

# The spatial distribution and emission of nitrous oxide (N<sub>2</sub>O) in a large eutrophic lake in eastern China: Anthropogenic effects

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## ABSTRACT

The emission of N<sub>2</sub>O from China is globally significant, but relatively few direct observations have been made in many of the fresh water environments most likely to be important sites of N<sub>2</sub>O production. In this paper, N<sub>2</sub>O saturations were examined in the ecologically heterogeneous, eutrophied, Lake Taihu, as well as in surrounding rivers in eastern China. The emissions of N<sub>2</sub>O were estimated and compared with those from other landscapes within the Lake Taihu drainage basin. We found that anthropogenically-enhanced inorganic N inputs act as a limited primary control on the spatial distribution of N<sub>2</sub>O saturations in heavily eutrophied parts of the lake only and that overall, lake N<sub>2</sub>O production and emission are not raised as significantly as expected due to high N inputs. In comparison, the heavily eutrophied river network is an important fraction of the local N<sub>2</sub>O budget, and when considered together with emissions of N<sub>2</sub>O from the lake, constitute a major (10–50% depending on season) fraction of total N<sub>2</sub>O emissions from the Lake Taihu drainage basin.

## 1. Introduction

Based on long-term records from ice cores, it is known that global concentrations of atmospheric N<sub>2</sub>O began to rapidly increase during the last century, and that this trend continues, with anthropogenic sources estimated to contribute approximately 1/3 of the total (Khalil and Rasmussen, 1992; Seitzinger et al., 2000). While it is postulated that a global increase in the frequency of occurrence and extents of hypoxic zones in continental shelf and estuarine settings may seriously impact the global N<sub>2</sub>O budget (de Wilde and de Bie, 2000; Naqvi et al., 2000), the potential effects of anthropogenic nutrient inputs on N<sub>2</sub>O production in lakes and rivers is poorly understood.

Aquatic  $N_2O$  production is complex and sensitive to a variety of processes and variables, as evidenced by the follow-

ing examples. Increased inorganic nitrogen  $(NH_4^++NO_3^-)$  concentrations can promote nitrification in oxic environments and denitrification in anoxic environments to enhance N<sub>2</sub>O production (Seitzinger, 1985; McMahon and Dennehy, 1999; Cole and Caraco, 2001). Proliferation of various genera of green algae (Chlorella, Scenedesmus, Coelastrum, and Chlorococcum) in response to eutrophication, can significantly enhance N<sub>2</sub>O production (Weathers, 1984). By increasing the supply of organic matter, eutrophication can drive hypoxia or anoxia in bottom waters, and stimulate denitrification (Naqvi et al., 2000; Liikanen and Martikainen, 2003). Enhanced denitrification and the addition of anthropogenic NO<sub>3</sub> lead to significant accumulations of N<sub>2</sub>O (Herbert, 1999; Cole and Caraco, 2001).

The importance of lacustrine systems in emitting or absorbing  $N_2O$  is poorly constrained. Generally, lakes are not

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Fig. 1–Map of Lake Taihu, China. The dashed lines delineate the boundaries of the three ecological zones, and the open and solid circles delineate the locations of sample sites in July and September, 2003, respectively. The sampling transects  $M-M_1$  and  $G-G_1$  are also shown.

considered significant sources of N<sub>2</sub>O to the atmosphere (Mengis et al., 1997; Seitzinger and Kroeze, 1998), and in some cases have been shown to serve as minor or seasonal sinks (Hendzel et al., 2005). However, higher average N<sub>2</sub>O saturations have been observed in surface waters of eutrophic as compared to oligo- and meso-trophic lakes (Mengis et al., 1997; Huttunen et al., 2003). Lakes are important sinks for watershed-derived N (e.g., Gulati and van Donk, 2002; Wang et al., 2006; Bunting et al., 2007), the extent to which lacustrine N<sub>2</sub>O saturation is enhanced by increasing N levels and expanding eutrophication is not well characterized.

In the aquatic environment, nitrous oxide is mainly produced by two biogenic mechanisms. The first is reductive, microbial production during denitrification (Knowles et al., 1981), and the second is oxidative, microbial production during nitrification (Yoh et al., 1988). As an intermediate in denitrification (NO<sub>3</sub> to N<sub>2</sub>), N<sub>2</sub>O may accumulate when O<sub>2</sub> is present, NO<sub>3</sub> concentrations are high and pH is low (Knowles, 1996). The rate of denitrification is influenced by temperature, abundance of organic carbon and supply of N (Knowles, 1996); all save N supply are interrelated with  $O_2$  concentrations. As such, it is not surprising that incubation experiments show that N<sub>2</sub>O dynamics in eutrophic lakes are regulated by the availability of O2 (Liikanen and Martikainen, 2003). In nitrification ( $NH_4^+$  to  $NO_2^-$ ),  $N_2O$  can be produced as a by-product when NH<sup>+</sup><sub>4</sub> is oxidized by nitrifying bacteria and methanotrophs (Mengis et al., 1997). Methane oxidation can interfere with NH<sub>4</sub><sup>+</sup> oxidation due to competition for O<sub>2</sub> (Liikanen and Martikainen, 2003). Methane and N<sub>2</sub>O concentrations have been compared to explore the processes governing N<sub>2</sub>O production in fresh waters (Mengis et al., 1997).

Calculated  $N_2O$  emissions from China are globally significant, accounting for over 90% of those in the Pacific Basin (Seitzinger and Kroeze, 1998). Despite this, few direct observations of  $N_2O$  production and emission have been made in the many impacted watersheds, lakes and estuaries of China (Xing et al., 2001; Wang et al., 2006). This shortage of data in key areas is a major contributor to the uncertainty of modeling results (Seitzinger and Kroeze 1998; Seitzinger et al., 2000). In this study, we investigated the spatial patterns of  $N_2O$ saturation in a large, eutrophied lake and its surrounding rivers in eastern China. The objectives of this study were to address the following hypotheses: (H<sub>1</sub>) variable nutrient inputs and degrees of eutrophication act as the primary control on the distribution of  $N_2O$  saturations and emissions throughout the lake, and (H<sub>2</sub>) when considered together, lake and river  $N_2O$  emissions constitute a considerable fraction of the total  $N_2O$  emissions from the Lake Taihu drainage basin.

## 2. Study site

Lake Taihu is the third largest lake in China with a surface area of approximately 2338 km<sup>2</sup>, an average water depth of 1.8 m and a volume of approximately  $5.77 \times 10^9$  m<sup>3</sup>. It is located in a heavily polluted region in eastern China (Fig. 1). A population of more than 37 million and a significant industrial complex are located within its drainage basin (~36,500 km<sup>2</sup>). Cultivated lands comprise approximately 15,100 km<sup>2</sup> of the drainage area and are heavily fertilized annually (~34.5 g N m<sup>-2</sup> a<sup>-1</sup>) (Xing et al., 2001; Qin and Luo, 2004). The lake annually receives approximately 30,635,000 kg total nitrogen (TN), 1,751,000 kg total phosphorus (TP) and 131,223,000 kg chemical oxygen demand on chromium (COD<sub>Cr</sub>) from a combination of municipal and industrial wastewaters and agricultural soil runoff (Qin and Luo, 2004). Consequently, the lake is eutrophied, particularly in the northern area. Lowlands comprise more

Table 1 – Summary total nitrogen, total phosphorus, dominant vegetation types,  $N_2O$  saturation,  $CH_4$  saturation and the estimated  $N_2O$  exchange fluxes in Lake Taihu and adjacent rivers.

Area	Vegetation	$TN^{a}$ (mg L <sup>-1</sup> )	TP <sup>a</sup> (mg L <sup>-1</sup> )	CH <sub>4</sub> saturation <sup>b</sup> (%)		N <sub>2</sub> O saturation <sup>b</sup> (%)		N <sub>2</sub> O fluxes <sup>b</sup> (μmol m <sup>-2</sup> d <sup>-1</sup> )	
				Jul.	Sept.	Jul.	Sept.	Jul.	Sept.
North L. Taihu	Algae	3.52	0.13	8746±9336 (98.0–757.5)	2616±2552 (39.9–332.7)	689±472 (21.9–98.8)	331±189 (11.5–46.3)	54.9±63.5	24.8±22.0
Middle L. Taihu	Algae/macrophyte	2.0	0.07	783±781 (1.5–64.7)	1330±1318 (3.5–115.3)	184±44 (6.6–16.8)	176±38 (9.4–21.1)	8.3±4.4	7.3±3.1
South L. Taihu	Macrophyte	0.77	0.03	3554±1731 (44.5–127.3)	1579±1350 (12.7–165.0)	80±10 (4.6–7.9)	144±12 (9.8–14.0)	$-1.7 \pm 0.9$	$5.0 \pm 1.4$
Lake averages <sup>c</sup> Rivers		2.08 4.27	0.07 0.16	2556 20,157±15,054 (17.1–1588)	1581 3148±3048 (25.7–286.1)	246 1501±744 (10.5–169.2)	195 342±223 (9.8–62.2)	14.0 142.1±72.9	9.7 28.8±26.4

<sup>a</sup> Data are from Zhai and Zhang (2006).

 $^{\rm b}\,$  Given are means  $\pm\,$  SE, concentration range (nmol L^-1) is shown in parentheses.

<sup>c</sup> Averages are weighted to surface areas.

than 80% of the drainage basin. A dense river network, with a total distributary length of approximately 120,000 km, covers 7% of the total drainage area. Water pollution is more severe in some of these rivers as compared to the lake, particularly those which flow into the lake, as evidenced by higher concentrations of TN and TP (Table 1) (Qin and Luo, 2004; Zhai and Zhang, 2006).

Spatial contrasts in ecology, vegetation and nutrient levels are evident, allowing Lake Taihu to be divided into three zones (Fig. 1): the hypereutrophic-eutrophic northern zone where pollution (TP, TN and COD) is most serious (Qin and Luo, 2004; Table 1), vegetation is dominated by large blue-green algae blooms (cyanobacteria) and sediments consist mainly of organic and nutrient-rich sludge; the middle zone where TN and TP concentrations and the degree of eutrophication are intermediate between the northern and southern zones (Table 1), vegetation consists of a mix of algae and macrophytes and sediments are sandy, with relatively little organic matter; and the mesotrophic southern zone where pollution is relatively light, concentrations of TN and TP are the lowest (Table 1), vegetation is dominated by submersed and emergent macrophytes and sediments are a thick, poorly consolidated and organic-rich slurry. These spatial heterogeneities provide an excellent setting to compare the importance of various factors on aquatic N<sub>2</sub>O production and emissions.

## 3. Materials and methods

Wind-driven mixing provides for a well mixed water column in this very shallow lake (McCarthy et al., 2007), which is confirmed by a high observed vertical velocity range (1.5 to 2 cm s<sup>-1</sup>) (Luo et al., 2004). Thus, concentrations of N<sub>2</sub>O in surface waters are nearly equivalent to those in bottom waters (Wang et al., 2008). As a result, only intermediate-depth lake waters (0.7–0.9 m depth) were sampled, and their N<sub>2</sub>O concentrations are considered representative of conditions at each site. Lake waters were sampled at 18 sites, encompassing all three lake zones, in July, 2003. Additional sampling sites were added in September, 2003 with the intention of profiling N<sub>2</sub>O concentrations along environmental gradients from both east to west and north to south (Fig. 1, transects  $G-G_1$  and  $M-M_1$ , respectively). Additionally, 19 major rivers flowing into the lake were sampled in both July and September, 2003 (Fig. 1).

Field measurements included total dissolved solids (TDS) and pH, which were determined using a PIONneer 65 multiparameter instrument (Radiometer Co.); and dissolved oxygen (DO), which was determined using a portable DO analyzer (HANNA Co.). Small water samples were collected at each station using serum bottles for the laboratory determination of NO<sub>3</sub> and NO<sub>2</sub> using an HPLC1100 liquid chromatograph; and for NH<sub>4</sub><sup>+</sup>, which was determined by a spectrophotometer (Unico 2000). Water samples for gas measurements were collected in serum bottles, followed by the addition of 10 M NaOH to raise the pH (10<sup>+</sup>), after which the bottles were sealed with no air bubbles. At the laboratory, the headspace equilibrium technique was used to determine concentrations of dissolved gases (Mengis et al., 1997; Wang et al., 2008). Approximately 20 ml of ultra-pure N2 was injected into the sample bottle and water was displaced. After equilibration at 25 °C by vigorously shaking for 30 min in a water bath, the N<sub>2</sub>O concentration in the headspace was analyzed using an ECD-GC (HP6890) equipped with a 4.5 m×3 mm packed Porapak Q (80/100 mesh) column. The column and ECD detector were conditioned at 50 °C and 320 °C, respectively. A mixture of Ar/CH<sub>4</sub> (95/5 v/v) was used as a carrier gas at a flow rate of 20 ml min<sup>-1</sup>. Headspace CH<sub>4</sub> concentrations were determined using an FID-GC (HP6890) equipped with a 2 m×3 mm carbon molecular sieve column (60/80 mesh). The column and FID detector were conditioned at 90 °C and 250 °C, respectively. Ultra-pure hydrogen gas was used as a carrier gas at a flow rate of 20 ml min<sup>-1</sup>.

From headspace concentrations, the formula described by Butler and Elkins (1991) was used to calculate gas concentrations, with a mean error of  $\pm 4\%$ . Gas concentrations are expressed by the saturation degree relative to air (Mengis et al., 1997):

 $N_2O$  saturation =  $(C_{N_2O}/C_{N_2Oatm.}) \times 100$ 

where  $C_{N2O}$  is the measured concentration of  $N_2O$ , and  $C_{N2Oatm.}$  is the saturation concentration of atmospheric  $N_2O$  at the given



Fig. 2–N<sub>2</sub>O versus NH<sup>4</sup><sub>4</sub> (A, B) and NO<sup>3</sup><sub>3</sub> (C, D) in Lake Taihu (September, 2003), and N<sub>2</sub>O versus NH<sup>4</sup><sub>4</sub> (E, F) (July, 2003). The solid circles denote samples from the eutrophied northern zone and the open circles from the less eutrophied middle and southern zones.

water temperature. Expression of the  $CH_4$  saturation degree is analogous to that of  $N_2O$ . The exchange flux of  $N_2O$  at the waterair interface is calculated using:

 $F = K \Delta C$ 

where F is the gas exchange flux,  $\Delta C$  is the difference between N<sub>2</sub>O concentrations in the air and water, and K is the gas transfer velocity (calculated from the wind speed and Schmidt number (Wanninkhof, 1992; de Wilde and de Bie, 2000)).

## 4. Results and discussion

#### 4.1. N<sub>2</sub>O production

The rate of increase of  $N_2O$  relative to  $NH_4^+$  that can be estimated as  $N_2O$  saturations was found to correlate linearly with  $NH_4^+$  concentrations. Regressing the September data for

the entire lake yielded a slope of  $0.08 \times 10^{-2}$ ; in comparison, the slope was  $0.11 \times 10^{-2}$  using only those data collected in the eutrophied northern zone (Fig. 2A). These relationships suggest that approximately 1 mol of N<sub>2</sub>O is formed during nitrification for every 900–1250 mol of NH<sup>+</sup><sub>4</sub> in a given volume during September in Lake Taihu. As a proportion of the NH<sub>4</sub> concentration, these data are within the range of N<sub>2</sub>O yields reported from the Scheldt Estuary by de Wilde and de Bie (2000) (0.04% to 0.42% (mol N<sub>2</sub>O per mol NH<sub>4</sub><sup>+</sup>)), and are slightly lower than the average N<sub>2</sub>O yield of 0.14% (1 mol N<sub>2</sub>O for 700 mol NH<sub>4</sub>) during nitrification in both fresh and salt water areas (de Wilde and de Bie, 2000). The same regression using the July data, for the entire lake resulted in a slope of  $0.05 \times 10^{-2}$  (Fig. 2E), which is presumably a result of the origin of N<sub>2</sub>O from denitrification. Despite strong positive relationships between N<sub>2</sub>O and both  $NH_4^+$  and  $NO_3^-$  in the lakes eutrophied northern zone (Fig. 2A, C, E), large increases in  $NH_4^+$  and  $NO_3^-$  resulted in only modest increases in N<sub>2</sub>O, particularly for the September data set.



Fig.  $3-N_2O$  saturation relative to TDS concentration in the heavily eutrophied northern zone (solid circles) and less eutrophied middle and southern zones (open diamond) during July (A) and September (B), 2003.

#### 4.2. N<sub>2</sub>O saturations and inorganic N

Strong, positive correlations of N<sub>2</sub>O saturations with NO<sub>3</sub> (r=0.85, p<0.01) and NH<sub>4</sub><sup>4</sup> concentrations (r=0.74, p<0.01) were observed in the September data set (with the sole exception at the site in Wulihu Bay). In the July data set, N<sub>2</sub>O saturations did not significantly correlate with NO<sub>3</sub>, but did with NH<sub>4</sub><sup>4</sup> concentrations (r=0.94, p<0.01). Overall, N<sub>2</sub>O saturations and inorganic nitrogen concentrations were most strongly associated in the heavily eutrophied northern zone of the lake (Fig. 2A, C and E), not in the less eutrophied middle and southern zones (Fig. 2B, D and F). Additionally, both the

September and July data sets show that N<sub>2</sub>O saturations are strongly related to TDS concentrations over a wide range (~190–220 mg l<sup>-1</sup> in July; ~350–440 mg l<sup>-1</sup> in September) in only the northern zone of the lake (Fig. 3). The exponential increase of N<sub>2</sub>O saturations relative to TDS concentrations is consistent with enhanced N<sub>2</sub>O production by increased N loading in the eutrophied northern zone of the lake.

#### 4.3. Distribution of N<sub>2</sub>O and CH<sub>4</sub> saturations

Significant spatial differences in N2O saturation were observed in Lake Taihu's three zones (Fig. 4A, Table 1), ranging from 161% to 1579% in the northern zone, 140% to 267% in the middle zone and 70% to 164% in the southern zone. Maximum saturations of N<sub>2</sub>O, 1597% (July) and 748% (September) were observed in Wulihu Bay (Figs. 4A and 5A) where nutrient pollution is most serious. In both the northern and middle zones, N2O was supersaturated with respect to the atmosphere during both July and September, 2003. In the macrophyte-dominated southern zone, lake water was not saturated with respect to atmospheric N<sub>2</sub>O during the growing season (July) and thus acts seasonally as a N<sub>2</sub>O sink (Table 1). From north to south, N<sub>2</sub>O saturations decrease along the M-M<sub>1</sub> transect (Wulihu to Meiliangwan Bays) (Fig. 5A) and are spatially consistent with concentrations of TN and TP, and the degrees of lake eutrophication during July and September (Table 1).

As previously noted, significant ecological contrasts were observed between the three lake zones, these are considered here as possible factors influencing the spatial distribution of N<sub>2</sub>O. Overall, algal blooms dominate in the lake's northern zone, corresponding to high N<sub>2</sub>O concentrations, while the abundant macrophytes of the southern zone correspond to low N<sub>2</sub>O concentrations (Table 1). Generally, floating algae mats are not major sites for denitrification (Schaller et al., 2004), however, high N<sub>2</sub>O concentrations and fluxes during algal blooms can result indirectly from the supply of dissolved and particulate organic carbon provided by the algae, which can stimulate denitrification (Harrison and Matson, 2003). In



Fig. 4-Variations in saturation degree (%) of N<sub>2</sub>O (A) and CH<sub>4</sub> (B) (September 2003).



Fig. 5 – Variations in N<sub>2</sub>O (solid circles) and  $CH_4$  (open diamonds) saturations along transects  $M-M_1$  (A) and  $G-G_1$  (B).

the lake's northern zone, N<sub>2</sub>O saturation does not show any obvious response to the gradient from algae to macrophyte dominance along the G-G1 transect (Fig. 5B). In rivers, Chenier et al. (2006)showed a significant relationship between the denitrification rate and the biomass of algae and heterotrophic bacteria, but not with cyanobacteria. Cyanobacteriadominated algal blooms in the lake's northern zone are likely not an important driver of high N<sub>2</sub>O concentrations. The lack of N<sub>2</sub>O saturation during the growing period (July) in the macrophyte-dominated southern zone of the lake (Table 1) is mainly attributed to consumption by denitrifiers, which can reduce  $N_2O$  to  $N_2$  while  $NO_3^-$  concentrations remain low (Herbert, 1999; Harris, 1999). Denitrification can be stimulated by macrophytes, which can trap labile organic detritus in the water column or release it from roots (Herbert, 1999; Li et al., 2008). The observed unsaturated conditions lead to the hypothesis that macrophyte-dominated lakes with low nitrate concentrations may act seasonally as a sink for atmospheric N<sub>2</sub>O. This requires further research and may play an important role in accurately estimating how aquatic ecosystems influence regional and global N2O loadings. At Lake Taihu, vegetative variations alone are likely not a primary factor affecting the spatial distribution of N<sub>2</sub>O. The indirect effects of vegetation variability on N<sub>2</sub>O production, including regulating concentrations of various N species, the supply and nature of organic carbon, and the condition of sediment-water interface, are likely more significant.

Saturations of  $N_2O$  in the rivers around Lake Taihu were generally higher than those in the lake, coinciding with higher concentrations of TN and TP (Table 1). The maximum observed  $N_2O$  saturation (2708%) exceeded the maximum value (2500%) documented in the N-enriched South Platte River, U.S.A. (McMahon and Dennehy, 1999). Saturations of  $N_2O$  in the heavily eutrophied rivers in the northern and eastern areas ranged from 1655% to 2708%, and were generally much higher than those in the less eutrophied rivers of the western and southern areas (631% to 1077%).

Saturations of CH<sub>4</sub> were similar to those of N<sub>2</sub>O in the heavily eutrophied northern zone of the lake during July (r=0.97, p<0.01) and September, 2003 (r=0.75, p<0.01) (Fig. 4A, B). As with N<sub>2</sub>O, maximum CH<sub>4</sub> saturations were found in the heavily eutrophied Wulihu Bay; and also decreased from north to south along the M–M<sub>1</sub> transect (Wulihu to Meiliangwan Bays) (Fig. 5A). In the less eutrophied, middle zone of the lake, similarities in spatial distributions of CH<sub>4</sub> and N<sub>2</sub>O were weak in September (r=0.48, p=0.02), and absent in July, 2003, and were absent entirely in the macrophyte-dominated southern zone of the lake. The overall north to south decreasing trend for N<sub>2</sub>O saturations, as well as for the similarities in spatial distributions of N<sub>2</sub>O and CH<sub>4</sub>, likely reflects the weakening influences of human activities as evidenced by TN, TP and COD concentrations, in particular, the increasing distance from sewage outlets.

#### 4.4. Emission flux of N<sub>2</sub>O

Exchange fluxes of N<sub>2</sub>O were estimated in the three zones of the lake in July and September (Table 1). The heavily eutrophied northern zone accounts for 40–60% of all N<sub>2</sub>O emissions from the lake, despite occupying only 16% of the total surface area (Table 1). Surface area-weighted, average N<sub>2</sub>O emission fluxes from Lake Taihu are 14.0 and 9.7 µmol  $m^{-2} d^{-1}$  in July, and September, 2003, respectively (Table 1), which equates to an annual emission flux of approximately  $1.01 \times 10^7$  mol N<sub>2</sub>O from Lake Taihu (at 11.8 µmol m<sup>-2</sup> d<sup>-1</sup>). The estimated emission rate (11.8 µmol m<sup>-2</sup> d<sup>-1</sup>) is slightly higher than the median value (10.4 µmol m<sup>-2</sup> d<sup>-1</sup>) of N<sub>2</sub>O emission measured during 2003–2004 (Wang et al., 2006).

Approximately  $4.46 \times 10^7$  kg TN was discharged into the lake in 2002, of which approximately  $1.15 \times 10^7$  kg TN flowed out of the lake and  $3.31 \times 10^7$  kg TN was lost in the lake (Zhai and Zhang, 2006), most of which was emitted into the atmosphere as N<sub>2</sub> and N<sub>2</sub>O. Assuming the same lake N budget in 2003, ~0.63% of the total N discharged into the lake was released to the atmosphere as N<sub>2</sub>O. While there are no available N<sub>2</sub>O emission data from sewage-fed lakes for direct comparison, the ~0.63% value is less than the suggested default emission factor (N<sub>2</sub>O:TN) from rivers (EF5-r, 0.75%, Houghton et al., 1996 (IPCC)). Estimated N<sub>2</sub>O emission rates for the eutrophied rivers in the Lake Taihu drainage basin are 142.1 and 28.8 µmol m<sup>-2</sup> d<sup>-1</sup> in July and September, respectively, (Table 1). These rates are much higher than those from either the lake or the fertilized lands in the drainage basin (16.5 to 90.7 µmol m<sup>-2</sup> d<sup>-1</sup>) (Zheng et al., 1997).

If the mean emissions in Lake Taihu are extrapolated to the total surface area (~3160 km<sup>2</sup>) of all lake water bodies in the Lake Taihu drainage basin (~36,500 km<sup>2</sup>), approximately 0.49– $0.73 \times 10^6$  kg N<sub>2</sub>O are being released annually. In comparison, estimated, annual N<sub>2</sub>O emissions from cultivated lands in the drainage basin are ~12.22 × 10<sup>6</sup> kg N<sub>2</sub>O. No data are available on the N<sub>2</sub>O emission flux from the forested lands in this area. It is clear from these data that eutrophic lakes are small contributors to regional N<sub>2</sub>O emissions, despite enhancements resulting from increased nutrient loads. Applying the river N<sub>2</sub>O emission rate to the entire drainage basin river network, gives approximately 1.19–5.89×10<sup>6</sup> kg N<sub>2</sub>O emisted annually,

which is 10–50% of the  $N_2O$  emissions estimated from all cultivated lands within the drainage basin. This estimate is in congruence with other findings which indicate that river systems with high nutrient loads are important sources of anthropogenic  $N_2O$  to the atmosphere (McMahon and Dennehy, 1999; Seitzinger et al., 2000; Garnier et al., 2006).

#### 5. Conclusion

High levels of anthropogenic N inputs have a limited, positive effect on N<sub>2</sub>O production and emission only in the eutrophied northern area of the lake and overall, N<sub>2</sub>O production and emission are not raised as significantly as expected  $(H_1)$ . In comparison, the heavily eutrophied river network is an important fraction of the local N<sub>2</sub>O budget, and when considered together with emissions of N<sub>2</sub>O from the lake, constitute a major (10-50% depending on season) fraction of total  $N_2O$  emissions from the Lake Taihu drainage basin ( $H_2$ ). Despite the indication that plant type has little impact on N<sub>2</sub>O production within the northern zone of the lake (G-G1 transect, Fig. 5B), future research should focus on longer and higher resolution time series field measurements as well as mesocosm experiments which focus specifically on the role of various flora (algae, macrophytes) as direct or indirect modifiers of the production rate of N<sub>2</sub>O in lacustrine systems.

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