

壳幔混合及花岗质岩浆的形成

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摘 要: 花岗岩成因与壳幔作用有着密切的联系, 花岗岩成分不单来源于地壳, 还有地幔物质的参与, 不同程度的壳幔混合可形成类型多样的花岗质岩石。本文简要总结了在俯冲作用、底侵作用、折沉作用及地幔柱活动过程中花岗质岩浆的形成, 并对此类花岗岩的岩石学及地球化学特征进行了描述。鉴于此, 可初步判断花岗岩的形成是否与壳幔相互作用有关, 但还没有一种统一的模式能对其进行有效的解释。

关 键 词: 花岗岩; 俯冲作用; 底侵作用; 混合; 壳-幔作用

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The Crust—Mantle Mixing and the Generation of Granitic Magmas

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Abstract: Components of granite maybe not only derived from the crust but also from the mantle. The genesis of granites has close relationship with the crust—mantle interactions, thus the crust-mantle mixing has been used to study the genesis of granites recently by many researchers. It is believed that the crust-mantle mixing in various extents might form many kinds of granitoids. This article has summarized the generation of granitic magmas in subduction process, underplating process, detachment process and mantle pluming process, and has described the petrological and geochemical characteristics of granites. These features could help us to basically judge if the genesis of granites is related to the crust-mantle interaction, though there is no uniform model which could explain the formation of granites effectively.

Key words: granite; subduction; underplating; mixing; crust—mantle interaction

目前对花岗岩的研究着重于壳幔作用与花岗岩的成因; 将花岗岩的形成与壳幔作用相联系将是 21 世纪前 20 年的研究重点^[1]。

Castro 等^[2]提出了花岗岩的 S、M 和 H 型分类, 其中 H 型即混合花岗岩(相当于壳-幔型花岗岩), 并建议以 H 型花岗岩代替 I 型花岗岩。20 世纪 90 年代以来, 业已认识到大多数花岗岩是地壳和地幔相互作用的产物^[3]; 地幔在提供热源的同时, 地幔物

质也与地壳物质混熔形成花岗岩质岩浆^[4,5]。这个过程中, 地幔提供的热能起了十分重要的作用^[1]。

近代地球化学研究特别是 Sr、Nd、Pb 同位素和 REE 研究, 指出花岗岩普遍具有壳、幔源特征^[2]。花岗岩的形成不仅是地壳物质再循环的结果, 而且有幔源物质的参与, 最常见的就是基性和酸性成分的混合^[2], 混合后会形成无数种不同的花岗岩类型, 犹如连续谱系。不同来源的混合程度及经历过

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程的差别,造成了花岗岩类岩石的多样性^[1]。

1 混 合

实验岩石学研究^[6,7]表明, I型花岗岩中最常见的花岗闪长岩过于富钙不大可能源自泥质岩, 过于富钾也不大可能源自角闪质岩石, 只有特殊的岩石(如富石英角闪岩)在特定的条件下才可得到花岗闪长岩熔体^[8]。可见, I型花岗岩浆不是一个独立的岩浆, 而是 M型与 S型两个端员岩浆的混合物^[2]。另外, 很多 S型花岗岩也是岩浆混合的产物, A型花岗岩也是上地幔部分熔融体混染了地壳物质所成^[8]。因此, 壳幔物质混合是花岗岩形成中的普遍现象。

“混合”(Hybrid)一词最早由 Durocher(1857)和 Harker(1904)用于火成岩, 它包含岩浆(熔体、液体)间的混合、溶合, 也包括岩浆(溶体、液体)对固态岩石的同化混染。所谓“混合”即混熔、混染、混杂, 并且这三种作用是共生的、渐变的^[9], 在花岗岩形成中起着主导作用。

2 壳幔物质的混合

壳—幔相互作用包括俯冲作用、底侵作用、拆沉作用、火山、地幔柱活动等, 其中任何一种或几种同时作用都能引起壳幔物质的混合。俯冲、底侵还经常出现在同一过程的不同阶段: 一方面, 地壳物质下插到上地幔深度, 与地幔熔体混合形成花岗质岩浆; 另一方面, 基性地幔熔体侵入地壳, 熔化长英质物质, 混合形成花岗岩。

2.1 俯冲作用

花岗岩可以形成于俯冲作用的多个阶段, 但早期形成的花岗岩并不具混合特征^[10]。在俯冲过程中, 洋壳脱水, 但不会部分熔融^[11]。当其俯冲到壳幔边界, 上地幔部分熔融便形成玄武质岩浆。玄武质岩浆和流体绝大部分底侵于下地壳, 加热岩石圈, 诱发前寒武纪变质岩的部分熔融, 产生的中酸性岩浆侵入地壳后形成花岗岩^[12]。

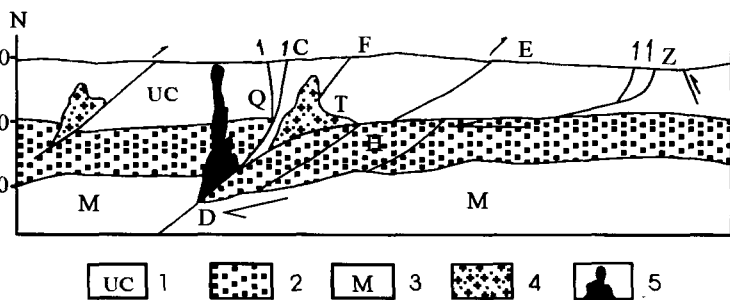
大陆内部挤压俯冲的典型例子是东秦岭燕山期 A型花岗岩^[13]。其形成过程是: 来自地幔环境的高温中-基性岩浆加入地壳, 与处于特定温度、压力条件下的陆壳岩石(太华群基底为主)相互作用形成花岗质岩浆(图 1)。

随着下插进入地幔, 俯冲盘逐渐开始含水分熔, 形成的含水岩浆进入地幔。若地幔岩处于 1000~1100℃条件, 来自消减地壳的含水岩浆加入将诱发地幔岩分熔形成高温中-基性岩浆, 产生同熔系列的初始岩浆, 随着同熔系列初始岩浆上升进入下地壳, 又可进一步诱发基底地层的分熔, 形成花岗闪长质岩浆。两种岩浆混合后进一步上升, 促使中上地壳内的岩石分熔形成花岗质熔体。

2.2 底侵作用

花岗岩质岩浆的形成过程中, 幔源岩浆的注入, 无论是长英质地壳岩浆和镁铁质幔源岩浆的混合, 还是由岩浆底侵混合形成的地壳岩石的一致熔融^[14], 都起着重要的作用。镁铁质岩是地幔岩浆的来源; 长英质岩是地壳岩石圈的主要组成部分, 一般都含有地幔物质。

混熔假说认为源岩的深熔与镁铁质岩浆有关: 镁铁质岩浆携带的热和热液使上覆地壳重熔形成花岗质熔浆, 继之而来的镁铁质岩浆上升, 进入花岗质岩浆内, 可产生混熔。太古代地壳以下的岩浆源地壳是以镁铁质为主要的下地壳。富轻稀土的大陆岩石圈地幔的部分熔融, 伴有轻微的地壳同化作用, 可产生镁铁质岩浆。接着, 岩浆上涌熔融下地壳, 形成奥长环斑花岗岩浆^[15], 如怀俄明州的奥长环斑花岗岩浆就是拉斑玄武岩(镁铁质)-中性火成岩熔融的产物^[16]。其中的地壳同化物极有可能是中性-长英质变质火成岩^[17,18]或镁铁质-中性变质火成岩^[16,19], 形成于早期的岩浆底侵作用。Sr、Nd、Pb 同位素研究



1. 上地壳; 2. 下地壳; 3. 地幔; 4. 改造系列花岗岩; 5. 同熔系列花岗岩;
ZHTD. 主滑脱带; CQD. 超壳断裂; EH, FT 为次级滑脱带; 据胡志宏等, 1990

图 1 东秦岭北部燕山期陆内挤压-俯冲示意图

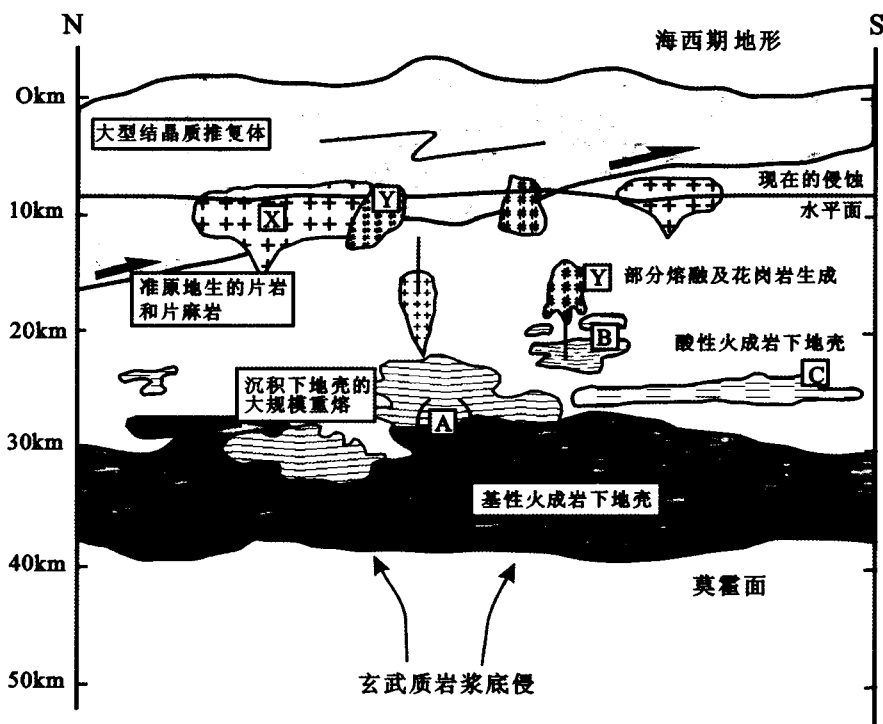
Fig. 1 A cross section showing the Yashanian intracontinental compression-subduction zone in the northern part of the eastern Qinling area

亦表明,花岗岩很大程度上来源于幔源岩浆底侵下地壳形成的混合源岩的熔融^[20]。

形成于拉张环境的花岗岩多见有玄武质岩石包体,并见有壳幔岩浆混合作用特征,表明大陆内部存在玄武质岩浆的底侵作用^[21]。另外,火山在喷发前的岩浆阶段也会产生对下地壳的底侵。法国中央高地花岗闪长岩和二长花岗岩就起源于底侵作用下基性变质火成岩和下地壳变质沉积岩的混合熔融^[22]。在混合过程中,亏损地幔成分注入镁铁质地壳和(或)为其提供热源^[23]。而且,变质岩系列源

岩的此种混合也可产生强过铝质花岗岩,如 Guéret 和 Millevaches 的花岗二长岩和闪长花岗岩^[24]。

图2为海西晚期法国中央高地的理想剖面。Downes等^[25]认为,此处基性变质火成岩下地壳形成于海西造山期玄武质岩浆的底侵作用。因此,在地壳深部,母岩浆可能与变质沉积下地壳、重熔的酸性变质岩混合。底侵岩浆不断受到地壳混染,并与变质沉积下地壳熔体混合形成花岗闪长岩和二长花岗岩。来自底侵作用的热补给也导致变质下地壳的大规模重熔,产生强过铝质花岗岩岩浆^[26]。



X. 花岗闪长岩和二长花岗岩; A. 基性底板; B. 变质沉积下地壳; C. 重熔的酸性变质火成岩; Y. 浅色花岗岩; 据 Downes 等, 1990

图2 325 Ma 法国中央高地剖面示意图

Fig. 2 Schematic cross-section through Massif Central at 325 Ma

2.3 拆沉作用与地幔柱

造山后期阶段的陆陆碰撞会造成下地壳和顶部地幔的剥离^[27];岩石圈增厚也可导致镁铁质突起拆离。据此,下地壳岩石圈重新回到地幔,极可能交代周围地幔。如果重新被激活,该地幔区就会有地壳混染的记录。

研究表明,地幔柱上升至岩石圈时会有地壳物质的加入^[28,29]。来自地幔的高热物质呈蘑菇云状上涌,导致大规模花岗岩岩浆活动,形成大面积花岗岩省和大火成岩省^[30]。

3 壳幔相互作用特征

经历过壳幔相互作用形成的花岗岩体多呈岩株、岩脉产出,有极少量呈不规则状或不规则条带状残留岩基产出,其余则呈大小不一的孤岛状零星分布^[31]。但其他成因花岗岩也有可能具这些特征,不能体现其独特性。

此类岩体内部多有暗色微粒包体^[32~38],以闪长质为主,具岩浆成因细粒半自形粒状结构,块状构造^[32,34~37]。研究表明,暗色微粒包体可作为岩浆

混合的证据,暗示壳幔岩浆的相互作用^[39,40]。花岗岩体一般与围岩呈突变接触,接触面倾向围岩,局部与围岩接触处粒度变细,围岩也有不同程度的热变质,具球状风化特征^[32,33,37]。暗色微粒包体与花岗岩呈弥漫型或截然型接触,前者接触部位暗色矿物增多,形成宽约几厘米到几十厘米的岩浆混合过渡带,后者局部见淬冷边。

4 混合的地球化学制约

花岗岩成分及其矿物、地球化学参数是花岗岩地质学的基础。其组分的细微差别可反映花岗岩多样的演化途径,并反映全面的地质环境变迁^[41]。

(1) Eu 的负异常:Rogers 等^[42] 和 Ellam 等^[43] 指出,轻微负 Eu 异常反映地幔源区有地壳来源成分(作为俯冲沉积物)的加入。在沉积成因岩石和地幔端员混合曲线图^[44] 中,有地幔物质参与的相对弱演化火成岩的 $\text{Eu}/\text{Eu}^* -^{143}\text{Nd}/^{144}\text{Nd}$ 值分布在这两种端员的范围内。

(2) HFSE:不同的 Th/Ta 值可区分不同的火成岩区。尽管 Th/Ta 值可因分异作用而变化,但也能直接反映不同的物质来源。若岩浆的 Th/Ta 值在 Th/Ta-Zr 图中发生重叠,则表明它们来自相同的源区,未重叠的部分则有其他不同成分的存在。地壳物质的加入量不同, Th/Ta-Zr 值就会变化。另外,较低的 Nb/Ta 值间接表明地壳混染,或地壳熔融,或(和)发生过熔化事件^[44]。

(3) Sr、Nd 同位素: $^{87}\text{Sr}/^{86}\text{Sr}$ 和 $^{87}\text{Sr}/^{86}\text{Sr}-1/\text{Sr}$ 图解可用于研究花岗岩类的壳幔物质混合。有些学者^[45] 以假定源区的 Sr、Nd 同位素成分来解释:拉斑玄武岩和中钾玄武岩的 $^{87}\text{Sr}/^{86}\text{Sr} = 0.705 \sim 0.709$, $^{143}\text{Nd}/^{144}\text{Nd} = 0.5125 \sim 0.5123$; 上地壳的 $^{87}\text{Sr}/^{86}\text{Sr} > 0.750$, $^{143}\text{Nd}/^{144}\text{Nd} < 0.5118$; 下地壳的 $^{87}\text{Sr}/^{86}\text{Sr} \approx 0.710$, $^{143}\text{Nd}/^{144}\text{Nd} < 0.5118$ 。

在初始 Nd 同位素成分的基础上,大陆地壳一般可以分成地幔来源的新生(juvenile)地壳和至少有部分古老地壳来源的进化(evolved)地壳。前者具有正 ϵ_{Nd} 值,类似亏损地幔来源;后者具有负 ϵ_{Nd} 值,类似古老地壳来源^[46]。

5 总 结

地幔物质以俯冲、拆沉、底侵、地幔柱、火山喷发等方式与地壳物质混合形成花岗质岩浆,使其具有混合特征。此类花岗岩多含有具岩浆成因细粒半自

形粒状结构的暗色微粒包体,能够体现壳幔岩浆的相互作用。稀土元素、HFSE、同位素特征可将地壳物质与地幔物质区分开来,限定花岗岩的物质来源范围,确定其来源是单一的还是混合的。

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