvements in the detection limits through purification of the flux components have been suggested (Hall and Pelchat, 1994). Compared with the NiS-FA technique, the Pb-FA technique has lower recoveries and higher detection limits (Juvonen et al., 1994), and only Au, Pt, and Pd can be determined using this method.

15.5.3.7.2 Nickel fire assay

The NiS-FA method is capable of collecting all of the PGEs in different types of samples and has been studied extensively over the past 20 years (Balaram et al., 2006; Bedard and Barnes, 2002; Bedard and Barnes, 2004; Farago et al., 1996; Frimpong et al., 1995; Gros et al., 2002; Haffty et al., 1977; Hall and Pelchat, 1994; Jackson et al., 1990; Jarvis et al., 1995; Jorge et al., 1998; Juvonen et al., 1994; 2002; Li and Ebihara, 2003; Li and Tong, 1995; Morcelli et al., 2004; Plessen and Erzinger, 1998; Ravizza and Pyle, 1997; Sun and Sun, 2005; Van Loon and Barefoot, 1991; Wichmann and Bahadir, 2001; Zereini et al., 1994; Zhou et al., 2001). In a typical NiS-FA, approximately 15 g of finely ground rock sample is mixed with 5 g of Ni, 3 g of S, 20 g of sodium tetraborate, 10 g of sodium carbonate, and 5 g of quartz in a clay crucible and is heated to 1100-1200 °C for approximately 75 min. During fusion, two phases are formed in the melt: the siliceous phase and the sulfide phase in which PGEs and Au are concentrated. Covering the crucible can prevent the possible contamination during fusion, and the size of the NiS button is increased because less of the sulfide is oxidized during fusion. The NiS button can be measured directly by laser ablation (LA)-ICP-MS (Jarvis et al., 1995). Alternatively, it can be dissolved in HCl to separate Ni from PGEs and then measured using the solution ICP-MS. The button is dissolved in HCl in an Erlenmeyer flask at approximately 90 °C, and this process is usually performed overnight. The insoluble PGE sulfides, along with the dissolved PGEs, are then collected by Te coprecipitation, which can effectively collect all of the PGEs in the form of a black precipitate (Jackson et al., 1990; Oguri et al., 1999; Sun et al., 1993); a good recovery of Os has even been reported (Gros et al., 2002).

Different sample type may require different fusion compositions for NiS-FA. Incomplete fusion has been observed in the case of chromite-bearing samples using normal FA techniques, with recoveries of PGEs ranged from 76% to 85% (Asif and Parry, 1991). Juvonen et al. (2002) changed the composition of the flux by adding oxidant to the fusion flux for black shale and a reducing reagent in the case of magnetite samples to improve the outcome of the fusion. Carbon can be oxidized either by roasting the sample prior to fusion (700 °C for 2 h) or by adding potassium nitrate to the fusion. For magnetite, a reducing agent (flour or potassium tartrate) is required. Base metals, As, Cu, Fe, Sb, and Zn, can also be collected with Ni into the sulfide button. If present in large quantities, these metals will lead to the formation of an unusually large sulfide button that is difficult to dissolve in HCl and that will remain with the PGEs in the final solution and cause interference in the ICP-MS determination.

Frimpong et al. (1995) found that the recovery of Ir, Os, and Au from a komatiite sample did not significantly depend

on the mass of the collector used; recoveries of Ru, Rh, Pd, and Pt, however, were reproducible and dependent upon the collector mass. The partition coefficients were found to depend upon the PGE concentrations. They concluded that, when PGE concentrations are determined at low levels, the NiS-FA collection method may cause significant biases.

Because of difficulties in establishing conditions for the complete fusion of all chromite grains, Zereini et al. (1994) reported that the best results were obtained when a 1:1 mixture of the lithium tetraborate and sodium tetraborate in flux was used. Bedard and Barnes (2004) reported a procedure based on the addition of sodium metaphosphate to the fusion mixture for the determination of PGEs in chromites. The optimum composition of the fusion mixture was found to be 10 g of sodium metaphosphate, 9 g of silica, 15 g of sodium carbonate, 30 g of lithium tetraborate, 7.5 g of nickel, and 4.5 g of sulfur for 10 g of sample to achieve complete dissolution of the chromite grains. The addition of sodium metaphosphate improved chromite dissolution in the flux and appeared to improve recoveries of the PGEs.

To accelerate the dissolution process, NiS buttons are crushed or pulverized before treatment with acid. However, this method introduces both the possibility of cross-contamination between samples and the possibility of losses of NiS together with PGEs during this step. Boisvert et al. (1991) reported that crushing buttons in a shatter-box resulted in losses of up to 20% of the NiS buttons. Some researchers crush the NiS buttons in plastic bags with a hammer to avoid mechanical losses (McDonald et al., 1994).

One disadvantage for samples that contain low levels of PGEs and Au is the contamination caused by reagent blanks as a result of the large amount of chemicals used as fluxes. Nickel produced by the carbonyl process must be used. Reagent blanks for PGEs and Au were reduced significantly when a much smaller quantity (0.5 g) of nickel was used. Because the high reagent blank of NiS-FA was found to originate primarily from the nickel, Sun et al. (1998) developed a method to purify the nickel oxide. They dissolved 500 g of nickel in HCl. The solution was then purified to be free of PGEs by Te coprecipitation or through the use of anion-exchange resin. The purified solution was neutralized with Na₂CO₃, and the precipitate by centrifugation of the solution was subsequently rinsed with water. The precipitate was finally ignited at 700 °C for approximately 2 h. The procedure blank was significantly reduced.

Because the recoveries of PGEs for a single FA are approximately 90%, approximately 10% of the PGEs still remain in the slag. Oguri et al. (1999) developed a duplicate NiS-FA technique to enhance the recovery of PGEs to better than 96.7%. Isotope dilution (ID) was also used for NiS-FA to compensate the loss and improve the accuracies of PGEs, especially Os (Ravizza and Pyle, 1997).

When dissolving the NiS bead in an open beaker, the Ru and Pd can possibly escape as chlorides at low temperatures. Sun et al. (1993) used a closed beaker to prevent the loss of HCl during dissolution of the NiS button. Gros et al. (2002) designed a semiopen system that combines the advantages of the open system and those of the closed-beaker system. They also found that the S/Ni ratio in the NiS button is very important. An increase in the S/Ni ratio results in the formation of nickel disulfide (NiS₂). This compound is poorly soluble in hot hydrochloric acid but is readily soluble in aqua regia. A decrease in the S/Ni ratio to 0.5 results in the formation of $Ni_{1-x}S$ in the NiS button, which is theoretically soluble in hot hydrochloric acid. This change in the S/Ni ratio significantly reduces the formation of insoluble nickel sulfide but does not alter the efficiency of PGE collection in the bead.

Sun and Sun (2005) reported an improved NiS-FA method that involves the addition of Fe to the flux. The Fe assay buttons are disintegrated into powder in water and then dissolved in HCl. This improved technique for the preconcentration of all six PGEs not only overcomes the problems of cross-contamination between samples and button losses at the mechanical crushing step but also speeds up the dissolution of the button. The button can be manipulated more easily than can a conventional sulfide button.

As previously reported, Re is poorly recovered by the NiS-FA technique (Frimpong et al., 1995). Sun et al. (2009) reported an improved Fe–Ni sulfide FA method in which $Na_2B_4O_7$ was used instead of Na_2CO_3 in the flux for the determination of Re, PGE, and Os isotopic ratios. Recovery of Re, to as high as 75%, was achieved, whereas it was less than 10% when Na_2CO_3 was used. When Na_2CO_3 is absent, the alkalinity of the slag phase is decreased, which favors the formation of Re^{4+} . The radius of Re^{4+} (0.77 Å) is virtually identical to that of Fe^{2+} . Consequently, Re^{4+} tends to occur in the Fe–Ni sulfide phase that contains FeS and thereby increasing Re recovery.

In summary, all of the PGEs can be collected using the NiS-FA technique. Large sample masses can be used to reduce the nugget effect. The NiS button can be analyzed directly by LA inductively coupled plasma mass spectrometry (LA-ICP-MS). The main disadvantage of NiS-FA is the relatively large mass of reagents used, which increase the reagent blanks of the PGEs, especially from the used Ni powder. Carbonyl nickel or purified nickel is preferred for fusion. Osmium may be lost as volatile OsO_4 during the fusion and the recovery of Au is usually low. Ru and Pd may be lost as chlorides or Cl-bearing complexes during the dissolution of the NiS bead with HCl. Finally, it must be borne in mind that fire assaying is a technique that requires skill on the part of the operator; different types of samples require the composition of the flux to be varied, and considerable operator experience is required to obtain good results.

15.5.3.8 Carius Tube and High-Pressure Asher (HPA-S)

The use of a sealed glass tube for the digestion of organic materials with concentrated nitric acid under elevated temperatures of 250–300 °C was first described in 1860 and is often referred to as the Carius tube technique. The glass tube is typically a thick-walled quartz ampule in which the sample and the acid have been mixed, and the tube has been subsequently sealed. The tube is transferred to a jacket and heated for several hours for digestion. The modern redesign and application of the Carius tube for the digestion of some refractory materials was accomplished at the US National Bureau of Standards during the 1940s (Gordon, 1943; Gordon et al., 1944; Wichers et al., 1944).

Since the 1990s, the Carius tube technique has been mainly used to determine PGEs and Re–Os isotopes in geologic samples. Shirey and Walker (1995) used the Carius tube technique to determine Re–Os in geologic samples. In a typical analysis,

sample powders are introduced via a glass funnel into the tube. Spikes and aqua regia are added through a funnel and allowed to freeze in the bottom of the tube. As a precaution against bursting, at least half of the tube volume should be left empty. The lower parts of the tubes are immersed in an ethanol-dry ice slush to reduce the vapor pressure of the reagents and to make the sealing of the tube easier and more reliable. This process precludes the possibility of Re or Os loss during loading before sealing, because the reagents can oxidize both Re and Os, which may form moderately volatile oxides. While the bottom parts of the tubes are still frozen, the tube is sealed and then placed in a stainless-steel jacket. The internal volume of the stainless-steel jacket is about twice the external volume of the Carius tube, so that if it bursts, the jacket will not become overpressurized. The tubes are then heated in an oven at 220-260 °C for approximately 12 h. Agitation between multiple heating cycles can aid in sample digestion. It should be noted that, once heated, the inside of the tube may be under high pressure even when the temperature is reduced to room conditions. Thus, the tubes should be treated gently. Protective face and body shields should be worn during subsequent handling.

To avoid the problem of loss of mechanical stability at high temperatures, vessels made of quartz are used in the highpressure ashing (HPA) technique, the digestion principles of which are similar to those of the Carius tube. This pressure digestion system has not only reduced the effective digestion time but has also enabled digestion of extremely resistant materials, such as carbon, carbon fibers, mineral oils, organic waste materials, and PGEs (Knapp, 1984, 1985). Nitric acid alone or combined with HCl is a sufficiently powerful reagent for a variety of samples. The maximum digestion temperature for high-pressure digestion is as high as 320 °C at a pressure of approximately 13 MPa. A commercial pressure digestion system, the HPA-S High Pressure Asher (Anton Paar GmbH, Austria), is available. The HPA-S technique is substantially safer than the use of Carius tubes, but its costs are higher.

The Carius tube technique has proven to be useful for Re-Os chronological dating and Os isotopes (Birck et al., 1997; Brauns, 2001; Cohen and Waters, 1996; Malinovsky et al., 2002; Meisel et al., 2001b, 2003b; Pearson and Woodland, 2000; Qi et al., 2010; Qu et al., 2008; Shirey and Walker, 1995; Sun et al., 2001a,b; Yang et al., 2005) and for the dissolution of most or all Re and PGEs from refractory platinum group minerals in geologic samples (Jensen et al., 2003; Meisel and Moser, 2004a, 2004b; Meisel et al., 2001a, 2003a; Norman et al., 2002; Pearson and Woodland, 2000; Pretorius et al., 2003; Qi and Zhou, 2008; Qi et al., 2007; Rehkamper et al., 1998). Because samples are dissolved in aqua regia under high temperatures (220-240 °C for Carius tubes and 300 °C for HPA-S), all the PGEs are at their highest oxidation state with complete equilibration between spikes and samples. A total of 1-2 g of sample is commonly used in a Carius tube and HPA-S, with agua regia added to less than one-third the total volume of the Carius tube to avoid a possible explosion. The total procedural blanks of HPA-S technique are as low as 0.04 ng for Pt and 0.005 ng for other PGEs (Meisel et al., 2003a) because small volumes of reagents are used and because all of the reagents can be purified. Osmium can be separated by normal distillation (Fritsche and Meisel, 2004; Malinovsky et al., 2002; Qi and Zhou, 2008;

Qi et al., 2007; Shirey and Walker, 1995; Sun et al., 2001a,b), in situ distillation (Brauns, 2001; Qi et al., 2010), or solvent extraction (Birck et al., 1997; Cohen and Waters, 1996). Recently, the use of HPA-S to fully dissolve rock samples with HF and ultimately to quantify REEs and HFSEs has also been reported (Cotta and Enzweiler, 2012). The HPA-S technique enables faster dissolution of samples that contain refractory phases (several hours), which otherwise require several days to digest in PTFE bombs.

Rehkamper et al. (1998) reported a modified Carius tube technique. The Carius tube was incorporated into a liner of high-purity quartz glass that retained the sample and acids during the digestion procedure. This technique achieves total procedural blanks at the 1–15 pg g^{-1} level for PGEs and lowers the blank values by a factor of 10-100 compared to the standard NiS-FA technique. The quartz vessel of HPA-S is stabilized during the digestion process by subjecting it to an external pressure roughly equivalent to or higher than the internal pressure by using N₂ to pressurize the exterior of the vessel to 13 MPa; the vessel is thus protected from explosion (Meisel et al., 2001a). Becker et al. (2006) used a steel pressure vessel that contained 20 g dry ice to digest 2-3 g of peridotite in a quartz glass Carius tubes at 345 °C. The CO₂ pressure that builds up inside the pressure vessel upon warming supports the internal pressure in the Carius tube. They demonstrated that the spinels in peridotites were either completely dissolved or oxidized. Qi et al. (2007, 2008) reported a modified Carius tube method that utilizes a sealed stainless-steel high-pressure autoclave filled with water to prevent possible explosion of the tube. This technique allows a higher temperature (up to 300 °C), a greater volume of aqua regia (up to two-thirds of the total volume of the Carius tube), and thus a larger sample mass (12 g) compared to the normal Carius tube technique. This improved technique is efficient for the digestion of mafic rocks, ultramafic rocks, and chromites. The stainless-steel highpressure autoclave is inexpensive and easy to maintain compared with the equipment required for the HPA-S technique.

In summary, the Carius tube and HPA-S techniques have the lowest procedural blanks of all the digestion techniques. This technique permits Re–Os dating and the determination of low concentrations of PGEs in geologic and environmental samples. However, the reproducibility is poor due to the small sample size (1–2 g) and the possible nugget effect in some sample types. Because only aqua regia is used, this technique may cause incomplete dissolution for some samples that contain refractory minerals, such as chromites and olivines. Carius tubes may explode under high temperatures and pressures, and the HPA-S technique requires expensive equipment.

15.5.3.9 Direct Fusion of Rock Powder

During the last decades, LA-ICP-MS has developed into an accurate and sensitive microanalytical technique for the analysis of solid samples. Because of its unique advantages (minimal sample preparation and rapid throughput) and the problems associated with wet chemical dissolutions, there is an increasing trend to use LA-ICP-MS for whole rock sample analysis. A homogeneous sample is a prerequisite to obtain reliable bulk analytical results using microanalytical techniques. The direct fusion of rock powder is a rather simple method to obtain a homogeneous sample for LA-ICP-MS analysis. Sylvester (2001) has reviewed the trace element analysis of fused whole rock glasses using LA-ICP-MS.

15.5.3.9.1 Flux-free fusion glasses

Rucklidge et al. (1970), Nicholls (1974), and Brown (1977) first reported the W-, Ir-, and Mo-strip fusion methods, respectively, for the determination of major element concentrations on glass beads using an electron microprobe. Fedorowich et al. (1993) prepared glass beads by melting small quantities $(\sim 25 \text{ mg})$ of rock powders on a tungsten-strip heater under an Ar atmosphere. The optimum melt conditions were achieved for most rock types by maintaining the temperature at 1700-1800 °C for 2 min. Hf and Ta contaminations from the tungsten strip and losses of Li and Pb at high strip temperatures and under longer melting conditions were reported. The Hf and Ta contaminations were avoided by using an iridium-strip heater instead of a tungsten-strip heater (Norman et al., 1996). Reid et al. (1999) used isotope dilution for the measurement of Zr and Hf in fused whole rock glasses. To maximize the homogeneity of the sample and enriched tracers, the resulting glass was ground and re-fused on an iridium-strip heater at 1600 °C for 2 min. Replicate analyses on these glasses resulted in precisions of approximately 1% for Zr and approximately 3.5% for Hf, which suggests a high degree of homogeneity in the samples. Jochum et al. (2000) have prepared large amounts (50-100 g) of silicate glasses (MPI-DING glasses) from rock powders to be used as reference materials in platinum crucibles by direct fusion at temperatures in the range of 1400-1600 °C. The resulting glasses are homogeneous for most elements at the micrometer to millimeter scale. Kurosawa et al. (2006) prepared fused glasses of seven Japanese silicate reference materials for LA-ICP-MS and particle-induced x-ray emission analysis by rapid fusion (2 min at 1200 °C) and subsequent quenching in welded platinum capsules. The fused glasses were found to be homogeneous for major elements and for trace elements with concentrations of more than $1 \mu g g^{-1}$ within the observed precision (10%) on a 70 µm sampling scale. The values obtained by LA-ICP-MS for almost all the elements (including the relatively volatile elements Zn, Ga, Rb, and Pb) agreed well with the reference values, except for B and Cu. Pack et al. (2007) prepared small meteorite material glasses ($\sim 10 \text{ mg}$) by means of a SYNRAD 50 W CO₂ laser fusion system and an aerodynamic levitation device. Switching off the laser and cooling via the gas stream ensured rapid quenching. Although the cooling rates were approximately 520 K s⁻¹, crystallization during the quenching of the Fe- and Ni-rich samples could not be avoided. Stoll et al. (2008) developed an automated Ir-strip heater, which enables the accurate control of the melting temperature and the melting time. This control makes it possible to achieve reproducible melting conditions. Elements with condensation temperatures (at a pressure of 10 Pa) higher than approximately 900 K (e.g., Zr, Hf, Ba, Sr, REEs, U, Mo, Ni, Rb, and Ga) were found not depleted at 1600 °C for 10 s.

The direct Ir-strip fusion method is suitable for mafic compositions ($<55 \text{ wt\% SiO}_2$) that can easily be fused at 1200 °C. The fusion of siliceous samples is more problematic, because rapid homogenization during melting on the strip is hindered by high melt viscosities (Nehring et al., 2008). High melting temperatures (1700-1800 °C), long fusion times (60-120 s), and stirring at high temperatures are required to achieve homogeneity in such samples (Nicholls, 1974), but these conditions lead to loss of volatile elements. For example, a high loss (20-90%) of highly volatile elements (e.g., Cs, Ge, Sn, and Pb) was observed at high melting temperatures ($\geq 1600 \,^{\circ}$ C) and long melting times (80 s) (Stoll et al., 2008). The dilution of a high SiO₂ concentration by the addition of MgO is another way to facilitate rapid fusion and homogeneity (Gumann et al., 2003; Nehring et al., 2008). Nehring et al. (2008) evaluated the melting behavior of a wide range of rock compositions (55-80 wt% SiO₂) and the dependence of trace element concentrations on the sample/MgO mixing ratio, melting time, and melting temperature. The melting conditions of 1600 °C and 20 s for samples are recommended when the SiO₂ content is between 55% and 70 wt%, whereas 1800 °C and 20-30 s are often required for samples with >70 wt% SiO₂ (Nehring et al., 2008). Pb and Cs due to volatilization on the Ir strip are partly lost. Zhu et al. (2011) proposed a modified method using a double iridiumstrip heater, which suppresses the volatilization of Pb and Zn to a certain extent, to fuse the rock powder without flux.

In summary, the flux-free fusion of a rock powder is a simple and rapid method that does not require acid digestion. It needs very small amounts of sample (10–50 mg), which makes it suitable for bulk analyses of valuable samples (e.g., lunar samples). The obtained precision and accuracy by LA-ICP-MS on fused whole rock glasses are comparable with those of other routine bulk analytical techniques. The limitations of this technique include the loss of some highly volatile trace elements (e.g., Pb, Zn, Sn, Ge, In, and Cs) during sample fusion and the requirement for high melting temperatures (>1600 °C), long fusion times, and/or the addition of MgO for samples with higher SiO₂ contents (>55%). Selection of an appropriate melting temperature and time for the sample preparation is an essential prerequisite for the accurate analysis of fused whole rock glasses by LA-ICP-MS.

15.5.3.9.2 Lithium-borate fusion glasses

Another method for bulk rock analysis by LA-ICP-MS is to generate glasses by adding the flux agents LiBO₂ (Perkins et al., 1993; Pickhardt et al., 2000; Sylvester, 2001), Li₂B₄O₇ (Günther et al., 2001; Ødegård and Hamester, 1997; Ødegård et al., 1998), or LiBO₂ and Li₂B₄O₇ (Beck and Dietze, 1999; Yu et al., 2003) to the rock powders prior to fusion. This technique is commonly used for making fused glass disks for major element analysis using x-ray fluorescence (Norrish and Hutton, 1969). A commonly used recipe for making fused disks involves mixing 0.1-1 g of rock powder with a fivefold excess of lithium-borate flux in a platinum/5% gold crucible. This mixture is then heated to 1000-1100 °C for 5-20 min in a furnace (Sylvester, 2001). The addition of the flux agents significantly reduces the temperature required for fusion compared with that needed for the rock powder alone, thereby reducing the possible lost of volatile elements. The flux agents can also sufficiently change the melt composition such that olivine crystals do not form upon melt quenching of komatiltes and peridotites and break down the rigid structure of the siliceous melt such that the viscosity is reduced and homogenization is rapidly achieved (Sylvester, 2001). This method is suitable for a wide variety of bulk compositions, including

mafic, intermediate, and silicic rocks (Yu et al., 2003). The analytical precision reported in the studies mentioned earlier typically ranged from 2% to 7% for most trace elements, which suggests that most of the lithium-borate fusion glasses are homogeneous at the scale of laser sampling. The disadvantages of this method include contamination of the rock sample by impurities in the flux agents, a lower detection power because of sample dilution due to flux, and long-term instrumental memory effects from lithium and boron contaminations. For example, for some elements (e.g., Be, Sc, Sr, La, Ti, V, Cr, Ni, Cu, Zn, Y, Zr, Ba, Nd, and Pb), impurities of up to several micrograms per gram have been observed in the lithiumborate flux (Ødegård and Hamester, 1997; Sylvester, 2001). For LA-ICP-MS analyses, these impurities necessitate significant blank subtractions for these elements. Moreover, longterm analyses of the lithium-borate fusion glasses preclude the precise measurement of Li and B isotope ratios and the detection of low concentrations of Li and B on the contaminated instrument. Thus, the lithium-borate fusion techniques are not popularly used for LA-ICP-MS analysis currently.

15.5.4 Summary and Overview

Sample digestion is a fundamental and critical stage in the process of sample analysis, particularly with respect to the digestion of most geologic samples (e.g., rocks, minerals, soils, and sediments) that have to be digested into the solution to satisfy the requirements for chemical analyses. Clearly, no single digestion technique will allow the determination of all elements. The performance of analytical instrumentation for chemical analysis has improved dramatically over the past several decades. However, methods of sample digestion have not changed to keep pace and have lagged far behind developments in analytical instrumentation. Future developments in digestion methods need to be focused on aspects of sample throughput, precision and accuracy, analytical cost, simplification of the operating procedure, and as many elements as possible for measurements, minimization of waste generation, and safety considerations. There is little doubt that automation technology will be increasingly used in sample digestion in the near future, which would especially improve sample throughput and analytical reproducibility.

Improved measurement techniques and instruments allow the use of smaller analytical test portions to determine analyte concentrations. However, smaller test portions mean greater difficulty in achieving representativeness of the population. The minimum sampling mass of the sample materials largely depends on the sample particle size. The processing and use of ultrafine samples ($<30 \,\mu$ m) will significantly influence the future development of digestion techniques as a consequence of the potential significant reduction in the masses of the test portions (Wang et al., 2004), which will provide the added advantages of enhanced digestion efficiency and speed, lower costs, and lower environmental impacts.

Finally, it is evident that sample digestion methods will remain an important area for development, in order to keep up with the rapidly evolving measurement techniques and instruments (see Chapters 15.10; 15.11; 15.12; 15.13; 15.14; 15.15; 15.16; 15.17; 15.18; 15.19; 15.20; 15.21; 15.22; and 15.23).

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References

- Abu-Samra A, Morris JS, and Koirtyohann SR (1975) Wet ashing of some biological samples in a microwave oven. *Analytical Chemistry* 47: 1475–1477.
- Adeloju SB (1989) Comparison of some wet digestion and dry ashing methods for voltammetric trace-element analysis. *Analyst* 114: 455–461.
- Anazawa K, Tomiyasu T, and Sakamoto H (2001) Simultaneous determination of fluorine and chlorine in rocks by ion chromatography in combination with alkali fusion and cation-exchange pretreatment. *Analytical Sciences* 17: 217–219.
- Arnaud J, Bouillet MC, Alary J, and Favier A (1992) Zinc determination in human-milk by flameless atomic-absorption spectrometry after dry ashing. *Food Chemistry* 44: 213–219.
- Asif M and Parry SJ (1991) Study of the digestion of chromite during nickel sulphide fire assay for the platinum group elements and gold. *Analyst* 116: 1071–1073.
- Awaji S, Nakamura K, Nozaki T, and Kato Y (2006) A simple method for precise determination of 23 trace elements in granitic rocks by ICP-MS after lithium tetraborate fusion. *Resource Geology* 56: 471–478.
- Aydin I (2008) Comparison of dry, wet and microwave digestion procedures for the determination of chemical elements in wool samples in Turkey using ICP-AES technique. *Microchemical Journal* 90: 82–87.
- Bakircioglu D, Kurtulus YB, and Ucar G (2011) Determination of some traces metal levels in cheese samples packaged in plastic and tin containers by ICP-AES after dry, wet and microwave digestion. *Food and Chemical Toxicology* 49: 202–207.
- Balaram V, Mathur R, Banakar VK, et al. (2006) Determination of the platinum-group elements (PGE) and gold (Au) in manganese nodule reference samples by nickel sulfide fire-assay and Te coprecipitation with ICP-MS. *Indian Journal of Marine Sciences* 35: 7–16.
- Balaram V, Raju PVS, Ramesh SL, et al. (1999) Rapid partial dissolution method in combination with AAS techniques for use in geochemical exploration. *Atomic Spectroscopy* 20: 155–160.
- Beck JS and Dietze HJ (1999) Determination of trace elements in geological samples by laser ablation-inductively coupled plasma-mass spectrometry. *Fresenius' Journal of Analytical Chemistry* 365: 429–434.
- Becker H, Horan MF, Walker RJ, Gao S, Lorand JP, and Rudnick RL (2006) Highly siderophile element composition of the Earth's primitive upper mantle: Constraints from new data on peridotite massifs and xenoliths. *Geochimica et Cosmochimica Acta* 70: 4528–4550.
- Bedard LP and Barnes SJ (2002) A comparison of the capacity of FA-ICP-MS and FA-INAA to determine platinum-group elements and gold in geological samples. *Journal of Radioanalytical and Nuclear Chemistry* 254: 319–329.
- Bedard LP and Barnes SJ (2004) Improved platinum-group element extraction by NiS fire assay from chromitite ore samples using a flux containing sodium metaphosphate. *Geostandards and Geoanalytical Research* 28: 311–316.
- Begum Z, Balaram V, Ahmad SM, Satyanarayanan M, and Rao TG (2007) Determination of trace and rare earth elements in marine sediment reference materials by ICP-MS: Comparison of open and closed acid digestion methods. *Atomic Spectroscopy* 28: 41–50.
- Belcher CB (1963) Sodium peroxide as a flux in refractory and mineral analysis. *Talanta* 10: 75–81.
- Bernas B (1968) A new method for decomposition and comprehensive analysis of silicates by atomic absorption spectrometry. *Analytical Chemistry* 40: 1682–1686.
- Bettinelli M, Baffi C, Beone GM, and Spezia S (2000) Soil and sediment analysis by spectroscopic techniques part I: Determination of Cd, Co, Cr, Cu, Mn, Ni, Pb, and Zn. Atomic Spectroscopy 21: 50–59.

Bettinelli M, Baroni U, and Pastorelli N (1989) Microwave-oven sample dissolution for the analysis of environmental and biological-materials. *Analytica Chimica Acta* 225: 159–174.

Birck JL, Barman MR, and Capmas F (1997) Re–Os isotopic measurements at the femtomole level in natural samples. *Geostandards Newsletter* 21: 19–27.

Boer RH, Beukes GJ, Meyer FM, and Smith CB (1993) Fluoride precipitates in silicate wet-chemistry: Implications on REE fractionation. *Chemical Geology* 104: 93–98.

Boisvert R, Bergeron M, and Turcotte J (1991) Re-examination of the determination of palladium, platinum and rhodium in rocks by nickel sulphide fire assay followed by graphite furnace atomic absorption measurements. *Analytica Chimica Acta* 246: 365–373.

Borosova D, Manova A, Mocak J, and Beinrohr E (2010) Determination of nickel in hair samples by graphite furnace atomic absorption spectrometry and flow-through stripping chronopotentiometry. *Analytical Methods* 2: 1913–1917.

Brauns CM (2001) A rapid, low-blank technique for the extraction of osmium from geological samples. *Chemical Geology* 176: 379–384.

Brown RW (1977) A sample fusion technique for whole rock analysis with electron microprobe. *Geochimica et Cosmochimica Acta* 41: 435–438.

Bugbee E (1957) A Text Book of Fire Assaying, 3rd edn. New York: Wiley. Chao TT and Sanzolone RF (1992) Decomposition techniques. Journal of Geochemical

Exploration 44: 65–106.

Chen DF, Huang YY, Yuan XL, and Cathles LM (2005) Seep carbonates and preserved methane oxidizing archaea and sulfate reducing bacteria fossils suggest recent gas venting on the seafloor in the northeastern South China Sea. *Marine and Petroleum Geology* 22: 613–621.

Chen LG, Song DQ, Tian Y, Ding L, Yu AM, and Zhang HQ (2008) Application of on-line microwave sample-preparation techniques. *TrAC Trends in Analytical Chemistry* 27: 151–159.

Church SE, Mosier EL, and Motooka JM (1987) Mineralogical basis for the interpretation of multielement (ICP-AES), oxalic-acid, and aqua regia partial digestions of stream sediments for reconnaissance exploration geochemistry. *Journal of Geochemical Exploration* 29: 207–233.

Cohen AS and Waters FG (1996) Separation of osmium from geological materials by solvent extraction for analysis by thermal ionisation mass spectrometry. *Analytica Chimica Acta* 332: 269–275.

Cotta AJB and Enzweiler J (2012) Classical and new procedures of whole rock dissolution for trace element determination by ICP-MS. *Geostandards and Geoanalytical Research* 36: 27–50.

Crocket JH and Paul DK (2004) Platinum-group elements in Deccan matic rocks: A comparison of suites differentiated by Ir content. *Chemical Geology* 208: 273–291.

Croudace IW (1980) A possible error source in silicate wet-chemistry caused by insoluble fluorides. *Chemical Geology* 31: 153–155.

Demirel S, Tuzen M, Saracoglu S, and Soylak M (2008) Evaluation of various digestion procedures for trace element contents of some food materials. *Journal of Hazardous Materials* 152: 1020–1026.

Diegor W, Longerich H, Abrajano T, and Horn I (2001) Applicability of a high pressure digestion technique to the analysis of sediment and soil samples by inductively coupled plasma-mass spectrometry. *Analytica Chimica Acta* 431: 195–207.

Dixon NE, Blakeley RL, and Zerner B (1980) Jack Bean Urease (Ec 3.5.1.5). I. A simple dry ashing procedure for the micro-determination of trace-metals in proteins – The nickel content of urease. *Canadian Journal of Biochemistry* 58: 469–473.

Dogan S and Kaya FND (2010) Determination of zinc and lead in human hair by atomic absorption spectrometry after digestion with tetramethylammonium hydroxide and conventional methods. *Trace Elements and Electrolytes* 27: 110–114.

Dulski P (2001) Reference materials for geochemical studies: New analytical data by ICP-MS and critical discussion of reference values. *Geostandards Newsletter* 25: 87–125.

Dupta JGS (1994) Characterization of major, minor and trace elements in six CCRMP (TDB-1, WGB-1, UMT-1, WPR-1, WMG-1 and WMS-1) and one IWG (ZW-C)

geochemical candidate reference materials. *Geostandards Newsletter* 18: 111–122. Duzgoren-Aydin NS, Wong CSC, Aydin A, Song Z, You M, and Li XD (2006) Heavy metal contamination and distribution in the urban environment of Guangzhou, SE China. *Environmental Geochemistry and Health* 28: 375–391.

Eggins SM, Woodhead JD, Kinsley LPJ, et al. (1997) A simple method for the precise determination of ≥40 trace elements in geological samples by ICPMS using enriched isotope internal standardisation. *Chemical Geology* 134: 311–326.

Ely JC and Neal CR (2003) Using platinum-group elements to investigate the origin of the Ontong Java Plateau, SW Pacific. *Chemical Geology* 196: 235–257.

Farago ME, Kavanagh P, Blanks R, et al. (1996) Platinum metal concentrations in urban road dust and soil in the United Kingdom. *Fresenius' Journal of Analytical Chemistry* 354: 660–663.

Fedorowich J, Richards J, Jain J, Kerrich R, and Fan J (1993) A rapid method for REE and trace-element analysis using laser sampling ICP-MS on direct fusion wholerock glasses. *Chemical Geology* 106: 229–249. Feng D, Chen DF, and Roberts HH (2009) Petrographic and geochemical characterization of seep carbonate from Bush Hill (GC 185) gas vent and hydrate site of the Gulf of Mexico. *Marine and Petroleum Geology* 26: 1190–1198.

Florian D, Barnes RM, and Knapp G (1998) Comparison of microwave-assisted acid leaching techniques for the determination of heavy metals in sediments, soils, and sludges. *Fresenius' Journal of Analytical Chemistry* 362: 558–565.

Frimpong A, Fryer BJ, Longerich HP, Chen Z, and Jackson SE (1995) Recovery of precious metals using nickel sulfide fire assay collection – Problems at nanogram per gram concentrations. *Analyst* 120: 1675–1680.

Fritsche J and Meisel T (2004) Determination of anthropogenic input of Ru, Rh, Pd, Re, Os, Ir and Pt in soils along Austrian motorways by isotope dilution ICP-MS. *Science* of the Total Environment 325: 145–154.

Fujimori E, Ichikawa K, Shiozawa R, Asai K, Chiba K, and Haraguchi H (2001) Multielement determination and distributions of major-to-ultratrace elements in industrial waste incineration bottom ash as studied by ICP-AES and ICP-MS. *Bulletin of the Chemical Society of Japan* 74: 2045–2051.

Gaines P (2012) Trace analysis guide – A guide for attaining reliable measurements. Inorganic Ventures, Inc. http://www.inorganicventures.com/tech/samplepreparation/default.

Gilman LB and Engelhart WG (1989) Recent advances in microwave sample preparation. Sprctroscopy 4(8): 14–21.

Gordon CL (1943) Modification of the Carius combustion tube to minimize losses by explosion: Pressures attained on heating nitric acid to 300 °C. *Journal of Research of the National Bureau of Standards* 30: 107–111.

Gordon CL, Schlecht WG, and Wichers E (1944) Use of sealed tubes for the preparation of acid solutions of samples for analysis, or for small-scale refining: Pressures of acids heated above 100 °C. *Journal of Research of the National Bureau of Standards* 33: 457–470.

Gorsuch TT (1970) *Destruction of Organic Matter*. New York: Pergamon Press. Govindaraju K and Mevelle G (1987) Fully automated dissolution and separation

methods for inductively coupled plasma atomic emission-spectrometry rock analysis-application to the determination of rare-earth elements – Plenary lecture. *Journal of Analytical Atomic Spectrometry* 2: 615–621.

Gray AL (1986) Mass spectrometry with an inductively coupled plasma as an ion source: The influence on ultra-trace analysis of background and matrix response. *Spectrochimica Acta Part B* 41: 151–167.

Gros M, Lorand JP, and Luguet A (2002) Analysis of platinum group elements and gold in geological materials using NiS fire assay and Te coprecipitation; the NiS dissolution step revisited. *Chemical Geology* 185: 179–190.

Guldas M (2008) Comparison of digestion methods and trace elements determination in chocolates with pistachio using atomic absorption spectrometry. *Journal of Food and Nutrition Research* 47: 92–99.

Gumann S, Lahaye Y, and Brey G (2003) Iridium strip – Rhyolithe enthüllen ihre details. Berichte des Europäischen Jahrbuchs für Mineralogie 15: 72.

Günther D, Quadt AV, Wirz R, Cousin H, and Dietrich VJ (2001) Elemental analyses using laser ablation-inductively coupled plasma-mass spectrometry (LA-ICP-MS) of geological samples fused with Li₂B₄0₇ and calibrated without matrix-matched standards. *Mikrochimica Acta* 136: 101–107.

Haffty J, Riley LB, and Goss WD (1977) *A Manual on Fire Assaying and Determination of the Noble Metals in Geological Materials. United States Geological Survey Bulletin* 1445: Washington, DC: US Government Printing Office.

Hall GEM and Pelchat JC (1994) Analysis of geological materials for gold, platinum and palladium at low ng/g levels by fire assay-ICP mass spectrometry. *Chemical Geology* 115: 61–72.

Hall GEM and Plant J (1992) Analytical errors in the determination of high field strength elements and their implications in tectonic interpretation studies. *Chemical Geology* 95: 141–156.

Hannaker P and Hou Q (1984) Dissolution of geological material with orthophosphoric acid for major-element determination by flame atomic-absorption spectroscopy and inductively coupled plasma atomic-emission spectroscopy. *Talanta* 31: 1153–1157.

Harju L, Lill JO, Saarela KE, Heselius SJ, Hernberg FJ, and Lindroos A (1997) Analysis of trace elements in trunk wood by thick-target PIXE using dry ashing for preconcentration. *Fresenius' Journal of Analytical Chemistry* 358: 523–528.

Heanes DL (1981) Determination of trace-elements in plant materials by a dry-ashing procedure. Part II. Copper, manganese and zinc. *Analyst* 106: 182–187.

Hoenig M (2001) Preparation steps in environmental trace element analysis – Facts and traps. *Talanta* 54: 1021–1038.

Hoenig M (2003) Dry ashing. In: Mester Z and Sturgeon R (eds.) Sample Preparation for Trace Element Analysis, pp. 235–255. The Netherlands: Elsevier.

Hu ZC and Gao S (2008) Upper crustal abundances of trace elements: A revision and update. *Chemical Geology* 253: 205–221.

- Hu ZC, Gao S, Hu SH, Yuan HL, Liu XM, and Liu YS (2005) Suppression of interferences for direct determination of arsenic in geological samples by inductively coupled plasma mass spectrometry. *Journal of Analytical Atomic Spectrometry* 20: 1263–1269.
- Hu ZC, Gao S, Liu YS, et al. (2010) NH₄F assisted high pressure digestion of geological samples for multi-element analysis by ICP-MS. *Journal of Analytical Atomic Spectrometry* 25: 408–413.
- Ingamell CO (1970) Lithium metaborate flux in silicate analysis. *Analytica Chimica Acta* 52: 323–334.
- Ito J (1962) A new method of decomposition for refractory minerals and its application to the determination of ferrous iron and alkalies. *Bulletin of the Chemical Society of Japan* 35: 225–228.
- Jackson SE, Fryer BJ, Gosse W, Healey DC, Longerich HP, and Strong DF (1990) Determination of the precious metals in geological-materials by inductively coupled plasma mass-spectrometry (ICP-MS) with nickel sulfide fire-assay collection and tellurium coprecipitation. *Chemical Geology* 83: 119–132.
- Jackwerth E and Gomiscek S (1984) Acid pressure decomposition in trace element analysis. *Pure and Applied Chemistry* 56: 478–489.
- Jain JC, Neal CR, and Hanchar JM (2001) Problems associated with the determination of rare earth elements of a "Gem" quality zircon by inductively coupled plasma-mass spectrometry. *Geostandards Newsletter* 25: 229–237.
- Jarvis KE (1988) Inductively coupled plasma mass spectrometry: A new technique for the rapid or ultra-trace level determination of the rare-earth elements in geological materials. *Chemical Geology* 68: 31–39.
- Jarvis KE (1990) A critical evaluation of two sample preparation techniques for low-level determination of some geologically incompatible elements by inductively coupled plasma-mass spectrometry. *Chemical Geology* 83: 89–103.
- Jarvis I (1992) Sample preparation for ICP-MS. In: Jarvis KE, Gray AL, and Houk RS (eds.) Handbook of Inductively Coupled Plasma-Mass Spectrometry, pp. 172–224. London: Blackie.
- Jarvis KE and Jarvis I (1988) Determination of the rare-earth elements and yttrium in 37 international silicate reference materials by inductively coupled plasma-atomic emission spectrometry. *Geostandards Newsletter* 12: 1–12.
- Jarvis I, Totland MM, and Jarvis KE (1997) Determination of the platinum-group elements in geological materials by ICP-MS using microwave digestion, alkali fusion and cation-exchange chromatography. *Chemical Geology* 143: 27–42.
- Jarvis KE, Williams JG, Parry SJ, and Bertalan E (1995) Quantitative-determination of the platinum-group elements and gold using NiS fire assay with laser ablation-inductively coupled plasma-mass spectrometry (LA-ICP-MS). *Chemical Geology* 124: 37–46.
- Jeffery PG (1975) *Chemical Methods for Rock Analysis*, 2nd edn., pp. 1–525. Oxford: Pergamon Press.
- Jenner GA, Longerich HP, Jackson SE, and Fryer BJ (1990) ICP-MS A powerful tool for high-precision trace-element analysis in earth sciences: Evidence from analysis of selected U.S.G.S. reference samples. *Chemical Geology* 83: 133–148.
- Jensen KK, Baker J, Waight T, Frei R, and Peate DW (2003) High precision Ru, Pd, Ir, Pt, Re and REE determinations in the Stevns Klint Cretaceous-Tertiary boundary reference material (FC-1) by isotope dilution multiple collector inductively coupled plasma-mass spectrometry. *Geostandards Newsletter* 27: 59–66.
- Jochum KP, Dingwell DB, Rocholl A, et al. (2000) The preparation and preliminary characterisation of eight geological MPI-DING reference glasses for *in-situ* microanalysis. *Geostandards Newsletter* 24: 87–133.
- Johnson WM and Maxwell JA (1981) Rock and Mineral Analysis, 2nd edn., pp. 1–489. New York: Wiley.
- Jorge APD, Enzweiler J, Shibuya EK, Sarkis JES, and Figueiredo AMG (1998) Platinumgroup elements and gold determination NiS fire assay buttons by UV laser ablation ICP-MS. *Geostandards Newsletter* 22: 47–55.
- Jorhem L (1993) Determination of metals in foodstuffs by atomic-absorption spectrophotometry after dry ashing – Nmkl interlaboratory study of lead, cadmium, zinc, copper, iron, chromium, and nickel. *Journal of AOAC International* 76: 798–813.
- Juvonen R, Kallio E, and Lakomaa T (1994) Determination of precious metals in rocks by inductively coupled plasma mass spectrometry using nickel sulfide concentration. Comparison with other pre-treatment methods. *Analyst* 119: 617–621.
- Juvonen R, Lakomaa T, and Soikkeli L (2002) Determination of gold and the platinum group elements in geological samples by ICP-MS after nickel sulphide fire assay: Difficulties encountered with different types of geological samples. *Talanta* 58: 595–603.
- Kaya G, Akdeniz I, and Yaman M (2008) Determination of Cu, Mn, and Pb in yogurt samples by flame atomic absorption spectrometry using dry, wet, and microwave ashing methods. *Atomic Spectroscopy* 29: 99–106.
- Khuder A, Ahmad M, Hasan R, and Saour G (2010) Improvement of X-ray fluorescence sensitivity by dry ashing method for elemental analysis of bee honey. *Microchemical Journal* 95: 152–157.

- Kingston HM and Haswell SJ (1997) Microwave-Enhanced Chemistry. Fundamentals, Sample Preparation and Applications, pp. 1–772. Washington, DC: American Chemical Society.
- Kingston HM and Jassie LB (1986) Microwave energy for acid decomposition at elevated temperatures and pressures using biological and botanical samples. *Analytical Chemistry* 58: 2534–2541.
- Kingston HM and Jassie LB (1988) *Introduction to Microwave Sample Preparation, Theory and Practice*, pp. 1–263. Washington, DC: American Chemical Society. Knapp G (1984) Routes to powerful methods of elemental trace analysis in
- environmental-samples. Fresenius' Journal of Analytical Chemistry 317: 213-219.
- Knapp G (1985) Sample preparation techniques An important part in trace-element analysis for environmental-research and monitoring. *International Journal of Environmental Analytical Chemistry* 22: 71–83.
- Knapp G and Schramel P (2003) Sources of analyte contamination and loss during the analytical process. In: Mester Z and Sturgeon R (eds.) Sample Preparation for Trace Element Analysis, pp. 23–45. The Netherlands: Elsevier.
- Koh S, Aoki T, Katayama Y, and Takada J (1999) Losses of elements in plant samples under the dry ashing process. *Journal of Radioanalytical and Nuclear Chemistry* 239: 591–594.
- Korn MDA, Morte ESD, dos Santos D, et al. (2008) Sample preparation for the determination of metals in food samples using spectroanalytical methods – A review. *Applied Spectroscopy Reviews* 43: 67–92.
- Kucera J and Soukal L (1995) Cadmium losses on wet and dry ashing of plant materials. *Journal of Radioanalytical and Nuclear Chemistry* 193: 33–38.
- Kurosawa M, Shima K, Ishii S, and Sasa K (2006) Trace element analysis of fused whole-rock glasses by laser ablation-ICP-MS and PIXE. *Geostandards and Geoanalytical Research* 30: 17–30.
- Kuss HM (1992) Applications of microwave digestion technique for elemental analyses. Fresenius' Journal of Analytical Chemistry 343: 788–793.
- Lamble KJ and Hill SJ (1998) Microwave digestion procedures for environmental matrices. *Analyst* 123: 103–133.
- Lamothe PJ, Fries TL, and Consul JJ (1986) Evaluation of a microwave-oven system for the dissolution of geologic samples. *Analytical Chemistry* 58: 1881–1886.
- Langmyhr FJ (1967) The removal of hydrofluoric acid by evaporation in the presence of sulfuric or perchloric acids. *Analytica Chimica Acta* 39: 516–518.
- Langmyhr FJ and Kringstad K (1966) An investigation of the composition of the precipitates formed by the decomposition of silicate rocks in 38–40% hydrofluoric acid. Analytica Chimica Acta 35: 131–135.
- Langmyhr FJ and Paus PE (1968) The analysis of inorganic siliceous materials by atomic absorption spectrophotometry and the hydrofluoric acid decomposition technique. Part I. The analysis of silicate rocks. *Analytica Chimica Acta* 43: 397–408.
- Langmyhr FJ and Sveen S (1965) Decomposability in hydrofluoric acid of the main and some minor and trace minerals of silicate rocks. *Analytica Chimica Acta* 32: 1–7.
- Lee CS, Li XD, Shi WZ, Cheung SC, and Thornton I (2006) Metal contamination in urban, suburban, and country park soils of Hong Kong: A study based on GIS and multivariate statistics. *Science of the Total Environment* 356: 45–61.
- Lee SB, Yoshimura E, Tanaka Y, Saitoh J, Yamazaki S, and Toda S (1986) Combination of wet charring and dry ashing suitable for the determination of trace-metals in oily foodstuffs by graphite-furnace AAS. *Bunseki Kagaku* 35: T120–T123.
- Li XL and Ebihara M (2003) Determination of all platinum-group elements in mantlederived xenoliths by neutron activation analysis with NiS fire-assay preconcentration. *Journal of Radioanalytical and Nuclear Chemistry* 255: 131–135.
- Li XL and Tong CH (1995) Neutron-activation analysis after a nickel sulfide fire assay preconcentration for determination of all platinum-group elements in rocks. *Journal* of Radioanalytical and Nuclear Chemistry 196: 11–14.
- Lightfoot PC and Keays RR (2005) Siderophile and chalcophile metal variations in flood basalts from the Siberian trap, Noril'sk region: Implications for the origin of the Ni–Cu–PGE sulfide ores. *Economic Geology* 100: 439–462.
- Llorente MJF and Garcia JEC (2006) Concentration of elements in woody and herbaceous biomass as a function of the dry ashing temperature. *Fuel* 85: 1273–1279.
- Longerich HP, Jenner GA, Fryer BJ, and Jackson SE (1990) Inductively coupled plasma–mass spectrometric analysis of geological samples: A critical evaluation based on case studies. *Chemical Geology* 83: 105–118.
- Luo XS, Yu S, and Li XD (2011) Distribution, availability, and sources of trace metals in different particle size fractions of urban soils in Hong Kong: Implications for assessing the risk to human health. *Environmental Pollution* 159: 1317–1326.
- Mader P, Haber V, and Zelinka J (1997) Classical dry ashing of biological and agricultural materials. 1. Following the course of removal of organic matrix. *Analusis* 25: 175–183.
- Mader P, Szakova J, and Miholova D (1998) Classical dry ashing of biological and agricultural materials. Part II. Losses of analytes due to their retention in an insoluble residue. *Analusis* 26: 121–129.

- Makishima A, Nakamura E, and Nakano T (1999) Determination of zirconium, niobium, hafnium and tantalum at ng g⁻¹ levels in geological materials by direct nebulization of sample HF solution into FI-ICP-MS. *Geostandards Newsletter* 23: 7–20.
- Makishima A, Tanaka R, and Nakamura E (2009) Precise elemental and isotopic analyses in silicate samples employing ICP-MS: Application of hydrofluoric acid solution and analytical techniques. *Analytical Sciences* 25: 1181–1187.
- Malde MK, Bjorvatn K, and Julshamn K (2001) Determination of fluoride in food by the use of alkali fusion and fluoride ion-selective electrode. *Food Chemistry* 73: 373–379.
- Malinovsky D, Rodushkin I, Baxter D, and Ohlander B (2002) Simplified method for the Re–Os dating of molybdenite using acid digestion and isotope dilution ICP-MS. *Analytica Chimica Acta* 463: 111–124.
- Maqueda C and Rodriquez JLP (1986) Problems in the dissolution of silicates by acid mixtures. Analyst 111: 1107–1108.
- Mariet C, Belhadj O, Leroy S, Carrot F, and Métricha N (2008) Relevance of NH₄F in acid digestion before ICP-MS analysis. *Talanta* 77: 445–450.
- Masson P, Dalix T, and Bussiere S (2010) Determination of major and trace elements in plant samples by inductively coupled plasma-mass spectrometry. *Communications* in Soil Science and Plant Analysis 41: 231–243.
- Matos-Reyes MN, Cervera ML, Campos RC, and de la Guardia M (2010) Total content of As, Sb, Se, Te and Bi in Spanish vegetables, cereals and pulses and estimation of the contribution of these foods to the Mediterranean daily intake of trace elements. *Food Chemistry* 122: 188–194.
- Matsumoto K and Koura T (2001) Rapid decomposition and analysis of refractory barium sulfate by fusion with ammonium salt. *Bunseki Kagaku* 50: 807–811.
- Matusiewicz H (1994) Development of a high pressure/temperature focused microwave heated Teflon bomb for sample preparation. *Analytical Chemistry* 66: 751–755.
- Matusiewicz H (2003) Wet digestion methods. In: Mester Z and Sturgeon R (eds.) *Sample Preparation for Trace Element Analysis*, pp. 193–233. The Netherlands: Elsevier.
- Matusiewicz H and Sturgeon RE (1989) Present status of microwave sample dissolution and decomposition for elemental analysis. *Progress in Analytical Spectroscopy* 12: 21–39.
- McDonald I, Hart RJ, and Tredoux M (1994) Determination of the platinum-group elements in South African kimberlites by nickel sulphide fire-assay and neutron activation analysis. *Analytica Chimica Acta* 289: 237–247.
- McGrath D (1998) Use of microwave digestion for estimation of heavy metal content of soils in a geochemical survey. *Talanta* 46: 439–448.
- Meisel T, Fellner N, and Moser J (2003) A simple procedure for the determination of platinum group elements and rhenium (Ru, Rh, Pd, Re, Os, Ir and Pt) using ID-ICP-MS with an inexpensive on-line matrix separation in geological and environmental materials. *Journal of Analytical Atomic Spectrometry* 18: 720–726.
- Meisel T and Moser J (2004a) Platinum-group element and rhenium concentrations in low abundance reference materials. *Geostandards and Geoanalytical Research* 28: 233–250.
- Meisel T and Moser J (2004b) Reference materials for geochemical PGE analysis: New analytical data for Ru, Rh, Pd, Os, Ir, Pt and Re by isotope dilution ICP-MS in 11 geological reference materials. *Chemical Geology* 208: 319–338.
- Meisel T, Moser J, Fellner N, Wegscheider W, and Schoenberg R (2001) Simplified method for the determination of Ru, Pd, Re, Os, Ir and Pt in chromitites and other geological materials by isotope dilution ICP-MS and acid digestion. *Analyst* 126: 322–328.
- Meisel T, Moser J, and Wegscheider W (2001) Recognizing heterogeneous distribution of platinum group elements (PGE) in geological materials by means of the Re–Os isotope system. *Fresenius' Journal of Analytical Chemistry* 370: 566–572.
- Meisel T, Reisberg L, Moser J, Carignan J, Melcher F, and Brugmann G (2003) Re–Os systematics of UB-N, a serpentinized peridotite reference material. *Chemical Geology* 201: 161–179.
- Meisel T, Schöner N, Paliulionyte V, and Kahr E (2002) Determination of rare earth elements, Y, Th, Zr, Hf, Nb and Ta in geological reference materials G-2, G-3, SCo-1 and WGB-1 by sodium peroxide sintering and inductively coupled plasma-mass spectrometry. *Geostandards Newsletter* 26: 53–61.
- Mihaljevic M, Sebek O, Lukesova E, and Bouzkova A (2001) Direct determination of boron and zirconium in ceramic materials by flame atomic absorption spectrometry after alkali sintering and fusion. *Fresenius' Journal of Analytical Chemistry* 371: 1158–1160.
- Mindak WR and Dolan SP (1999) Determination of arsenic and selenium in food using a microwave digestion—Dry ash preparation and flow injection hydride generation atomic absorption spectrometry. *Journal of Food Composition and Analysis* 12: 111–122.
- Ming Y and Bing L (1998) Determination of rare earth elements in human hair and wheat flour reference materials by inductively coupled plasma mass spectrometry with dry ashing and microwave digestion. *Spectrochimica Acta Part B* 53: 1447–1454.

- Mitsumata H and Aoki I (2004) Quantitative determination of barium, magnesium, aluminium and europium in blue phosphors for fluorescent lamp using ICP-AES after alkali fusion. *Journal of the Ceramic Society of Japan* 112: 608–611.
- Moody JR and Lindstrom RM (1977) Selection and cleaning of plastic containers for storage of trace element samples. *Analytical Chemistry* 49: 2264–2267.
- Morcelli CPR, Figueiredo AMG, Enzweiler J, Sarkis JES, Jorge APS, and Kakazu M (2004) Determination of platinum-group elements in geological reference materials by high resolution-ICP-MS after nickel sulfide fire-assay collection and Te co-precipitation. *Geostandards and Geoanalytical Research* 28: 305–310.
- Moreno IM, Gonzalez-Weller D, Gutierrez V, et al. (2008) Determination of Al, Ba, Ca, Cu, Fe, K, Mg, Mn, Na, Sr and Zn in red wine samples by inductively coupled plasma optical emission spectroscopy: Evaluation of preliminary sample treatments. *Microchemical Journal* 88: 56–61.
- Münker C (1998) Nb/Ta fractionation in a Cambrian arc/back arc system, New Zealand: Source constraints and application of refined ICPMS techniques. *Chemical Geology* 144: 23–45.
- Murphy TJ (1974) The role of the analytical blank in accurate trace analysis. In: National Bureau of Standards Special Publication 422, Accuracy in Trace Analysis, Sample Handling, and Analysis, Proceedings of the 7th IMR Symposium, Gaithersburg, MD, October 7–11, pp. 509–539.
- Mutsuga M, Sato K, Hirahara Y, and Kawamura Y (2011) Analytical methods for SiO₂ and other inorganic oxides in titanium dioxide or certain silicates for food additive specifications. *Food Additives and Contaminants. Part A* 28: 423–427.
- Nadkarni RA (1984) Application of microwave sample dissolution in analysis. *Analytical Chemistry* 56: 2233–2237.
- Nandy AK, Manjhi JK, and Roy NK (2008) Stream sediment and soil samples by microwave digestion followed by ICP-MS measurement. *Atomic Spectroscopy* 29: 115–123.
- Navarro MS, Andrade S, Ulbrich H, Gomes CB, and Girardi VAV (2008) The direct determination of rare earth elements in basaltic and related rocks using ICP-MS: Testing the efficiency of microwave oven sample decomposition procedures. *Geostandards and Geoanalytical Research* 32: 167–180.
- Neas ED and Collins MJ (1988) Microwave heating. Theoretical concepts and equipment design. In: Kingston HM and Jassie LB (eds.). *Introduction to Microwave Sample Preparation, Theory and Practice*, pp. 7–32. Washington, DC: American Chemical Society.
- Nehring F, Jacob DE, Barth MG, and Foley SF (2008) Laser-ablation ICP-MS analysis of siliceous rock glasses fused on an iridium strip heater using MgO dilution. *Microchimica Acta* 160: 153–163.
- Nicholls IA (1974) A direct fusion method of preparing silicate rock glasses for energydispersive electron microprobe analysis. *Chemical Geology* 14: 151–157.
- Nna-Mvondo D, Martin-Redondo M, and Martinez-Frias J (2008) New application of microwave digestion-inductively coupled plasma-mass spectrometry for multi-element analysis in komatiites. *Analytica Chimica Acta* 628: 133–142.
- Nóbrega JA, Santos MC, Sousa RA, Cadore S, Barnes RM, and Tatro M (2006) Sample preparation in alkaline media. Spectrochimica Acta Part B 61: 465–495.
- Norman M, Bennett V, McCulloch M, and Kinsley L (2002) Osmium isotopic compositions by vapor phase sample introduction using a multi-collector ICP-MS. *Journal of Analytical Atomic Spectrometry* 17: 1394–1397.
- Norman MD, Pearson NJ, Sharma A, and Griffin WL (1996) Quantitative analysis of trace elements in geological materials by laser ablation ICPMS: Instrumental operating conditions and calibration values of NIST glasses. *Geostandards Newsletter* 20: 247–261.
- Norrish K and Hutton JT (1969) An accurate X-ray spectrographic method for the analysis of a wide range of geological samples. *Geochimica et Cosmochimica Acta* 33: 431–453.
- Ødegård M, Dundas SH, Flem B, and Grimstvedt A (1998) Application of a doublefocusing magnetic sector inductively coupled plasma-mass spectrometer with laser ablation for the bulk analysis of rare earth elements in rocks fused with Li₂B₄O₇. *Fresenius' Journal of Analytical Chemistry* 362: 477–482.
- Ødegård M and Hamester M (1997) Preliminary investigation into a high resolution inductively coupled plasma-mass spectrometer with laser ablation for bulk analysis of geological materials fused with Li₂B₄O₇. *Geostandards Newsletter* 21: 245–252.
- Oguri K, Shimoda G, and Tatsumi Y (1999) Quantitative determination of gold and the platinum-group elements in geological samples using improved NiS fire-assay and tellurium coprecipitation with inductively coupled plasma-mass spectrometry (ICP-MS). *Chemical Geology* 157: 189–197.
- Ohlweile OA, Meditsch JO, and Piatnick CMS (1973) Fusion with boron trioxide for silicate analysis by atomic-absorption spectrometry – Determination of potassium. *Analytica Chimica Acta* 67: 283–288.
- Olive V, Ellam RM, and Wilson L (2001) A protocol for the determination of the rare earth elements at picomole level in rocks by ICP-MS: Results on geological reference materials USGS PCC-1 and DTS-1. *Geostandards Newsletter* 25: 219–228.

Pack A, Russell SS, Michael J, Shelley G, and Zuilen MV (2007) Geo- and cosmochemistry of the twin elements yttrium and holmium. *Geochimica et Cosmochimica Acta* 71: 4592–4608.

Panteeva SV, Gladkochoub DP, Donskaya TV, Markova VV, and Sandimirova GP (2003) Determination of 24 trace elements in felsic rocks by inductively coupled plasma mass spectrometry after lithium metaborate fusion. *Spectrochimica Acta Part B* 58: 341–350.

Pearson DG and Woodland SJ (2000) Solvent extraction/anion exchange separation and determination of PGEs (Os, Ir, Pt, Pd, Ru) and Re–Os isotopes in geological samples by isotope dilution ICP-MS. *Chemical Geology* 165: 87–107.

Perkins WT, Pearce NJG, and Jeffries TE (1993) Laser ablation-inductively coupled plasma-mass spectrometry: A new technique for the determination of trace and ultra-trace elements in silicates. *Geochimica et Cosmochimica Acta* 57: 475–482.

Pichler U, Haase A, Knapp G, and Michaelis M (1999) Microwave-enhanced flow system for high-temperature digestion of resistant organic materials. *Analytical Chemistry* 71: 4050–4055.

Pickhardt G, Becker JS, and Dietze H-J (2000) A new strategy of solution calibration in laser ablation inductively coupled plasma-mass spectrometry for multielement trace analysis of geological samples. *Fresenius' Journal of Analytical Chemistry* 368: 173–181.

Pin C and Joannon S (1997) Low-level analysis of lanthanides in eleven silicate rock reference materials by ICP-MS after group separation using cation-exchange chromotography. *Geostandards Newsletter* 21: 43–50.

Plessen HG and Erzinger J (1998) Determination of the platinum-group elements and gold in twenty rock reference materials by inductively coupled plasma-mass spectrometry (ICP-MS) after pre-concentration by nickel sulfide fire assay. *Geostandards Newsletter* 22: 187–194.

Potts PJ (1987) A Handbook of Silicate Rock Analysis, pp. 1–622. Glasgow: Blackie.

Potts PJ, Gowing CJB, and Govindaraju K (1992) Preparation, homogeneity evaluation and cooperative study of two new chromitite reference samples CHR-Pt⁺ and CHR-Bkg. *Geostandards Newsletter* 16: 81–108.

Potts PJ and Robinson P (2003) Sample preparation of geological samples, soils and sediments. In: Mester Z and Sturgeon R (eds.) Sample Preparation for Trace Element Analysis, pp. 723–763. The Netherlands: Elsevier.

Pretorius W, Chipley D, Kyser K, and Helmstaedt H (2003) Direct determination of trace levels of Os, Ir, Ru, Pt and Re in kimberlite and other geological materials using HR-ICP-MS. *Journal of Analytical Atomic Spectrometry* 18: 302–309.

Pretorius W, Weis D, Williams G, Hanano D, Kieffer B, and Scoates J (2006) Complete trace elemental characterization of granitoid (USGA G-2, GSP-2) reference materials by high resolution inductively coupled plasma-mass spectrometry. *Geostandards* and *Geoanalytical Research* 30: 39–54.

Qi L and Grégoire DC (2000) Determination of trace elements in twenty six Chinese geochemistry reference materials by inductively coupled plasma-mass spectrometry. *Geostandards Newsletter* 24: 51–63.

Qi L, Gregoire DC, Zhou MF, and Malpas J (2003) Determination of Pt, Pd, Ru and Ir in geological samples by ID-ICP-MS using sodium peroxide fusion and Te coprecipitation. *Geochemical Journal* 37: 557–565.

Qi L, Hu J, and Grégoire DC (2000) Determination of trace elements in granites by inductively coupled plasma mass spectrometry. *Talanta* 51: 507–513.

Qi L and Zhou MF (2008) Determination of platinum-group elements in OPY-1: Comparison of results using different digestion techniques. *Geostandards and Geoanalytical Research* 32: 377–387.

Qi L, Zhou MF, Gao JF, and Zhao Z (2010) An improved Carius tube technique for determination of low concentrations of Re and Os in pyrites. *Journal of Analytical Atomic Spectrometry* 25: 585–589.

Qi L, Zhou MF, Wang CY, and Sun M (2007) Evaluation of a technique for determining Re and PGEs in geological samples by ICP-MS coupled with a modified Carius tube digestion. *Geochemical Journal* 41: 407–414.

Qu WJ, Du AD, and Jing R (2008) Influence of H₂O₂ on the signal intensity of rhenium, osmium and Re–Os age in the process of dissolution for pyrite. *Chinese Journal of Analytical Chemistry* 36: 223–226.

Rafter TA (1950) Sodium peroxide decomposition of minerals in platinum vessels. *Analyst* 75: 485–492.

Ravizza G and Pyle D (1997) PGE and Os isotopic analyses of single sample aliquots with NiS fire assay preconcentration. *Chemical Geology* 141: 251–268.

Rehkamper M, Halliday AN, and Wentz RF (1998) Low-blank digestion of geological samples for platinum-group element analysis using a modified Carius tube design. *Fresenius' Journal of Analytical Chemistry* 361: 217–219.

Reid JE, Horn I, Longerich HP, Forsthye L, and Jenner GA (1999) Determination of Zr and Hf in a flux-free fusion of whole rock samples using laser ablation inductively coupled mass spectrometry (LA-ICP-MS) with isotope dilution calibration. *Geostandards Newsletter* 23: 149–155. Reimann C, Birke M, and Filzmoser P (2010) Bottled drinking water: Water contamination from bottle materials (glass, hard PET, soft PET), the influence of colour and acidification. *Applied Geochemistry* 25: 1030–1046.

Reimann C, Grimstvedt A, Frengstad B, and Finne TE (2007) White HDPE bottles as source of serious contamination of water samples with Ba and Zn. *Science of the Total Environment* 374: 292–296.

Reimann C, Siewers U, Skarphagen H, and Banks D (1999) Does bottle type and acidwashing influence trace element analyses by ICP-MS on water samples? A test covering 62 elements and four bottle types: High density polyethene (HDPE), polypropylene (PP), fluorinated ethene propene copolymer (FEP) and perfluoroalkoxy polymer (PFA). *Science of the Total Environment* 239: 111–130.

Révillon S and Hureau-Mazaudier D (2009) Improvements in digestion protocols for trace element and isotope determinations in stream and lake sediment reference materials (JSd-1, JSd-2, JSd-3, JLk-1 and LKSD-1). *Geostandards and Geoanalytical Research* 33: 397–413.

Robinson P, Higgins NC, and Jenner GA (1986) Determination of rare-earth elements, yttrium and scandium in rocks by an ion exchange-X-ray fluorescence technique. *Chemical Geology* 55: 121–137.

Robinson P, Townsend AT, Yu Z, and Munker C (1999) Determination of scandium, yttrium and rare earth elements in rocks by high resolution inductively coupled plasma-mass spectrometry. *Geostandards Newsletter* 23: 31–46.

Rosner M, Pritzkow W, VogI J, and Voerkelius S (2011) Development and validation of a method to determine the boron isotopic composition of crop plants. *Analytical Chemistry* 83: 2562–2568.

Roy P, Balaram V, Kumar A, Satyanarayanan M, and Gnaneshwar Rao T (2007) New REE and trace element data on two kimberlitic reference materials by ICP-MS. *Geostandards and Geoanalytical Research* 31: 261–273.

Rucklidge JC, Gibb FGF, Fawcett JJ, and Gasparrini EL (1970) Rapid rock analysis by electron microprobe. *Geochimica et Cosmochimica Acta* 34: 243–247.

Saarela KE, Lill JO, Hernberg FJ, Harju L, Lindroos A, and Heselius SJ (1995) Preconcentration of trace elements in biological materials by dry ashing for TTPIXE analysis – A study of matrix effects. *Nuclear Instruments and Methods in Physics Research Section B: Beam Interactions with Materials and Atoms* 103: 466–472.

Sahayam AC, Chaurasia SC, and Venkateswarlu G (2010) Dry ashing of organic rich matrices with palladium for the determination of arsenic using inductively coupled plasma-mass spectrometry. *Analytica Chimica Acta* 661: 17–19.

Saracoglu S, Saygi KO, Uluozlu OD, Tuzen M, and Soylak M (2007) Determination of trace element contents of baby foods from Turkey. *Food Chemistry* 105: 280–285.

Sastre J, Sahuquillo A, Vidal M, and Rauret G (2002) Determination of Cd, Cu, Pb and Zn in environmental samples: Microwave-assisted total digestion versus aqua regia and nitric acid extraction. *Analytica Chimica Acta* 462: 59–72.

Seelye FT and Rafter TA (1950) Low-temperature decomposition of rocks, ores and minerals by sodium peroxide using platinum vessels. *Nature* 165: 317.

Sekimoto A, Horie K, Matsumoto T, and Haraguchi H (2002) Multielement determination of major-to-ultratrace elements in sediment, rock and clay samples by ICP-AES and ICP-MS. *Bunseki Kagaku* 51: 1075–1084.

Sen Gupta JG and Bertrand NB (1995a) Direct ICP-MS determination of trace and ultratrace elements in geological materials after decomposition in a microwave oven. I. Quantitation of Y, Th, U and the lanthanides. *Talanta* 42: 1595–1607.

Sen Gupta JG and Bertrand NB (1995b) Direct ICP-MS determination of trace and ultratrace elements in geological materials after decomposition in a microwave oven. Part II. Quantitation of Ba, Cs, Ga, Hf, In, Mo, Nb, Pb, Rb, Sn, Sr, Ta and TI. *Talanta* 42: 1947–1957.

Sen Gupta JG and Bouvier JL (1995) Direct determination of traces of Ag, Cd, Pb, Bi, Cr, Mn, Co, Ni, Li, Be, Cu and Sb in environmental waters and geological materials by simultaneous multi-element graphite furnace atomic absorption spectrometry with Zeeman-effect background correction. *Talanta* 42: 269–281.

Shimizu K, Chang Q, and Nakamura K (2011) Flux-free fusion of silicate rock preceding acid digestion for ICP-MS bulk analysis. *Geostandards and Geoanalytical Research* 35: 45–55.

Shirey SB and Walker RJ (1995) Carius tube digestion for low-blank rhenium–osmium analysis. Analytical Chemistry 67: 2136–2141.

Shotyk W, Krachler M, and Chen B (2006) Contamination of Canadian and European bottled waters with antimony from PET containers. *Journal of Environmental Monitoring* 8: 288–292.

Smith FE and Arsenault EA (1996) Microwave-assisted sample preparation in analytical chemistry. *Talanta* 43: 1207–1268.

Soylak M, Tuzen M, Narin I, and Sari H (2004) Comparison of microwave, dry and wet digestion procedures for the determination of trace metal contents in spice samples produced in Turkey. *Journal of Food and Drug Analysis* 12: 254–258.

Srogi K (2006) A review: Application of microwave techniques for environmental analytical chemistry. *Analytical Letters* 39: 1261–1288. Srogi K (2007) Microwave-assisted sample preparation of coal and coal fly ash for subsequent metal determination. *Analytical Letters* 40: 199–232.

Stoll B, Jochum KP, Herwig K, et al. (2008) An automated iridium-strip heater for LA-ICP-MS bulk analysis of geological samples. *Geostandards and Geoanalytical Research* 32: 5–26.

- Stone WE and Crocket JH (1993) Determination of noble and allied trace-metals using radiochemical neutron-activation analysis with tellurium coprecipitation. *Chemical Geology* 106: 219–228.
- Šulcek Z and Povondra P (1989) Methods of Decomposition in Inorganic Analysis, pp. 1–325. Boca Raton, FL: CRC Press.
- Sun YC, Chi PH, and Shiue MY (2001a) Comparison of different digestion methods for total decomposition of siliceous and organic environmental samples. *Analytical Sciences* 17: 1395–1399.

Sun YL, Chu ZY, Sun M, and Xia XP (2009) An improved Fe–Ni sulfide fire assay method for determination of Re, platinum group elements, and Os isotopic ratios by inductively coupled plasma- and negative thermal ionization-mass spectrometry. *Applied Spectroscopy* 63: 1232–1237.

Sun YL, Guan XY, and Du AD (1998) Determination of platinum group elements by inductively coupled plasma-mass spectrometry combined with nickel sulfide fire assay and tellurium coprecipitation. *Spectrochimica Acta Part B* 53: 1463–1467.

Sun M, Jain J, Zhou MF, and Kerrich R (1993) A procedural modification for enhanced recovery of precious metals (gold, PGE) following nickel sulfide fire assay and tellurium co-precipitation: Applications for analysis of geological samples by inductively coupled plasma mass spectrometry. *Canadian Journal of Applied Spectroscopy* 38: 103–108.

Sun YL and Sun M (2005) Nickel sulfide fire assay improved for pre-concentration of platinum group elements in geological samples: A practical means of ultra-trace analysis combined with inductively coupled plasma-mass spectrometry. *Analyst* 130: 664–669.

Sun YL, Zhou MF, and Sun M (2001b) Routine Os analysis by isotope dilutioninductively coupled plasma mass spectrometry: OsO₄ in water solution gives high sensitivity. *Journal of Analytical Atomic Spectrometry* 16: 345–349.

Suzuki T and Sensui M (1991) Application of the microwave acid digestion method to the decomposition of rock samples. *Analytica Chimica Acta* 245: 43–48.

Sylvester PJ (2001) Trace element analysis of fused whole rock glasses by laser ablation ICP-MS. In: Sylvester PJ (ed.) Laser Ablation-ICP-MS in the Earth Sciences: Principle and Applications. Short Course Series, Vol., 29, pp. 147–162. Québec: Mineralogical Association of Canada.

Takei H, Yokoyama T, Makishima A, and Nakamura E (2001) Formation and suppression of AIF₃ during HF digestion of rock samples in Teflon bomb for precise trace element analyses by ICP-MS and ID-TIMS. *Proceedings of the Japan Academy, Series B* 77: 13–17.

Tan SH and Horlick G (1986) Background spectral features in inductively coupled plasma mass spectrometry. *Applied Spectroscopy* 40: 445–460.

Tanaka R, Makishima A, Kitagawa H, and Nakamura E (2003) Suppression of Zr, Nb, Hf and Ta coprecipitation in fluoride compounds for determination in Ca-rich materials. *Journal of Analytical Atomic Spectrometry* 18: 1458–1463.

Tang YQ, Jarvis KE, and Williams JG (1992) Determination of trace elements in 11 Chinese geological reference materials by ICP-MS. *Geostandards Newsletter* 16: 61–70.

Tarafder PK, Khorge CR, and Saran R (1997) Fluorimetric determination of fluoride. *Chemia Analityczna (Warsaw)* 42: 391–396.

Taylor VF, Toms A, and Longerich HP (2002) Acid digestion of geological and environmental samples using open-vessel focused microwave digestion. *Analytical* and Bioanalytical Chemistry 372: 360–365.

Terashima S, Okai T, and Taniguchi M (1998) Alkali fusion/ICP-AES for the determination of boron in geological reference materials and a comparison with spectrophotometry. *Bunseki Kagaku* 47: 451–454.

Tonarini S, Pennisi M, and Leeman WP (1997) Precise boron isotopic analysis of complex silicate (rock) samples using alkali carbonate fusion and ion-exchange separation. *Chemical Geology* 142: 129–137.

Totland M, Jarvis I, and Jarvis K (1992) An assessment of dissolution techniques for the analysis of geological samples by plasma spectrometry. *Chemical Geology* 95: 35–62.

Totland MM, Jarvis I, and Jarvis KE (1995) Microwave digestion and alkali fusion procedures for the determination of the platinum-group elements and gold in geological-materials by ICP-MS. *Chemical Geology* 124: 21–36.

Tschoepel P and Toelg G (1982) Comments on the accuracy of analytical results in nanogram and picogram trace analysis of the elements. *Journal of Trace and Microprobe Techniques* 1: 1–77.

Tuzen M (2003) Determination of heavy metals in fish samples of the middle Black Sea (Turkey) by graphite furnace atomic absorption spectrometry. *Food Chemistry* 80: 119–123.

Tuzen M, Sari H, and Soylak M (2004) Microwave and wet digestion procedures for atomic absorption spectrometric determination of trace metals contents of sediment samples. *Analytical Letters* 37: 1925–1936.

- Tuzen M, Silici S, Mendil D, and Soylak M (2007) Trace element levels in honeys from different regions of Turkey. *Food Chemistry* 103: 325–330.
- Vaisanen A, Laatikainen P, Ilander A, and Renvall S (2008) Determination of mineral and trace element concentrations in pine needles by ICP-AES: Evaluation of different sample pre-treatment methods. *International Journal of Environmental Analytical Chemistry* 88: 1005–1016.

Van Eenbergen A and Bruninx E (1978) Losses of elements during sample decomposition in an acid-digestion bomb. Analytica Chimica Acta 98: 405–406.

Van Loon JC and Barefoot RR (1991) *Determination of the Precious Metals*. New York: Wiley.

Van Paemel MR, De Rycke H, Millet S, Hesta M, and Janssens GPJ (2005) Evaluation of dry ashing in conjunction with ion chromatographic determination of transition metal ions in pig feed samples. *Journal of Agricultural and Food Chemistry* 53: 1873–1877.

Vassileva E, Docekalova H, Baeten H, Vanhentenrijk S, and Hoenig M (2001) Revisitation of mineralization modes for arsenic and selenium determinations in environmental samples. *Talanta* 54: 187–196.

Vogl J, Rosner M, and Pritzkow W (2011) Development and validation of a single collector SF-ICPMS procedure for the determination of boron isotope ratios in water and food samples. *Journal of Analytical Atomic Spectrometry* 26: 861–869.

Walsh JN, Gill R, and Thirlwall MF (1997) Dissolution procedures for geological and environmental samples. In: Gill R (ed.) *Modern Analytical Geochemistry*, pp. 29–40. Edinburgh Gate, Harlow: Pearson Education Limited.

Wang YM, Gao YS, Wang XH, Wu SQ, and Gu TX (2004) Investigations into the preparation of ultra-fine particle size geochemical reference materials. *Geostandards* and *Geoanalytical Research* 28: 113–121.

Wang L, He HL, and Li B (2003) Multi-element determination in geological samples by inductively coupled plasma mass spectrometry after fusion–precipitation treatment. *Rock and Mineral Analysis* 22: 86–92.

Wang CF, Huang MF, Chang EE, and Chiang PC (1996) Assessment of closed vessel digestion methods for elemental determination of airborne particulate matter by ICP-AES. *Analytical Sciences* 12: 201–207.

Wang XP and Li B (2010) Analysis of 27 mineral elements in the rice samples collected from China and Japan by using ICP-AES and ICP-MS. Spectroscopy Spectral Analysis 30: 2260–2264.

Wang J, Nakazato T, Sakanishi K, Yamada O, Tao H, and Saito I (2006) Single-step microwave digestion with HNO₃ alone for determination of trace elements in coal by ICP spectrometry. *Talanta* 68: 1584–1590.

Watanabe M (2001) Addition of iron (III) oxide in the alkali fusion of silicon carbide prepared by chemical vapor deposition for a highly precise determination of total silicon. *Bunseki Kagaku* 50: 263–266.

Watanabe M and Otsuki S (2008) Highly precise determination of silica in silicate by alkali fusion/coagulation gravimetric method. *Bunseki Kagaku* 57: 31–34.

Wei R and Haraguchi H (1999) Multielement determination of major-to-ultratrace elements in river and marine sediment reference materials by inductively coupled plasma atomic emission spectrometry and inductively coupled plasma mass spectrometry. *Analytical Sciences* 15: 729–735.

Wichers E, Schlecht WG, and Gordon CL (1944) Attack of refractory platiniferous materials by acid mixtures at elevated temperatures. *Journal of Research of the National Bureau of Standards* 33: 363–381.

Wichmann H and Bahadir M (2001) Determination of the noble metals platinum, rhodium, and palladium in airborne particles. *Fresenius Environmental Bulletin* 10: 106–108.

Wu S, Zhao YH, and Fang X (1996) Application of inductively coupled plasma-mass spectrometry for total metal determination in silicon containing solid samples using the microwave-assisted nitric acid-hydrofluoric acid-hydrogen peroxide-boric acid digestion system. *Journal of Analytical Atomic Spectrometry* 11: 287–296.

Xie QL and Kerrich R (1995) Application of isotope dilution for precise measurement of Zr and Hf in low-abundance samples and international reference materials by inductively coupled plasma mass spectrometry: Implications for Zr (Hf)/REE fractionations in komatiites. *Chemical Geology* 123: 17–27.

Xu YH, Iwashita A, Nakajima T, Yamashita H, Takanashi H, and Ohki A (2005) Effect of HF addition on the microwave-assisted acid-digestion for the determination of metals in coal by inductively coupled plasma-atomic emission spectrometry. *Talanta* 66: 58–64.

Yamada H, Sugahara M, Kosaka H, Katayama A, Takahashi K, and Yonebayashi K (1996) Determination of total and water soluble iodine in soil by high performance liquid chromatography. *Soil Science and Plant Nutrition* 42: 367–374.

Yamagata C, Andrade JB, Ussui V, Lima NB, and Paschoal JOA (2008) High purity zirconia and silica powders via wet process: Alkali fusion of zircon sand. *Materials Science Forum* 591–593: 771–776.

Yamaguchi H, Itoh S, Igarashi S, Naitoh K, and Hasegawa R (2000) Determination of trace impurities in graphite and silicon carbide by total reflection X-ray fluorescence spectrometry after homogeneous liquid–liquid extraction. *ISIJ International* 40: 779–782.

- Yaman M and Cokol N (2004) Determination of trace elements in human milk, cow's milk, and baby foods by flame AAS using wet ashing and microwave oven sample digestion procedures. *Atomic Spectroscopy* 25: 185–190.
- Yang G, Du AD, Lu J, Qu WJ, and Chen JF (2005) Re–Os (ICP-MS) dating of the massive sulfide ores from the Jinchuan Ni–Cu–PGE deposit. *Science in China Series D* 48: 1672–1677.
- Ybanez N, Cervera ML, and Montoro R (1992) Determination of arsenic in dry ashed seafood products by hydride generation atomic-absorption spectrometry and a critical comparative-study with platform furnace zeeman-effect atomic-absorption spectrometry and inductively coupled plasma atomic emission-spectrometry. *Analytica Chimica Acta* 258: 61–71.
- Yokoyama T, Makishima A, and Nakamura E (1999) Evaluation of the coprecipitation of incompatible trace elements with fluoride during silicate rock dissolution by acid digestion. *Chemical Geology* 157: 175–187.
- Yoshida S, Muramatsu Y, Tagami T, and Uchida S (1996) Determination of major and trace elements in Japanese rock reference samples by ICP-MS. *International Journal* of Mass Spectrometry 63: 195–206.
- Yu ZS, Norman MD, and Robinson P (2003) Major and trace element analysis of silicate rocks by XRF and laser ablation ICP-MS using lithium borate fused glasses: Matrix effects, instrument response and results for international reference materials. *Geostandards Newsletter* 27: 67–89.
- Yu ZS, Robinson P, and McGoldrick P (2001) An evaluation of methods for the chemical decomposition of geological materials for trace element determination using ICP-MS. *Geostandards Newsletter* 25: 199–217.

- Yu ZS, Robinson P, Townsend AT, Munker C, and Crawford AJ (2000) Determination of high field strength elements, Rb, Sr, Mo, Sb, Cs, TI and Bi at ng levels in geological reference materials by magnetic sector ICPMS after HF-HCIO₄ high pressure digestion. *Geostandards Newsletter* 24: 39–50.
- Yuan HL, Hu SH, Tong J, Zhao LH, Lin SL, and Gao S (2000) Preparation of ultra-pure water and acids and investigation of background of an ICP-MS laboratory. *Talanta* 52: 971–981.
- Zachariadis GA, Stratis JA, Kaniou I, and Kalligas G (1995) Critical comparison of wet and dry digestion procedures for trace-metal analysis of meat and fish-tissues. *Mikrochimica Acta* 119: 191–198.
- Zereini F, Skerstupp B, and Urban H (1994) Comparison between the use of sodium and lithium tetraborate in platinum-group element determination by nickel sulfide fireassay. *Geostandards Newsletter* 18: 105–109.
- Zhang W, Hu ZC, Liu YS, et al. (2012) Reassessment of HF/HNO₃ decomposition capability in high pressure digestion of felsic rocks for multi-element analysis by ICP-MS. *Geostandards and Geoanalytical Research* 36: 271–289.
- Zhou MF, Malpas J, Sun M, Liu Y, and Fu X (2001) A new method to correct Ni- and Cu-argide interference in the determination of the platinum-group elements, Ru, Rh, and Pd, by ICP-MS. *Geochemical Journal* 35: 413–420.
- Zhu LY, Liu YS, Hu ZC, Gao S, Wang XH, and Tian T (2011) Quick and accurate LA-ICP-MS analyses of major and trace elements in basalt, coupled with glass fusion method by double iridium strip heater. *Geochimica* 40: 407–417.